



Volcanic eruptions, global cooling
and ozone loss: historic examples
and what we need to measure after
the next eruption

Brian Toon

Department of Atmospheric and Oceanic Sciences
Laboratory for Atmospheric and Space Physics,
University of Colorado, Boulder

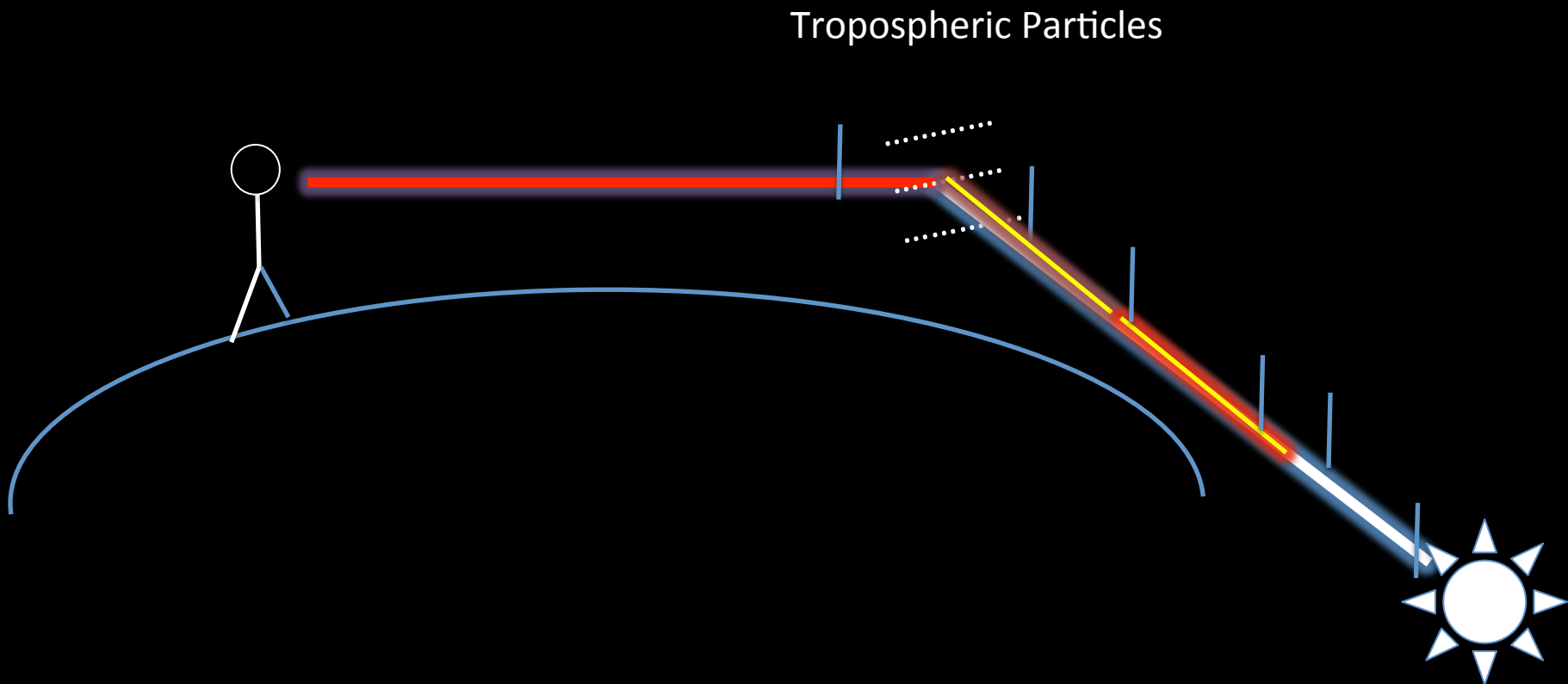
Volcanic clouds impact climate by scattering sunlight;
twilight is delayed and turned purple



Sunset in Zimbabwe on July 27, 2015 after the April 22, 2015 eruption of Calbuco in
Chile

Peter Lowenstein

Physics of normal sunsets involves
Rayleigh scattering of blue light and
tropospheric particles scattering light
toward you



Volcanic clouds impact climate by scattering sunlight;
twilight is delayed and turned purple

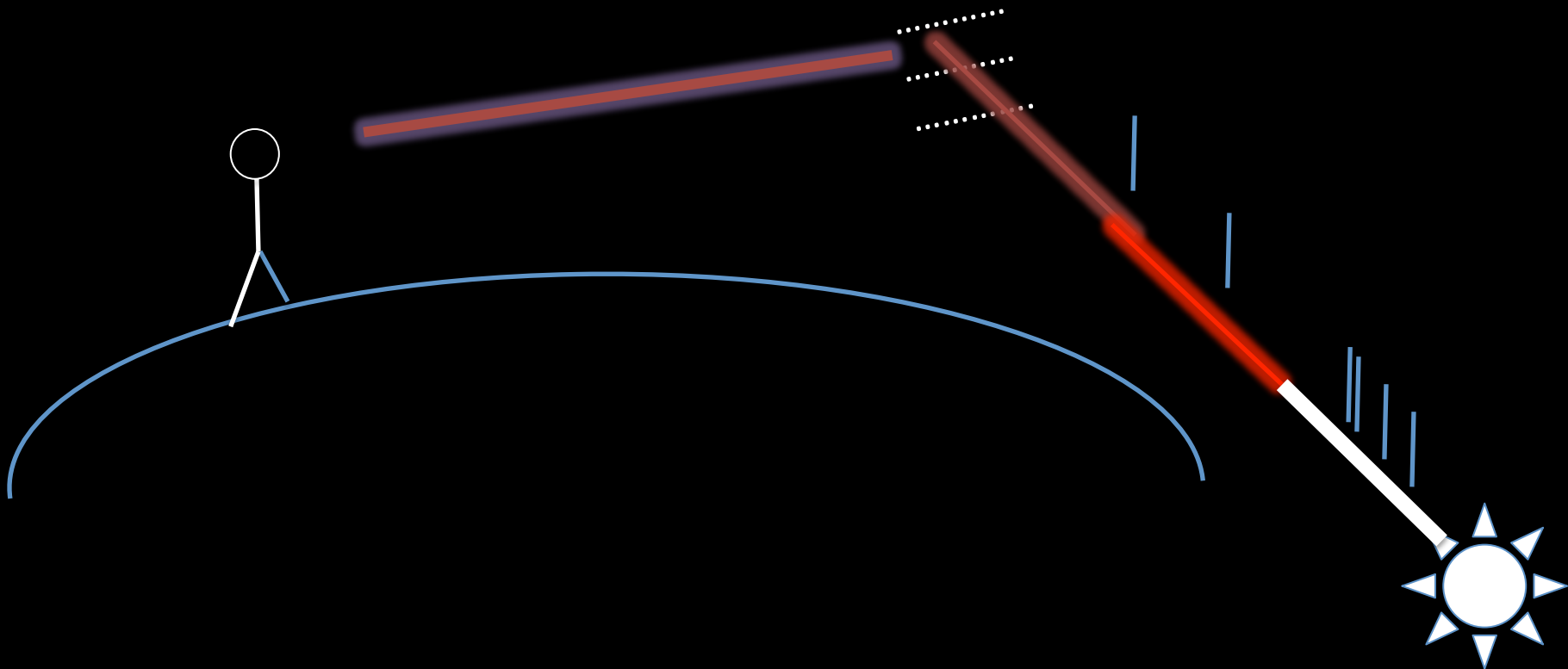


Sunset in Zimbabwe on July 27, 2015 after the April 22, 2015 eruption of Calbuco in
Chile

Peter Lowenstein

Physics of volcanic sunsets involves Rayleigh scattering of blue light, ozone absorption of red light, volcanic particles scattering towards you

Stratospheric Volcanic Particles + Ozone



Volcanic clouds impact climate by scattering sunlight;
twilight is delayed and turned purple

