

Dynamical and Thermodynamic Controls on Smoke-Cloud Interactions over the Amazon

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Observational Evidence of Smoke-Cloud Interactions over Amazon

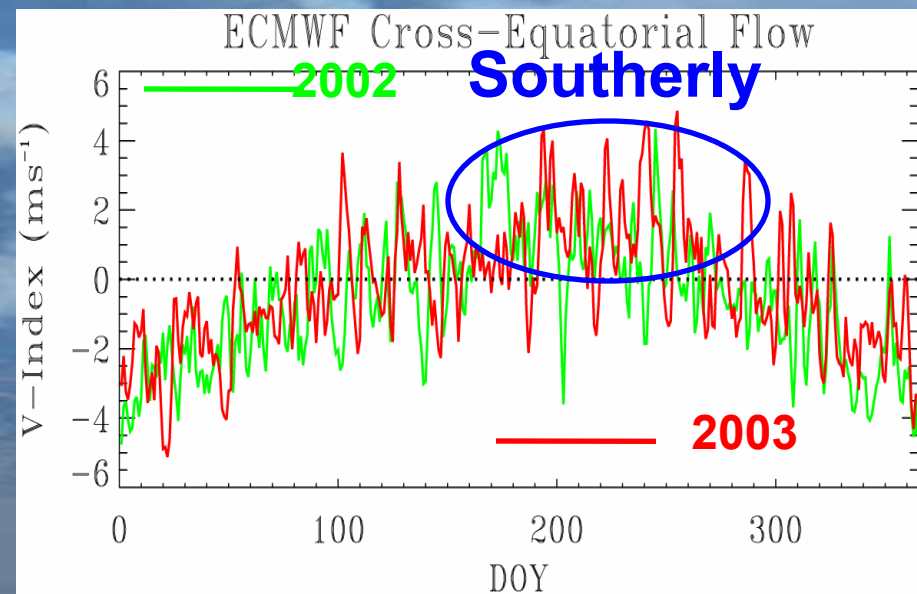
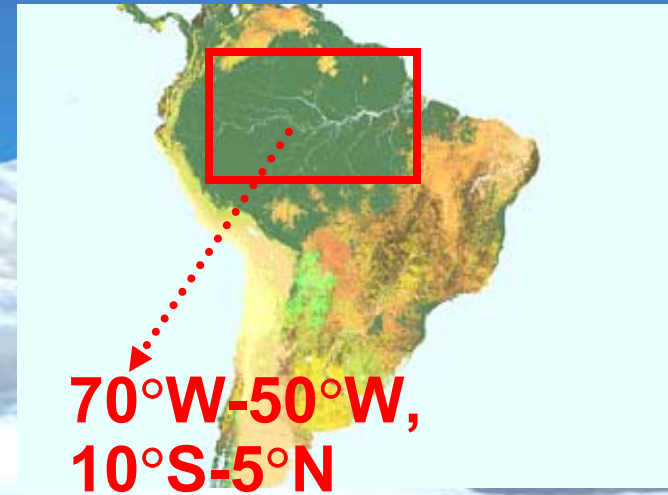
- Smoke particles act as CCN, resulting in highly concentrated but narrowly distributed cloud droplet spectra (larger reflectivity)
- Warm rain processes are suppressed, promoting vertical transport of water and pollutants (“smoking clouds”) and modifying circulations
- Smoke absorption both cools the surface and heats the atmosphere, changing the thermodynamics and likely reducing cloud cover (“cloud burning”)

Objectives:

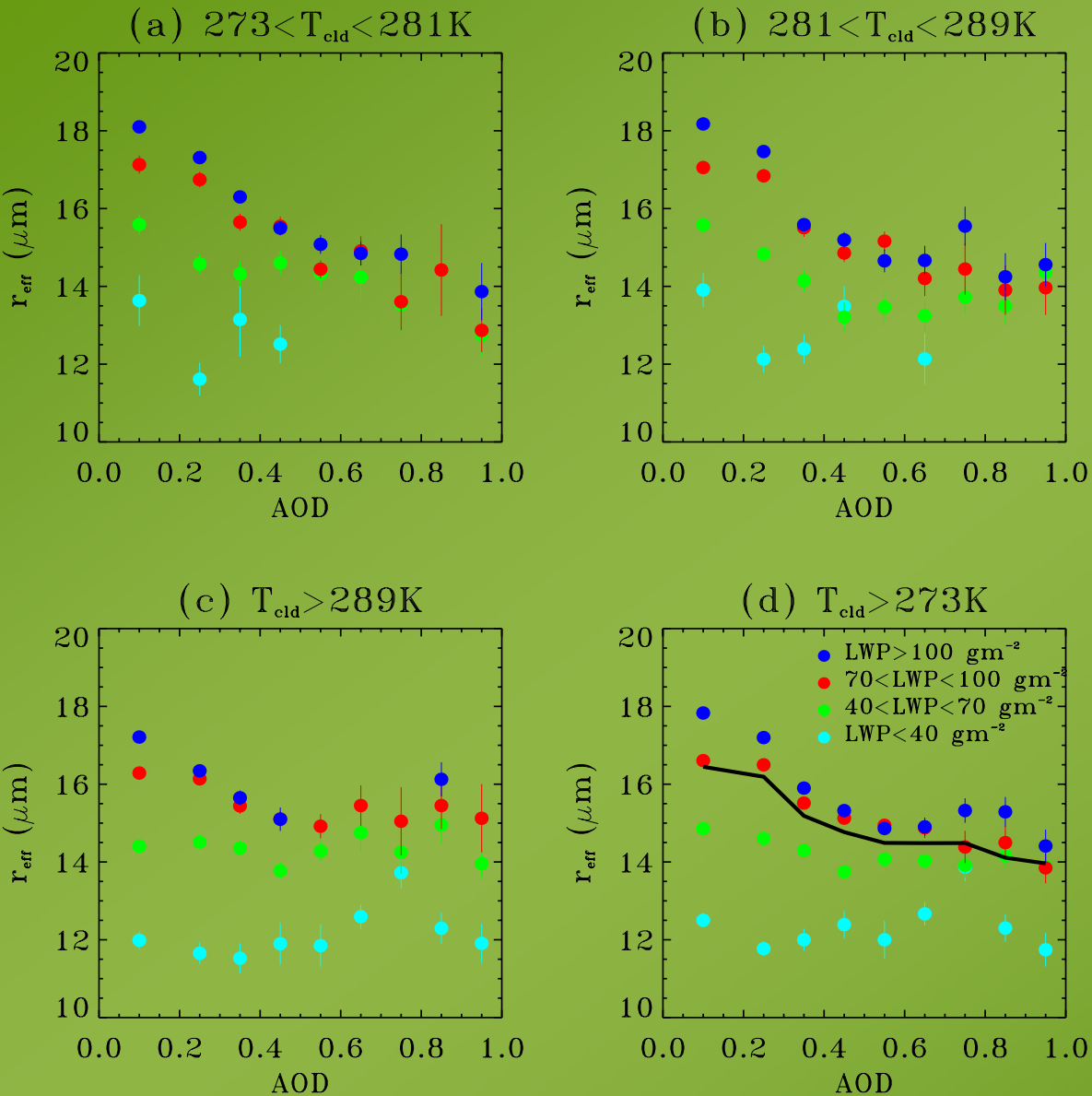
- To demonstrate the complexity of smoke-cloud interactions using MODIS data
- To explore
 - whether or not these complex interactions vary with the interannual changes of large-scale background;
 - what processes might control how changes of aerosol influence warm clouds.