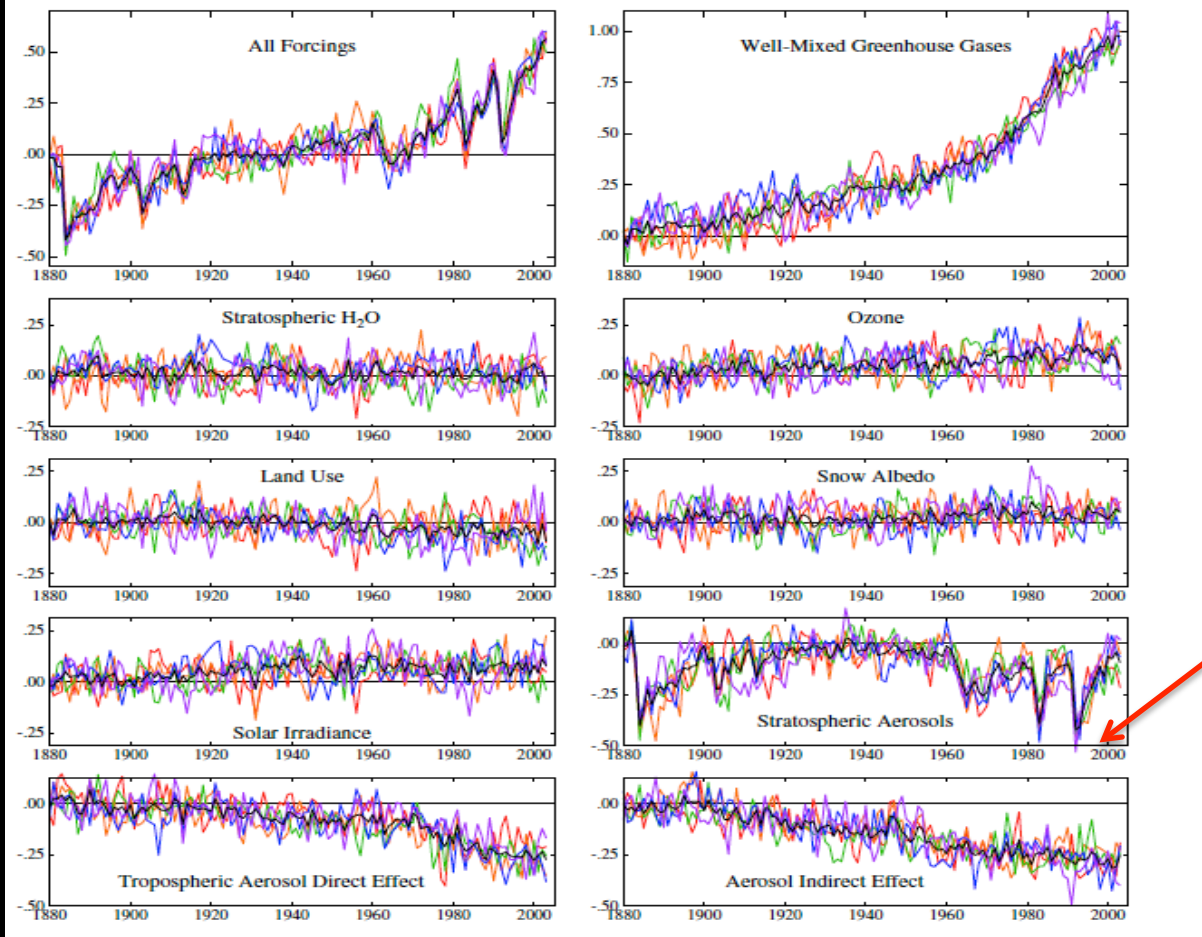


Sunset in Zimbabwe on July 27, 2015 after the April 22, 2015 eruption of Calbuco in
Chile

Peter Lowenstein

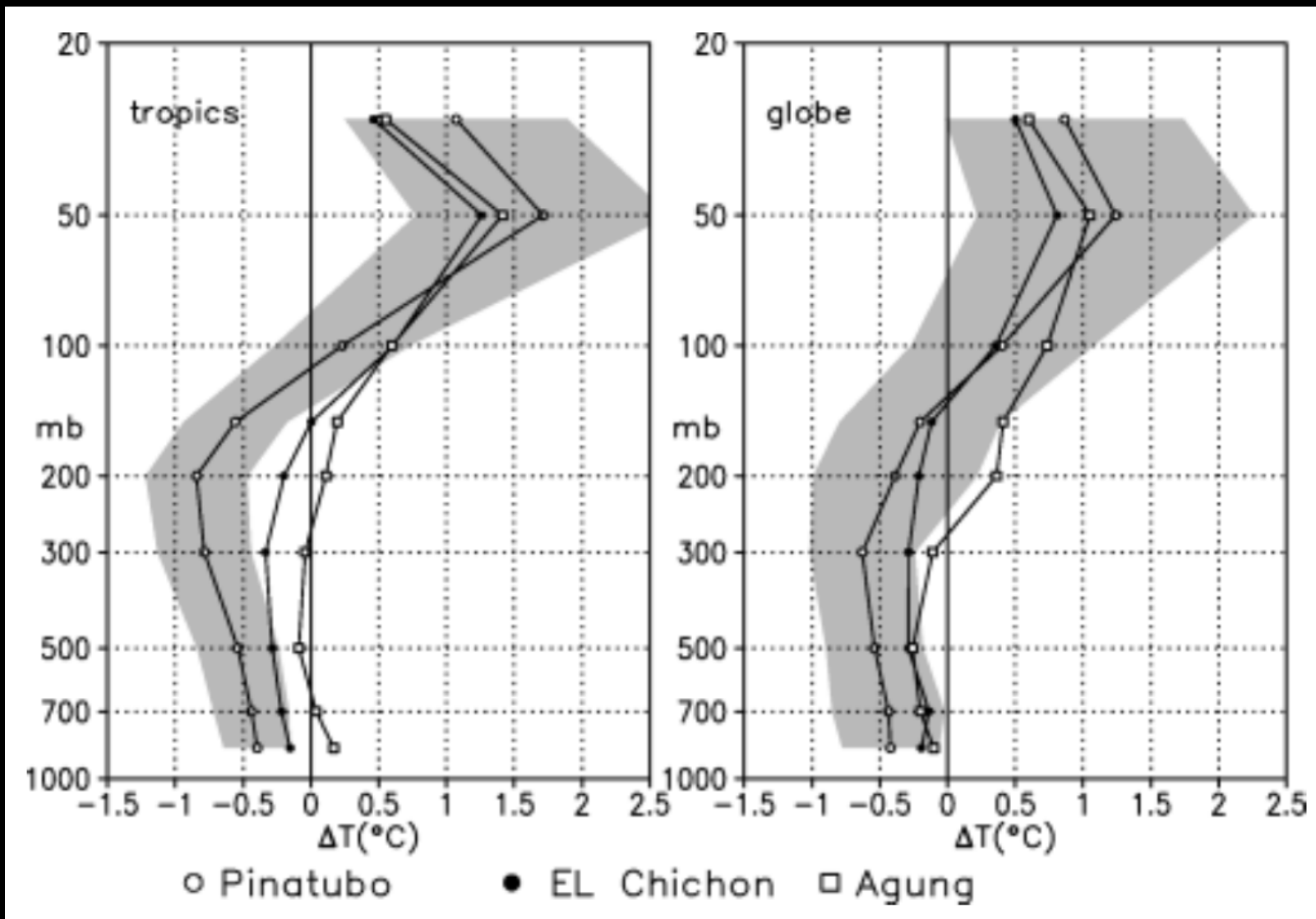
Volcanic aerosols cool the planet

Annual-Mean Global-Mean Surface Temperature Change (°C)