MODIS Analysis

- MODIS/Aqua daily $1^\circ \times 1^\circ$
- ECMWF, radiosondes
- Dry-to-wet transition (*Aug-Oct*), tropical Amazon
- 2003 (*normal*) vs 2002 (*El Nino*)
- Warm clouds only
- Southerly cross-equatorial flow



2003: Normal Year

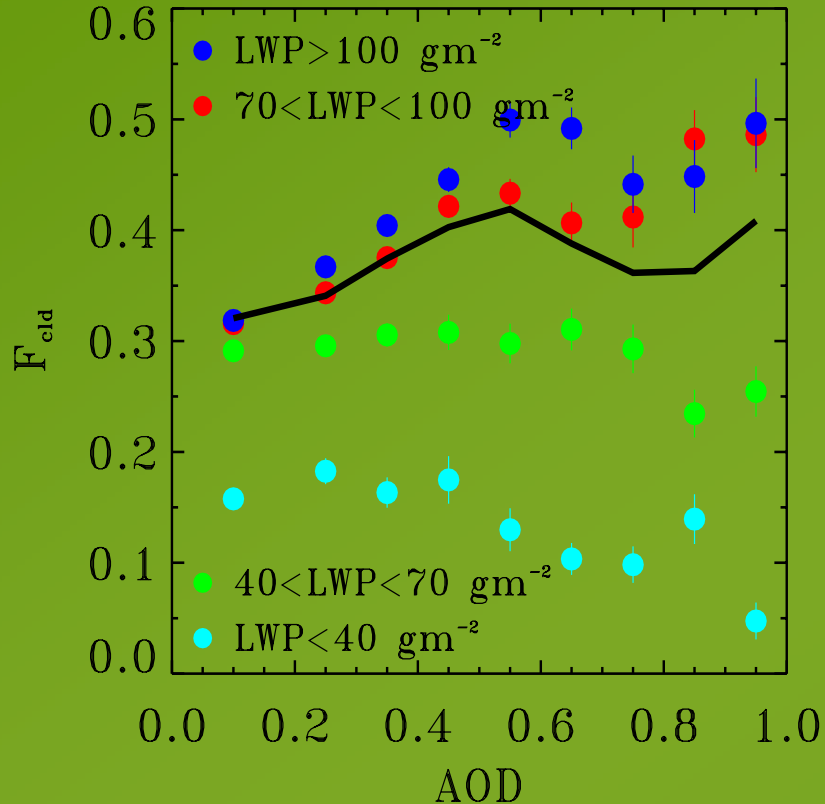


Twomey effect for
AOD < 0.6;

More significant at
high LWP than at
low LWP;

Droplet size levels
off at a high AOD.

2003: normal year



For $LWP > 70 \text{ gm}^{-2}$, cloud fraction increases with AOD (< 0.6)

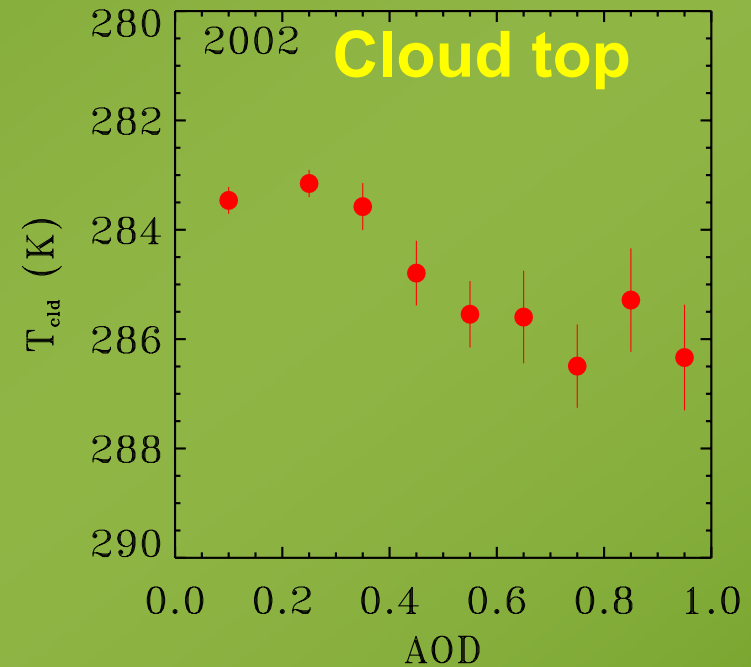
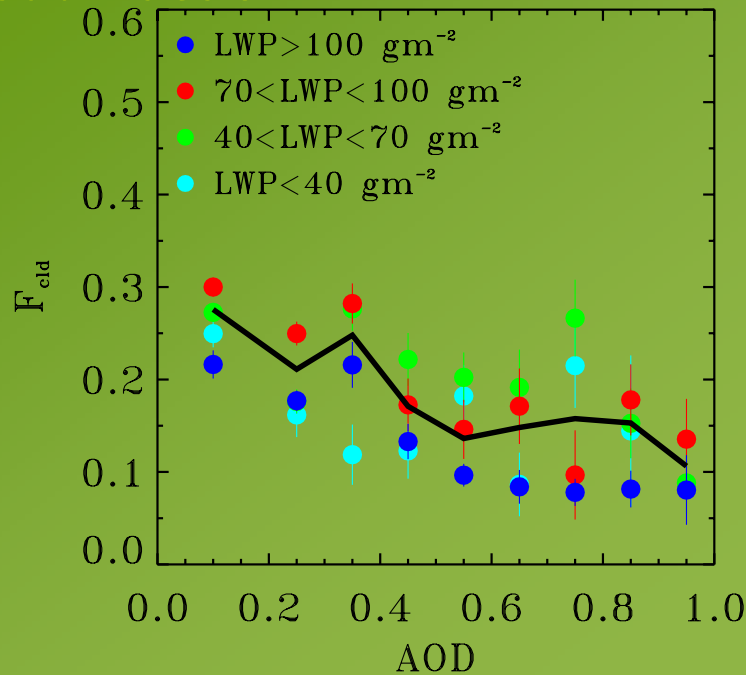
Thin clouds:

more or less constant fractions for $AOD < 0.6$

decrease for higher AOD conditions. (indication of cloud-burning effect)

2002 ---- El Nino year

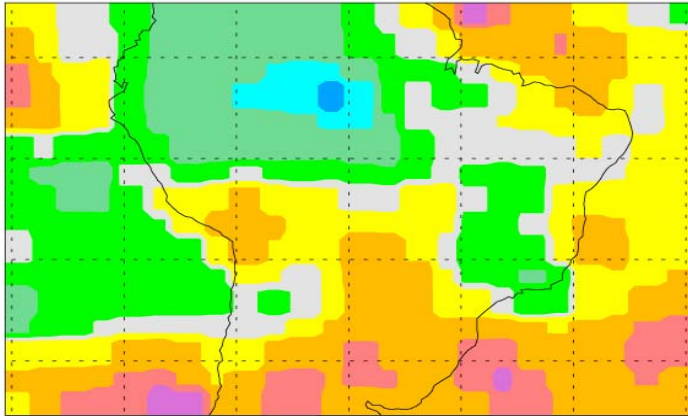
Cloud fraction (a) 2002



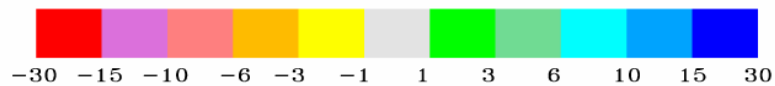
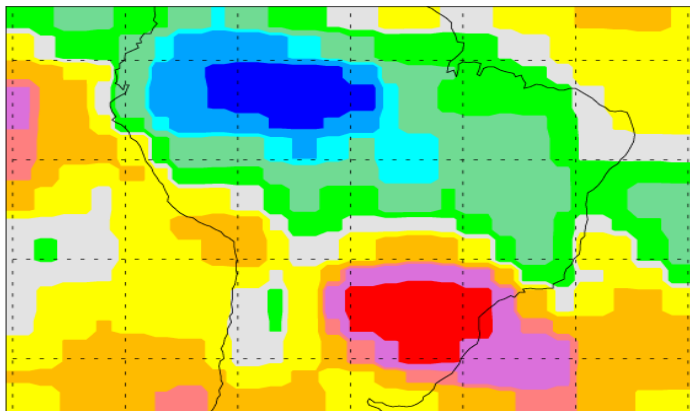
- Cloud fraction decreases with AOD, with greatest reduction for thicker and higher clouds. (semi-direct effect)
- The tops of warm clouds are apparently lowered by smoke
 - stabilization of ABL
 - enhanced capping inversion by smoke absorption

2003 was wetter than 2002

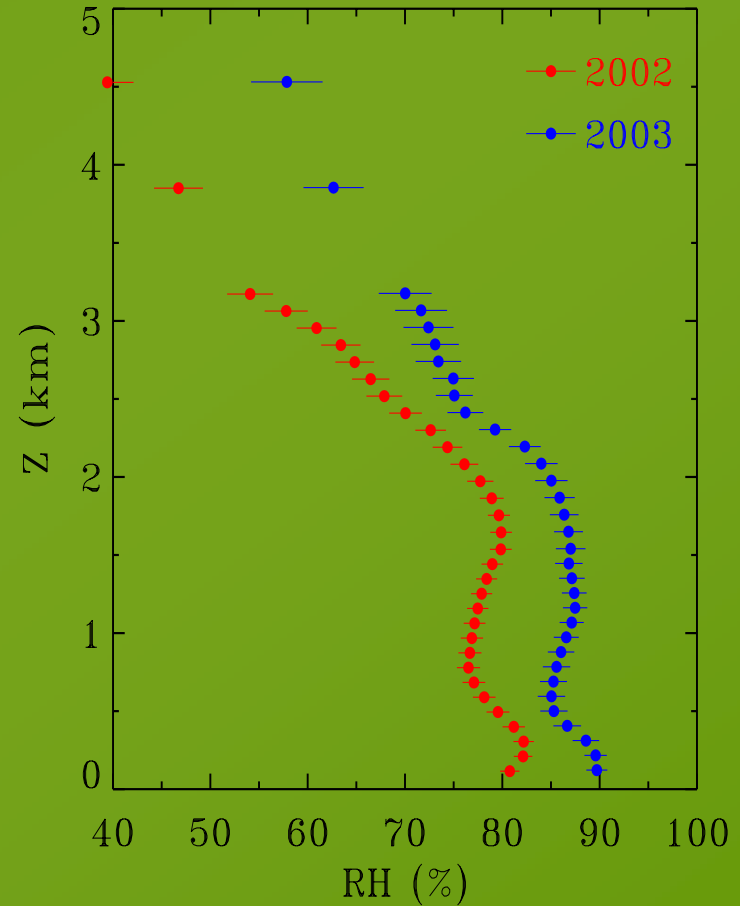
ECMWF 850 mb (2003-2002)



700 mb (2003-2002)