Hansen et al., 2007
Clim. Dyn. 29,661

Most focus has been on temperature



Mt. Tambora, April 1815 largest known eruption in 1000 years



Tambora and the Year Without Summer-1816

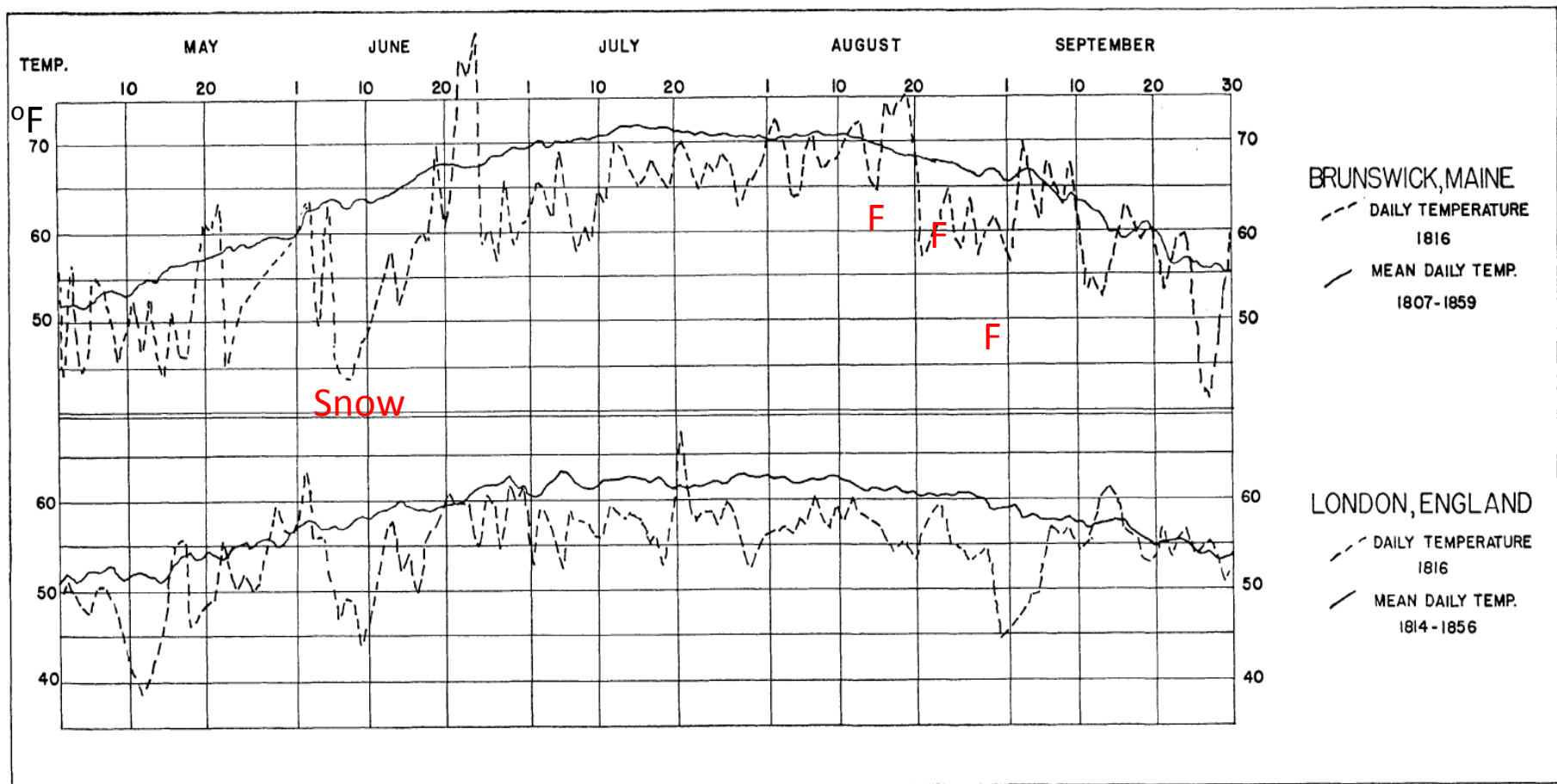


FIG. 6. Daily and mean daily temperatures—Brunswick, Maine, and London, England.

Food prices jumped after Tambora eruption

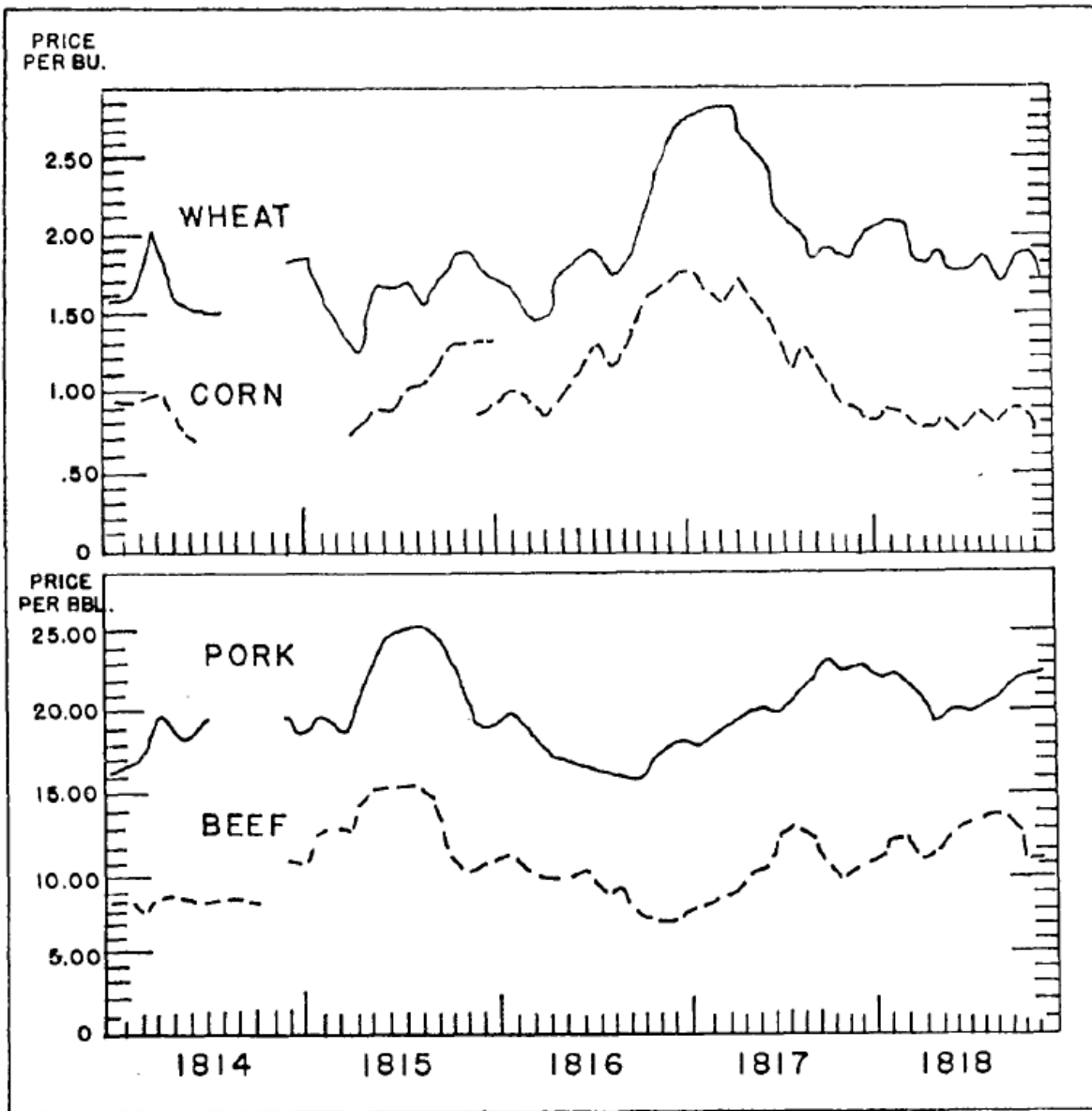


FIG. 5. Monthly average wholesale food prices, produce market, New York, 1814-18.

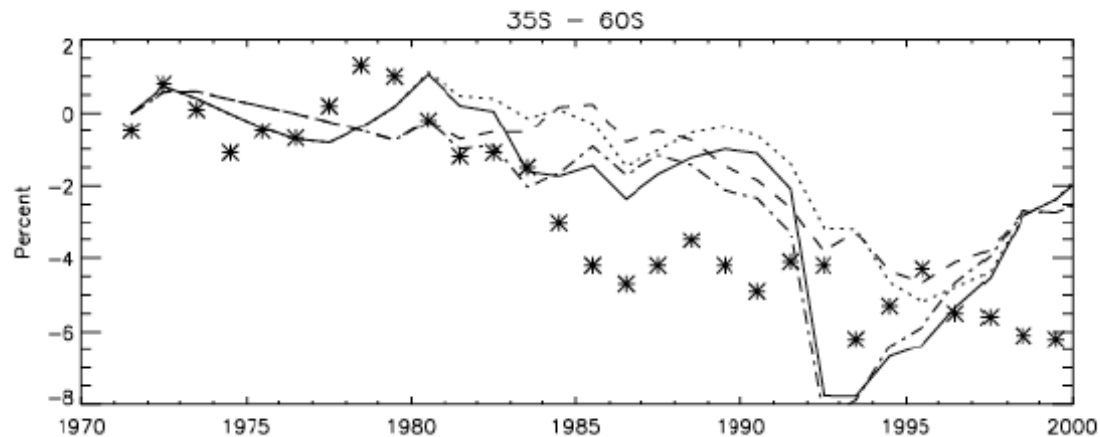
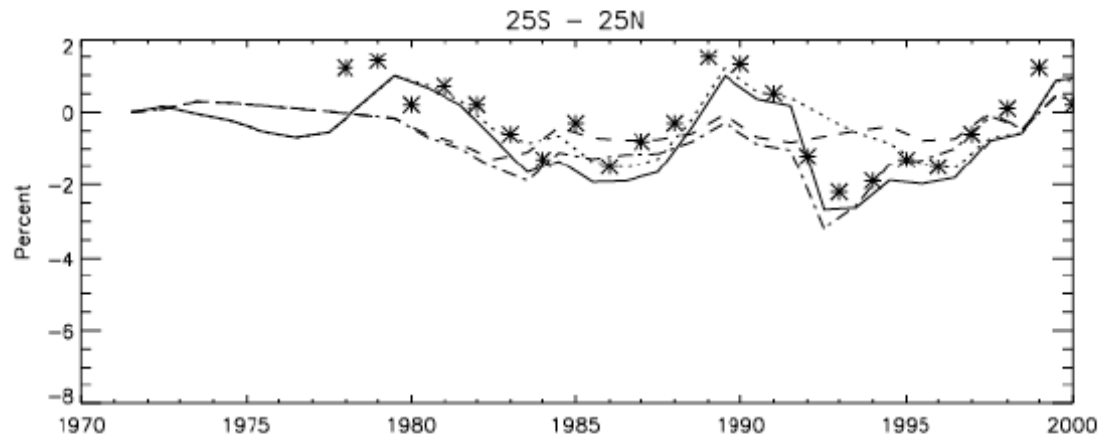
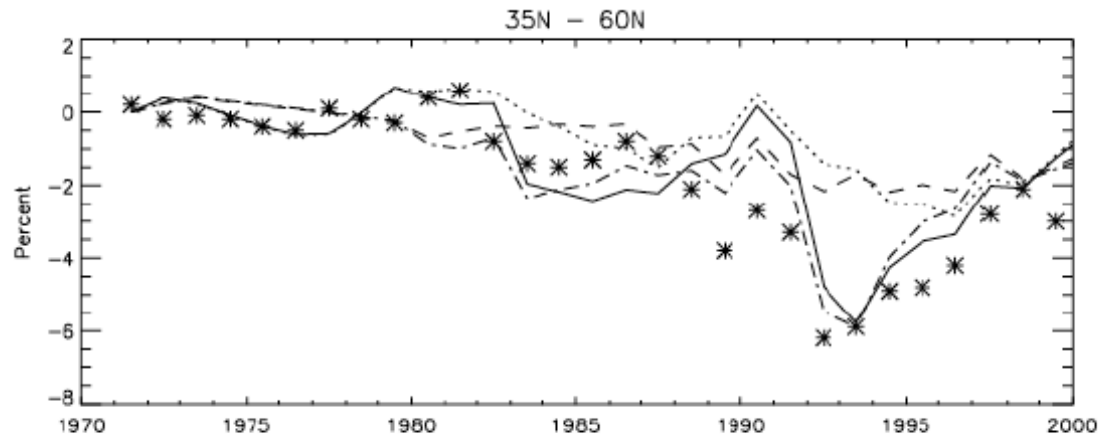
People fled New England post Tambora

TABLE 1.—RATES OF POPULATION INCREASE, NEW ENGLAND STATES, 1790–1860

Area	Rates of increase, in percentages						
	1790 to 1800	1800 to 1810	1810 to 1820	1820 to 1830	1830 to 1840	1840 to 1850	1850 to 1860
United States	35	36	33	33	32	36	36
New England	34	19	13	18	14	22	15
Maine	57	51	30	34	26	16	8
New Hampshire	30	16	14	10	6	12	3
Vermont	82	40	8	19	4	8	0.3
Massachusetts	12	11	10	17	21	35	24
Connecticut	6	4	5	8	4	19	24
Rhode Island	0.4	11	8	17	11	18	36

Many changes observed after eruptions

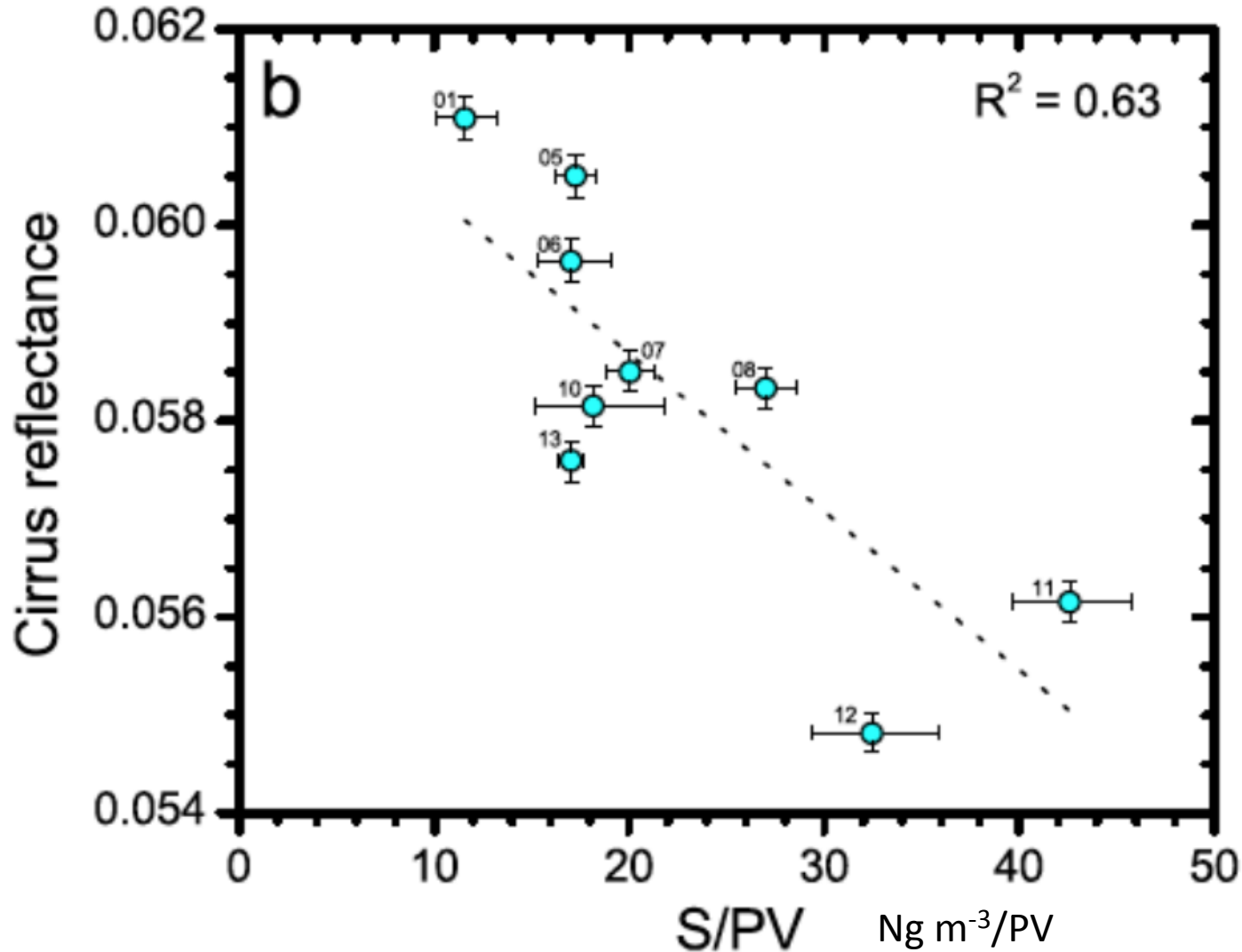
Observed	Probable cause
Cooling troposphere and surface	Reduction in shortwave forcing by aerosol
Tropopause and stratospheric warming	Sunlight and ir absorption by aerosol
Mid-lat. N.H. winter warming	Strat./troposphere dynamical interaction
Rapid spread of volcanic clouds	Alteration of atmospheric dynamics
Ozone loss	Heterogeneous reactions on sulfate aerosols
Hazy skies/bright twilights/ reduction in shortwave at surface	Scattering by aerosols
Change in stratospheric CH ₄ , H ₂ O	Change in dynamics, tropopause T
Change in tropospheric CO ₂ , CO, CH ₄	Increase/Reduction in UV in troposphere, drop in sea surface T, coincidence
Reduction in water vapor column	Sea surface cooling
Reduction in precipitation	Reduction of solar heating of sea surface
Expected	
Cirrus cloud increase/decrease	Seeding by large sulfate particles
Cooler days	Loss of sunlight
Cooler nights	Loss of sunlight, little IR change
Polar amplification	Decreased poleward energy flux
Increase in sea ice	Polar cooling