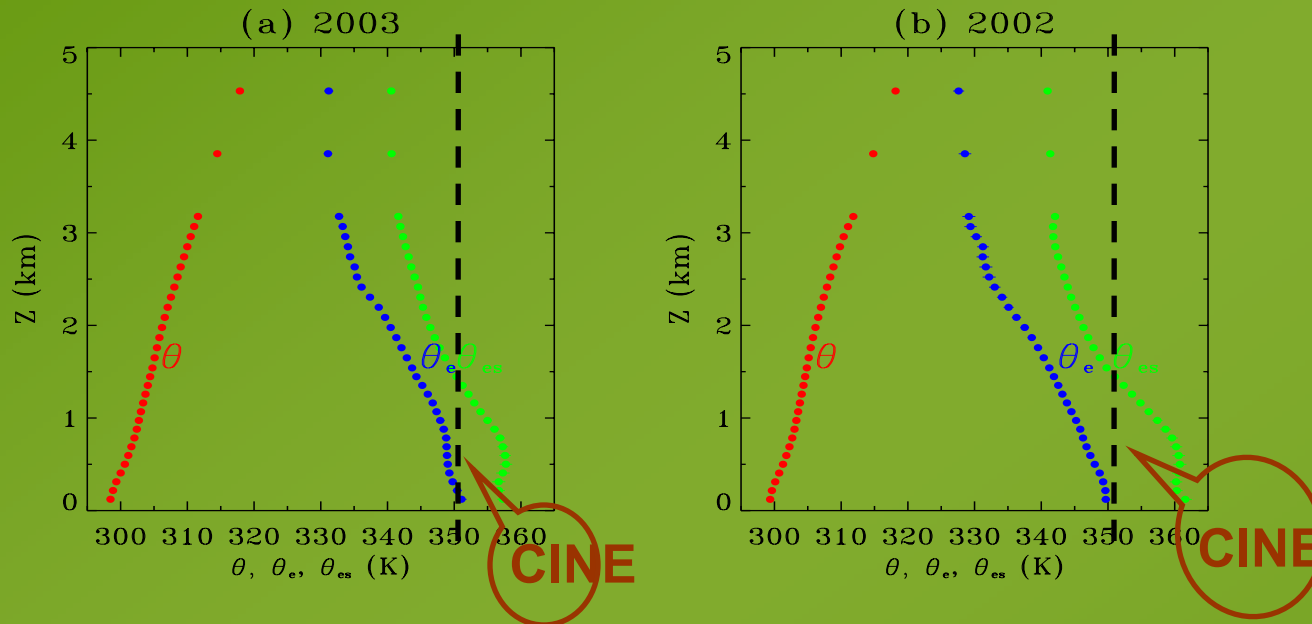


Manaus, Brazil



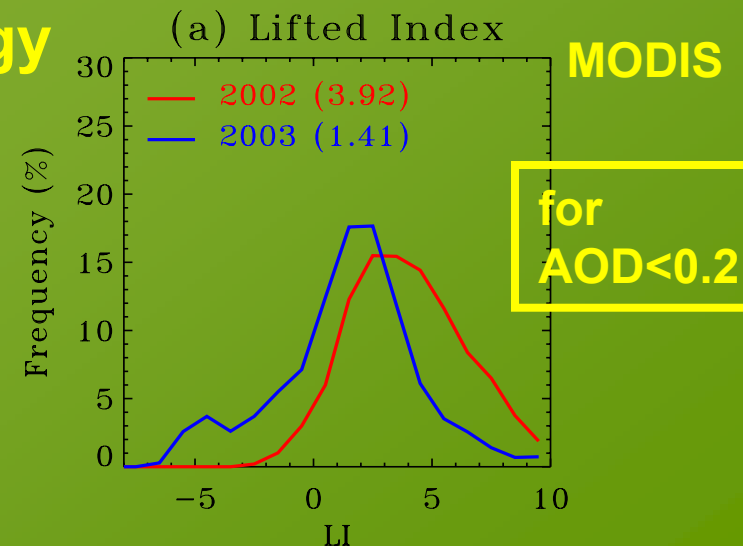
Radiosonde measurements

2003 was more unstable than 2002

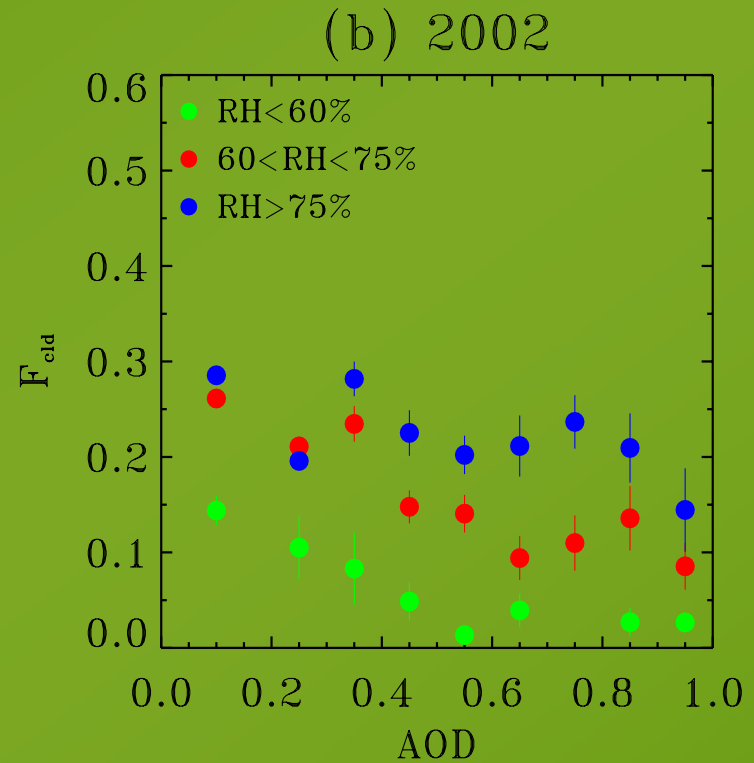
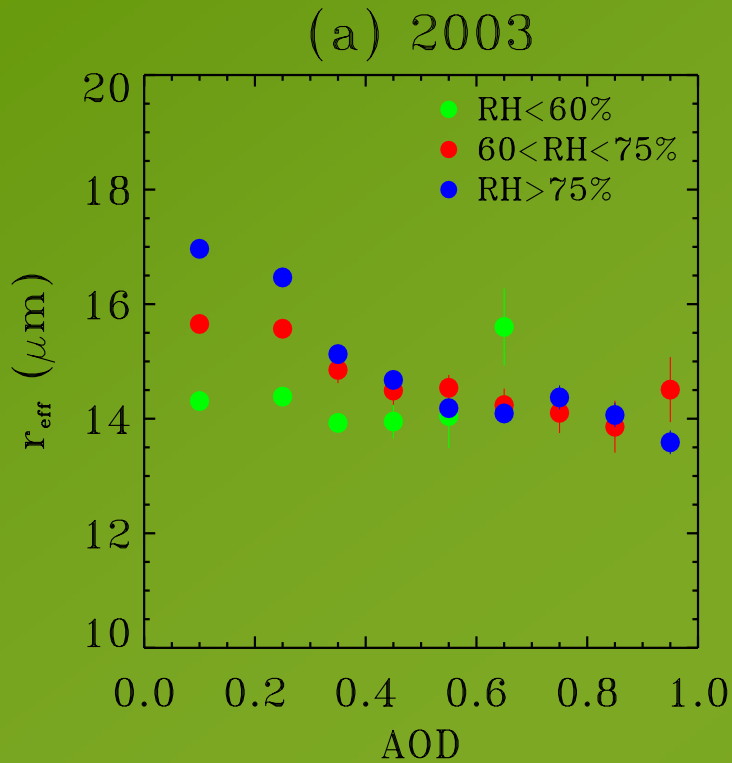