Column ozone depleted by El Chichon (82) and Pinatubo (91), but likely not important in earlier eruptions

Geller and Smyshlyaev

Cirrus reflectivity decreases as sulfur in lower stratosphere goes up

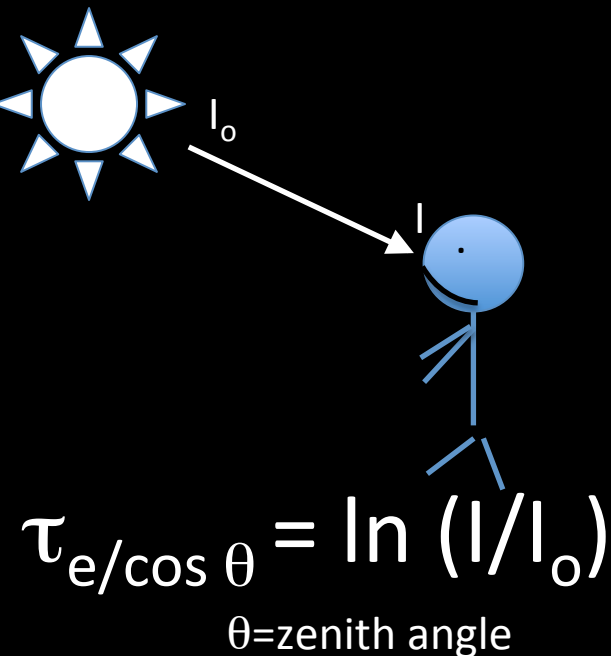


Friberg et al.,
2015

Need to know three things to calculate radiation field

AOD (AOT) = Aerosol extinction Optical Depth (Thickness)

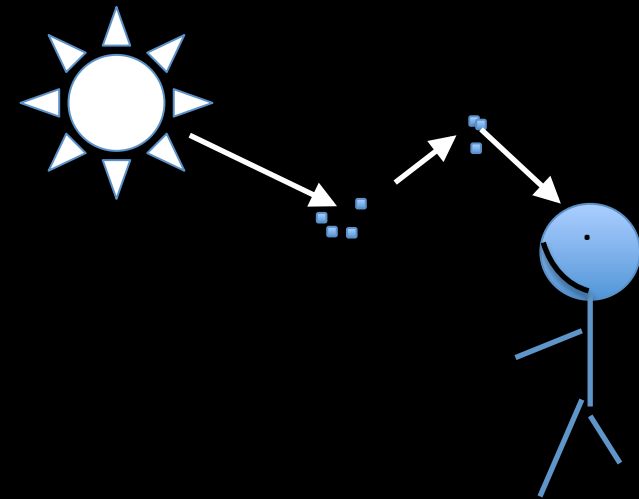
1. Extinction optical depth



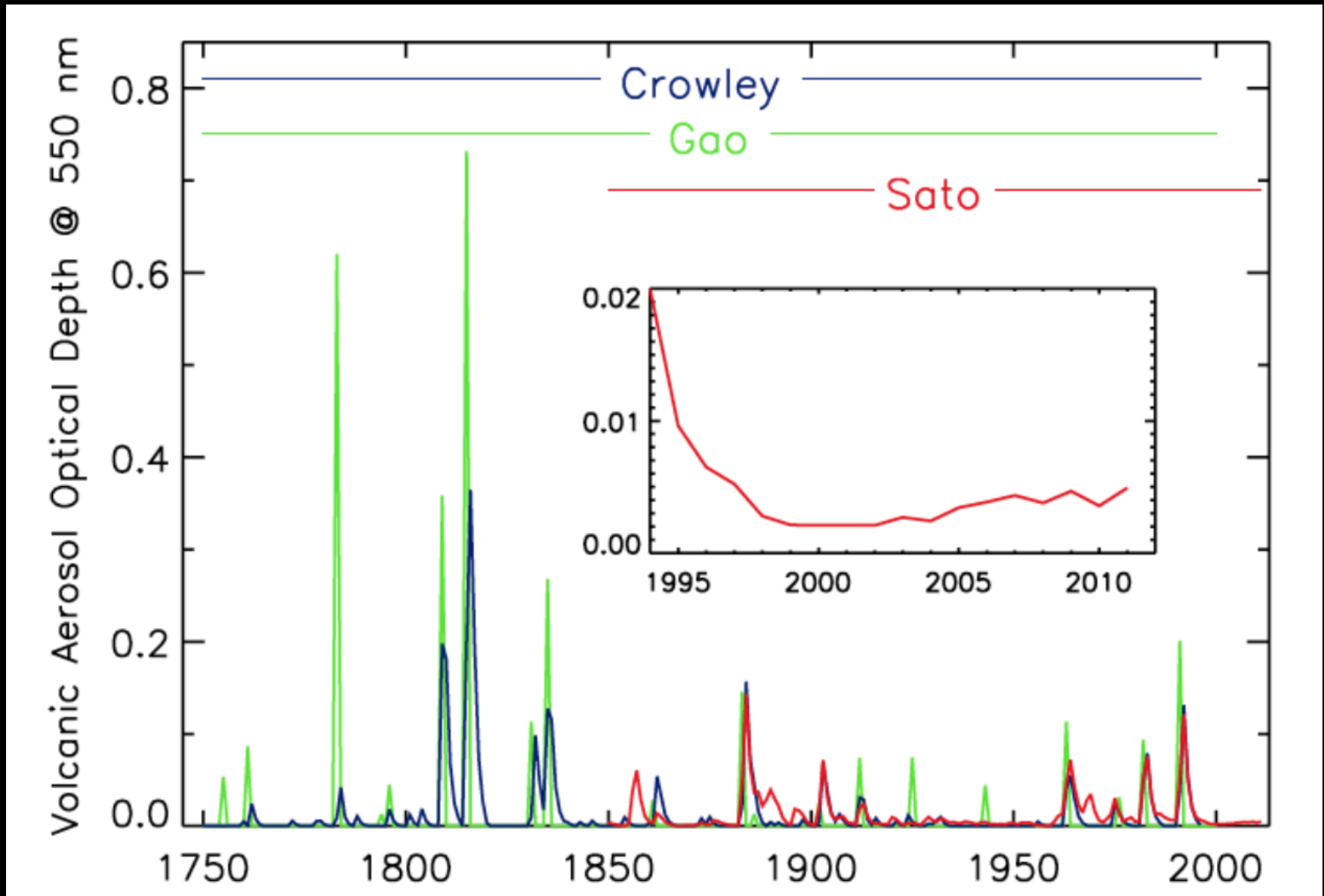
2. Absorption or scattering optical depth