Convective Inhibiting Negative Energy (CINE): 2003 < 2002

also corroborated by MODIS Lifted Index (LI). A larger LI means more stable atmosphere.



Dependence on RH (@850mb)



2003: more rapid decrease of droplet size at higher RH.

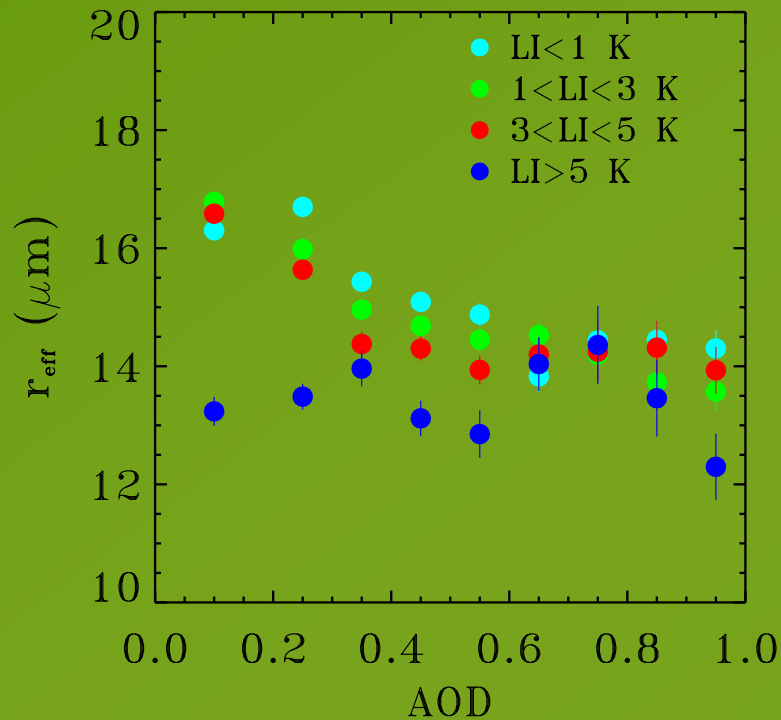
2002: more rapid reduction of cloud cover at lower RH.

Dependence on Lifted Index (LI)

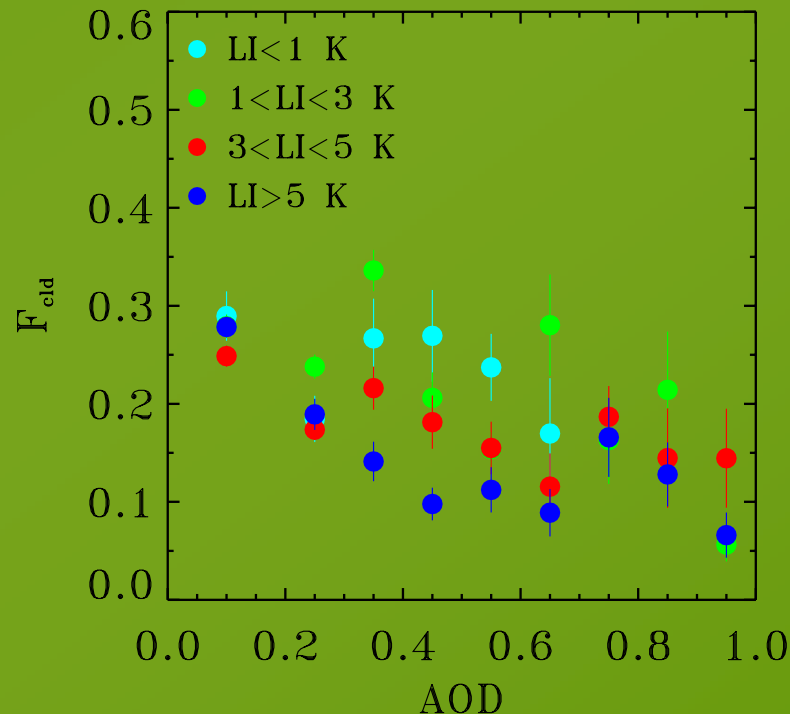
LI < 0 unstable;

LI > 0 stable

(a) 2003



(b) 2002

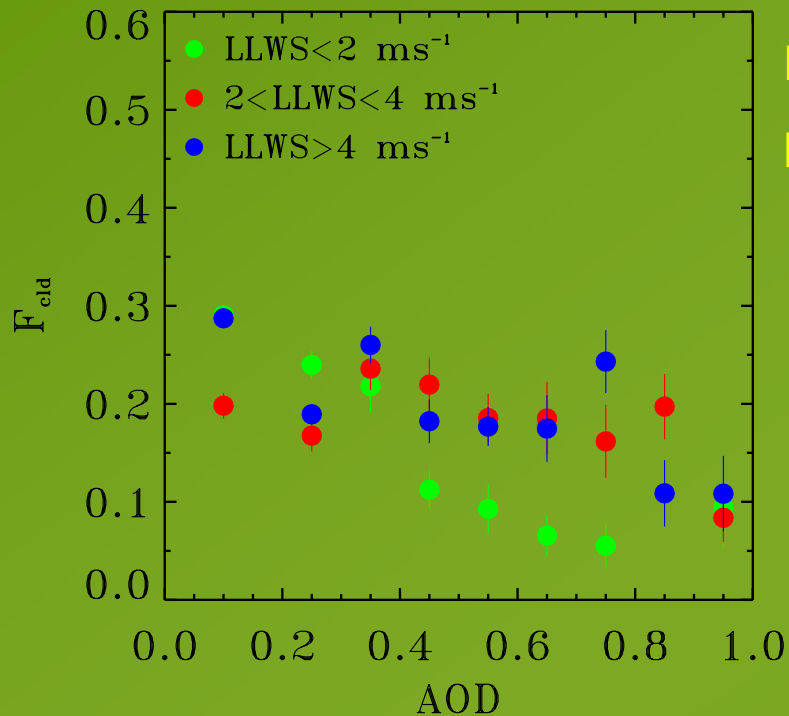


In 2003, Twomey effect is larger at smaller LI (unstable);

In 2002, the reduction of cloud fraction appears to be greater for larger LI (stable)

Dependence on Wind-Shear

(b) 2002



LLWS: from radiosonde measurements
between 925-700 mb level

* Cloud burning is more significant under weaker wind-shear, presumably because

- turbulence/convection are largely determined by surface fluxes for smaller wind-shear; and these can be significantly reduced by smoke absorption.

What have we learned from this study?

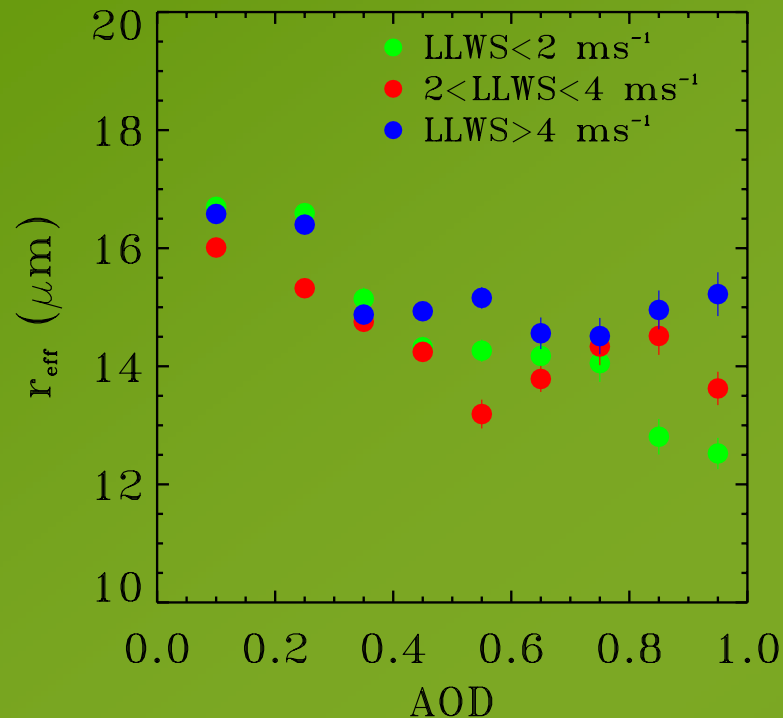
- ***Smoke influences clouds significantly but differently between a normal 2003 (indirect effects) and an ENSO 2002 (semi-direct effect).***
- ***Such differences correlate with differences in atmos. dynamics/thermodynamics:***
 - ***2003 was wetter and more unstable, favorable for activation of CN to CCN;***
 - ***2002 had a relatively dry ABL and less extensive clouds, presumably more susceptible to smoke absorption***
- ***Atmospheric dynamics & thermodynamics play a complex role in regulating the influences of aerosols on clouds. Findings at a specific time & location are not easily extrapolated to other times & places.***

END

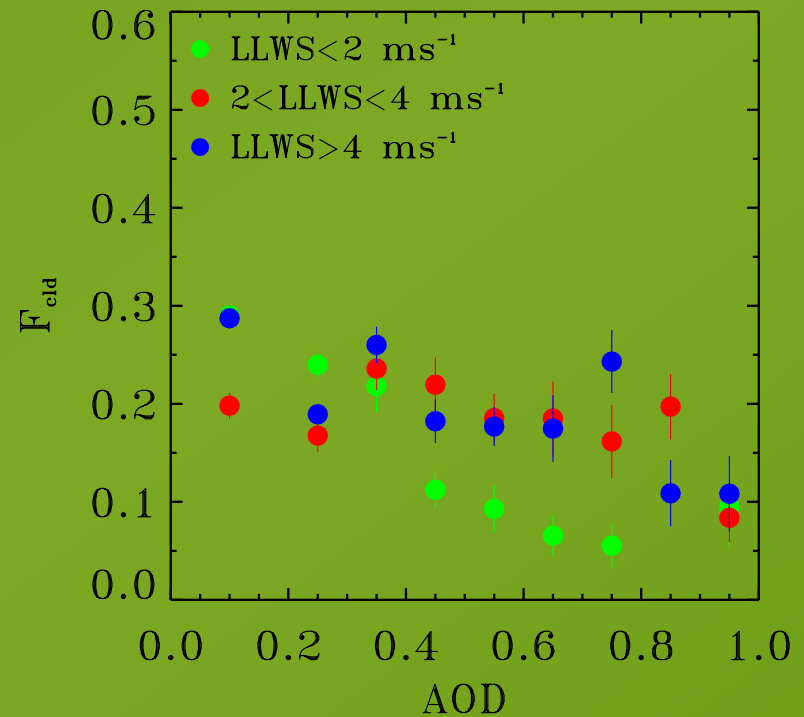
THANKS!