$$\tau_e = \tau_a + \tau_s$$

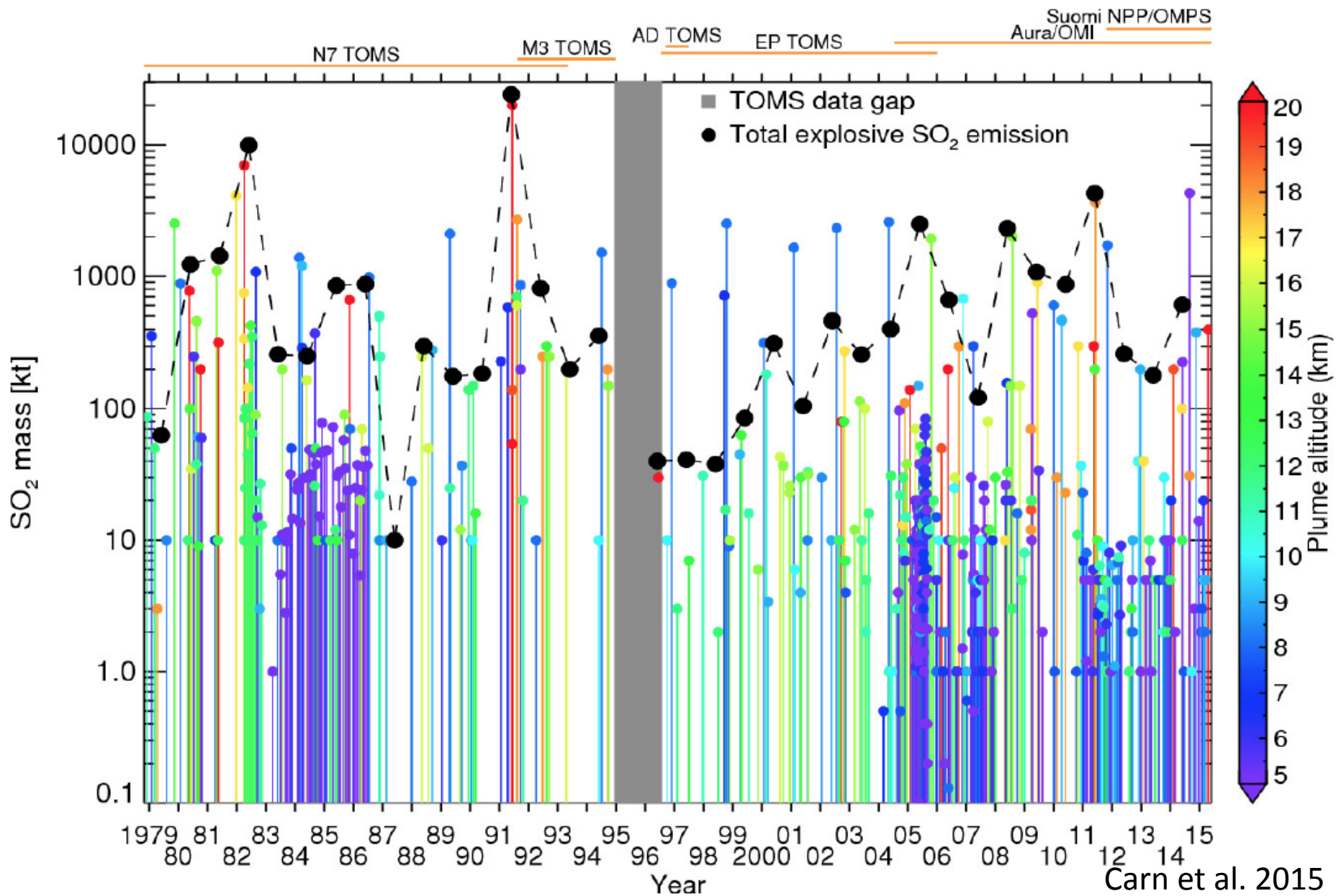
3. Scattering phase function



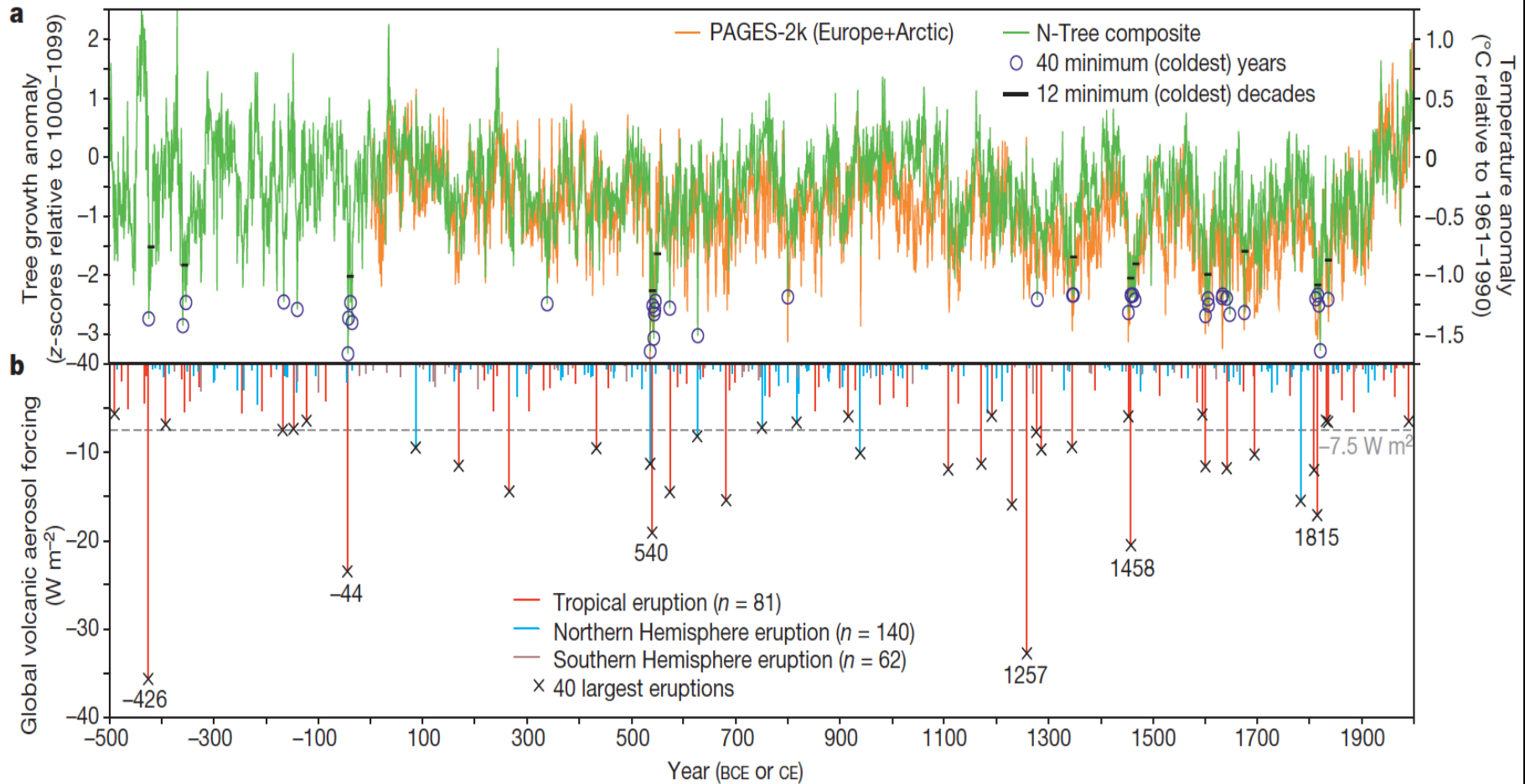
Early studies used volcanic explosivity to rank eruptions but it is not relevant to optical depth



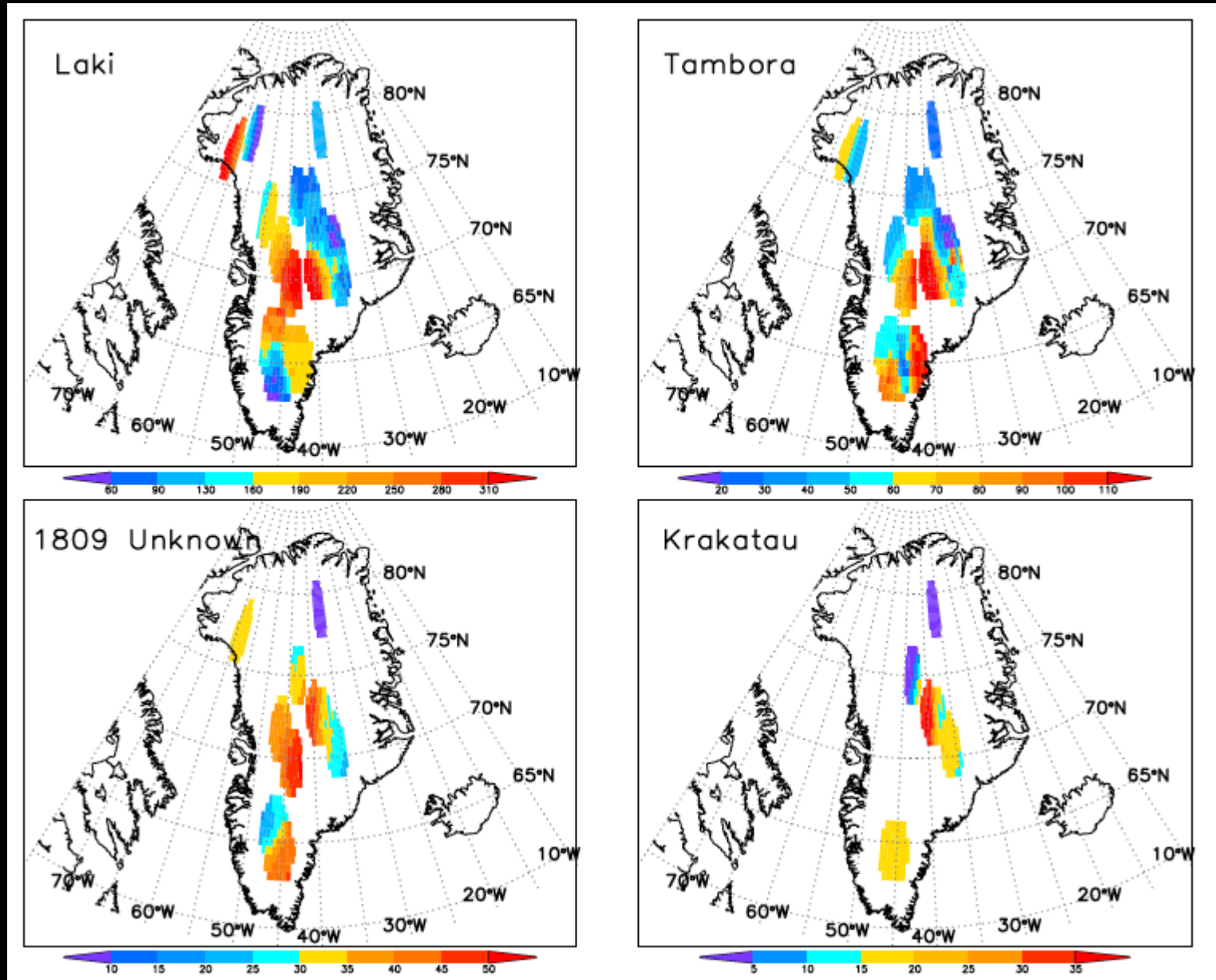
SO₂ injection is key to volcano climate effects



2500 years of volcanic forcing and climate change (Europe/Arctic)