Dependence on Wind-Shear

(a) 2003



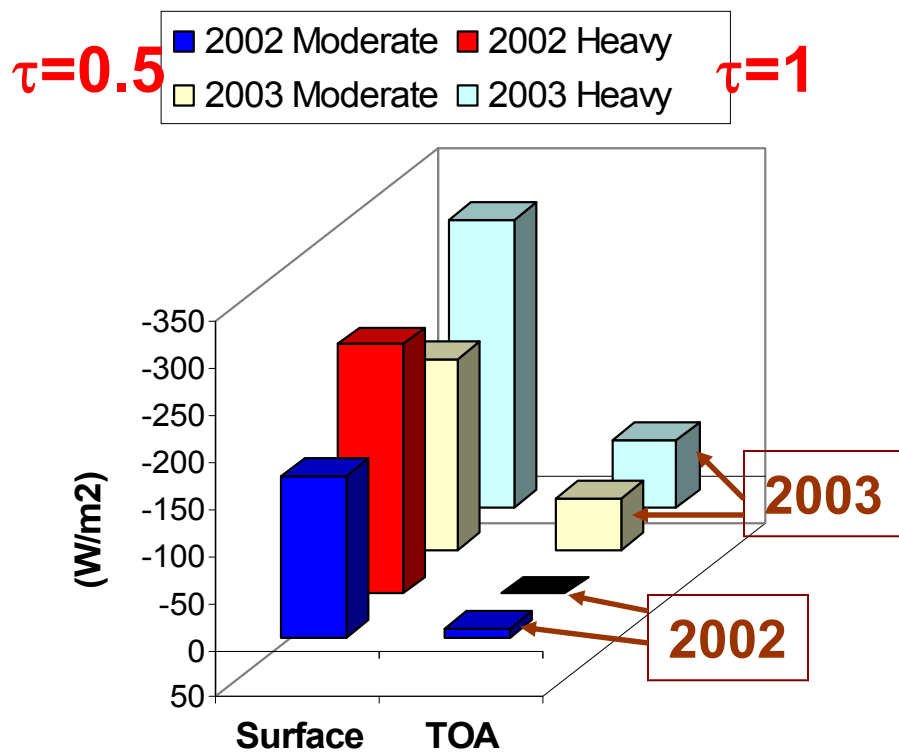
(b) 2002



- * LLWS: from radiosonde measurements between 925-700 mb level
- Cloud burning in 2002 is more significant under weaker wind-shear
 - presumably, turbulence/convection are largely determined by surface fluxes for smaller wind-shear; and these can be significantly reduced by smoke absorption.

Different smoke-cloud interactions result in interannual variations of aerosol forcing

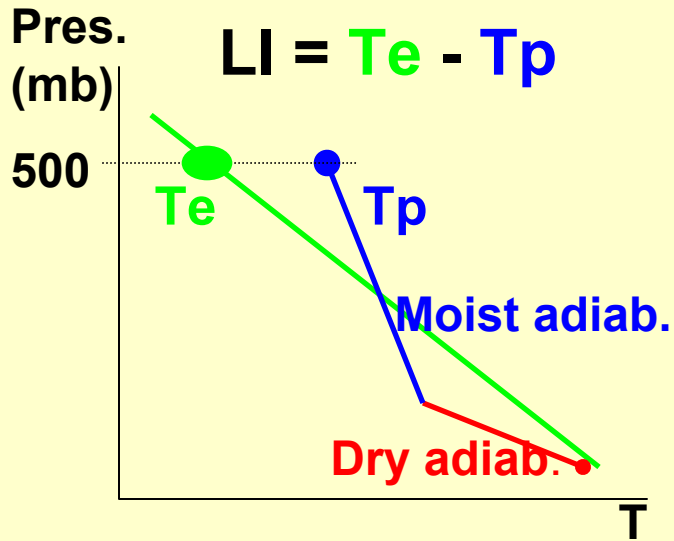
1:30 pm Instantaneous Forcing



- TOA cooling in 2003, but nearly zero in 2002
- Surface flux reduces by 20~30% (moderate smoke) and 35~45% (heavy smoke) (2003 more than 2002)
- Atmos. heating increases by 50-60% (moderate) and 80-90% (heavy).

• Possible influences on wet season onset (influenced by surface fluxes, thermodynamics)

MODIS Lifted Index (LI) for AOD < 0.2



2002

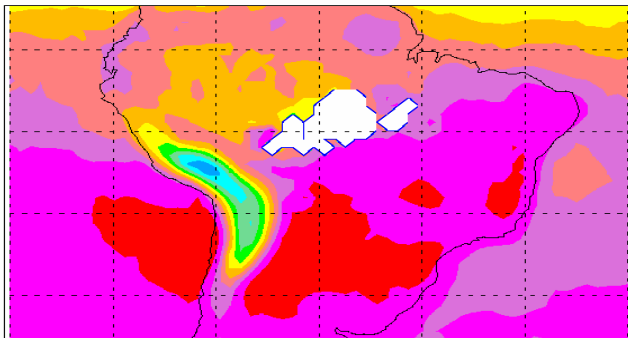
Tropics:

Smaller LI, less stable atmosphere in 2003;

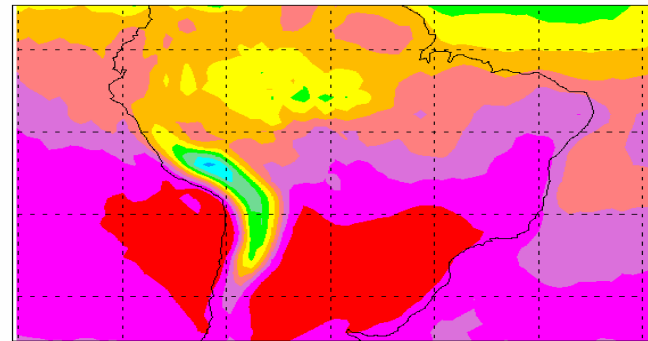
Stronger updrafts in 2003 could activate CN to CCN more efficiently.

2003

Lifted Index (LI)



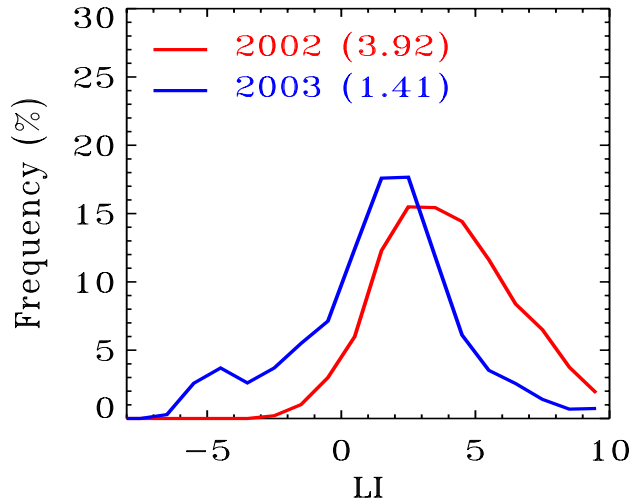
Lifted Index (LI)



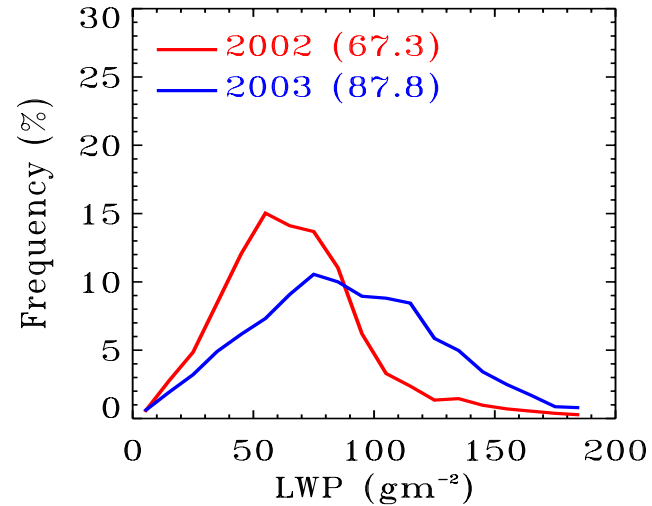
MODIS Observed PDFs

---- for AOD < 0.2

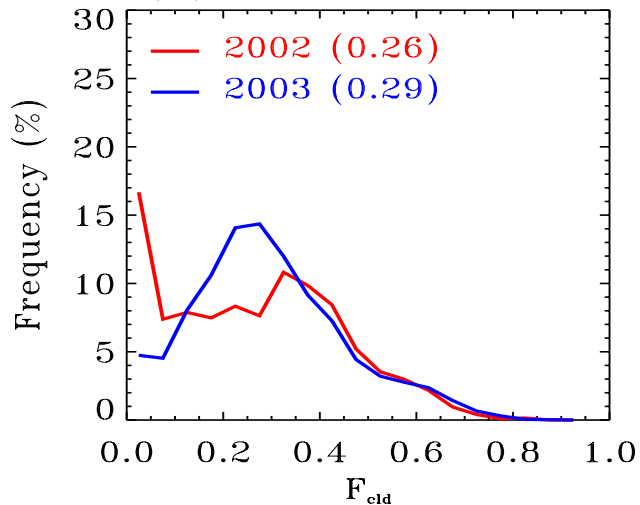
(a) Lifted Index



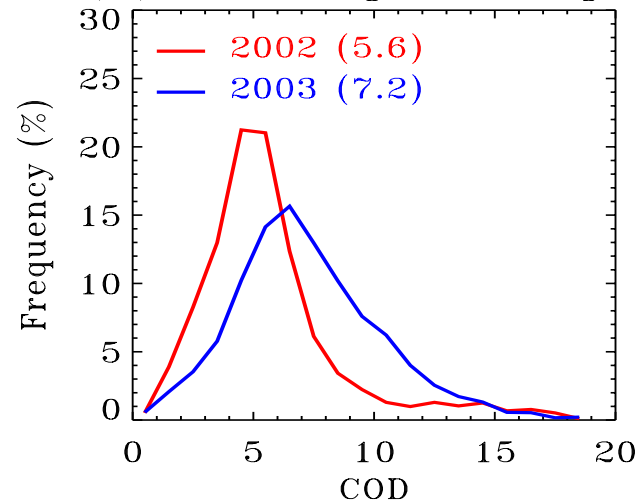
(b) Liquid Water Path



(c) Cloud Fraction

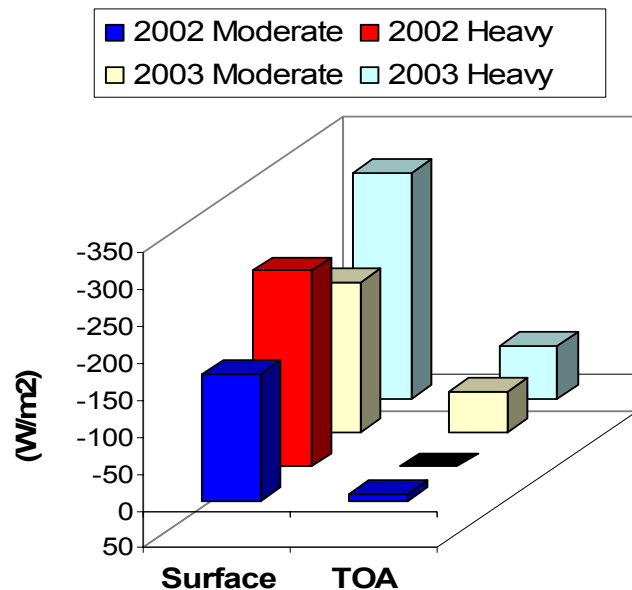


(d) Cloud Optical Depth

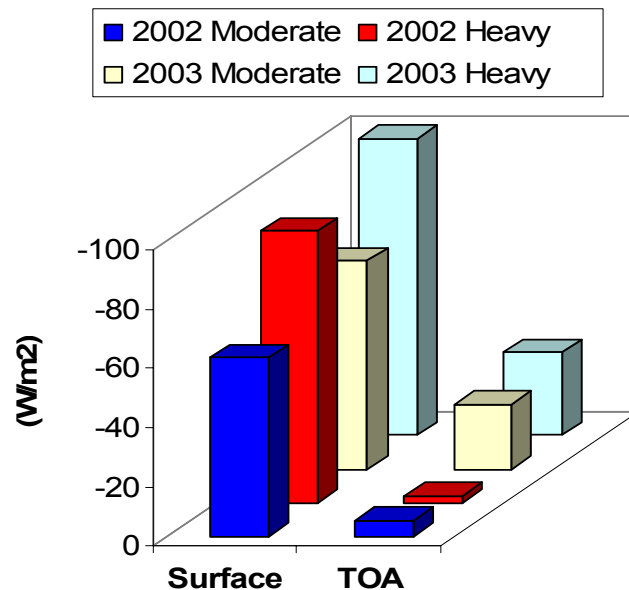


Different smoke-cloud interactions result in interannual variations of aerosol forcing

1:30 pm Instantaneous Forcing



24-h Average Forcing



- TOA cooling in 2003, but nearly zero in 2002
- Surface flux reduces by 20~30% (moderate smoke) and 35~45% (heavy smoke) (2003 more than 2002)
- Atmos. heating increases by 50-60% (moderate) and 80-90% (heavy).
- Possible influences on wet season onset (determined by surface fluxes, thermodynamics)