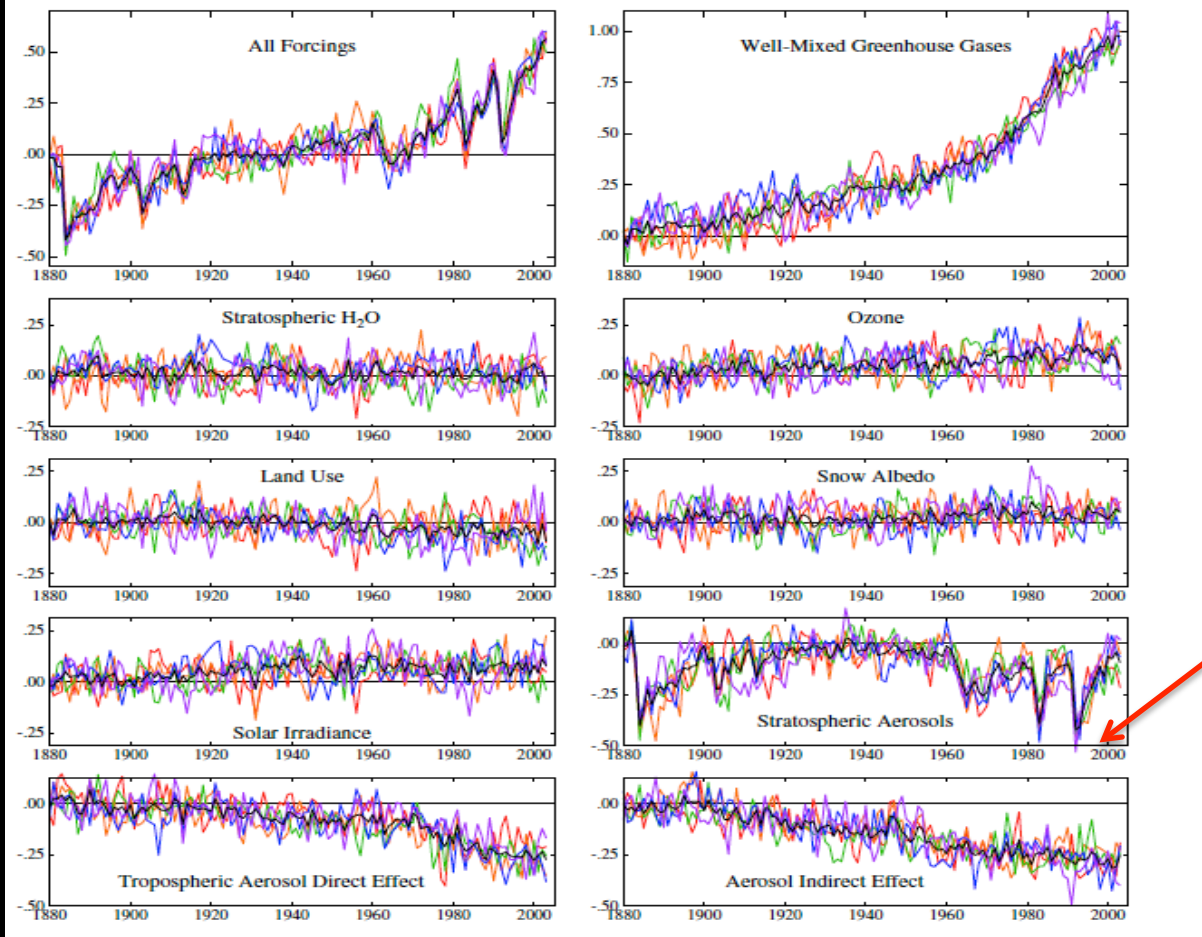
Polar core sulfate is highly variable



$\text{kg SO}_4/\text{km}^2$

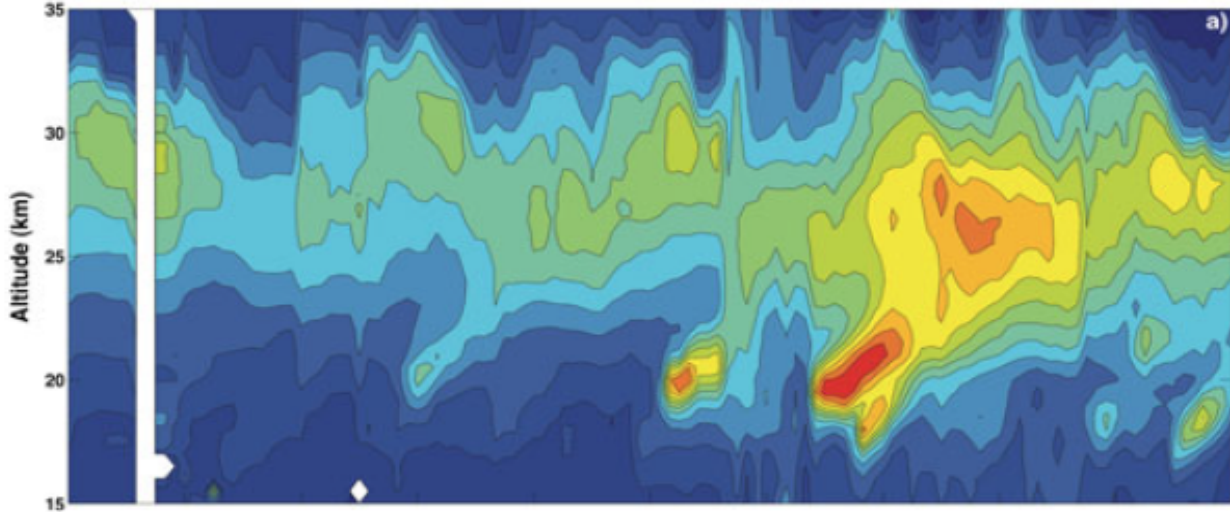
Large, rare, volcanic eruptions cool the planet

Annual-Mean Global-Mean Surface Temperature Change ($^{\circ}\text{C}$)

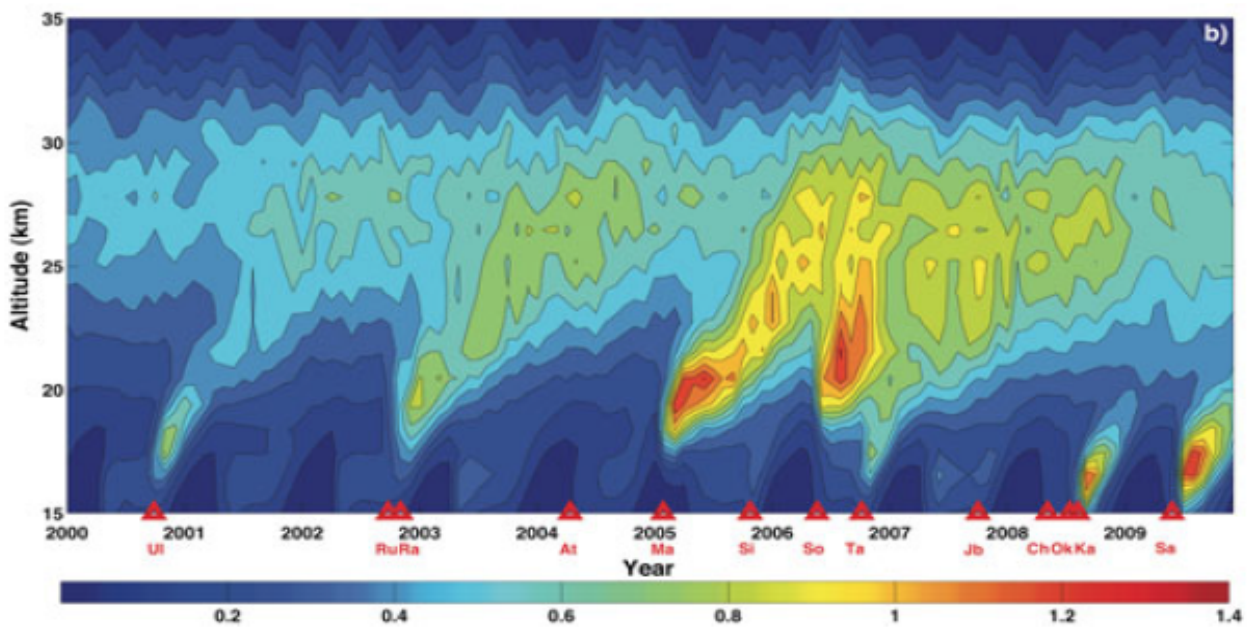


Hansen et al., 2007
Clim. Dyn. 29,661

Satellites detect and models can simulate extinction ratio for small volcanic clouds



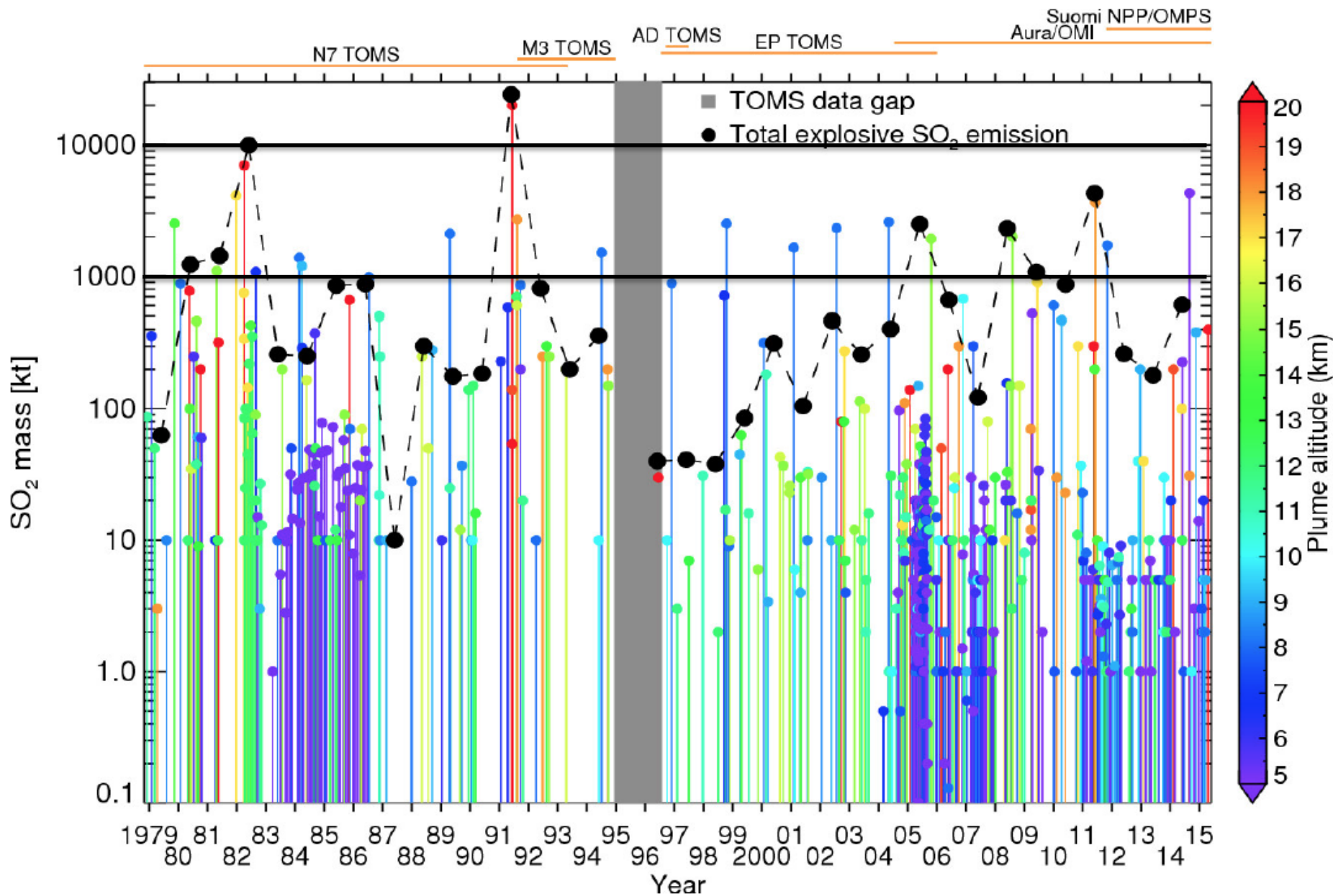
Data
Vernier et al., 2011



Model
Neely III et al., 2013

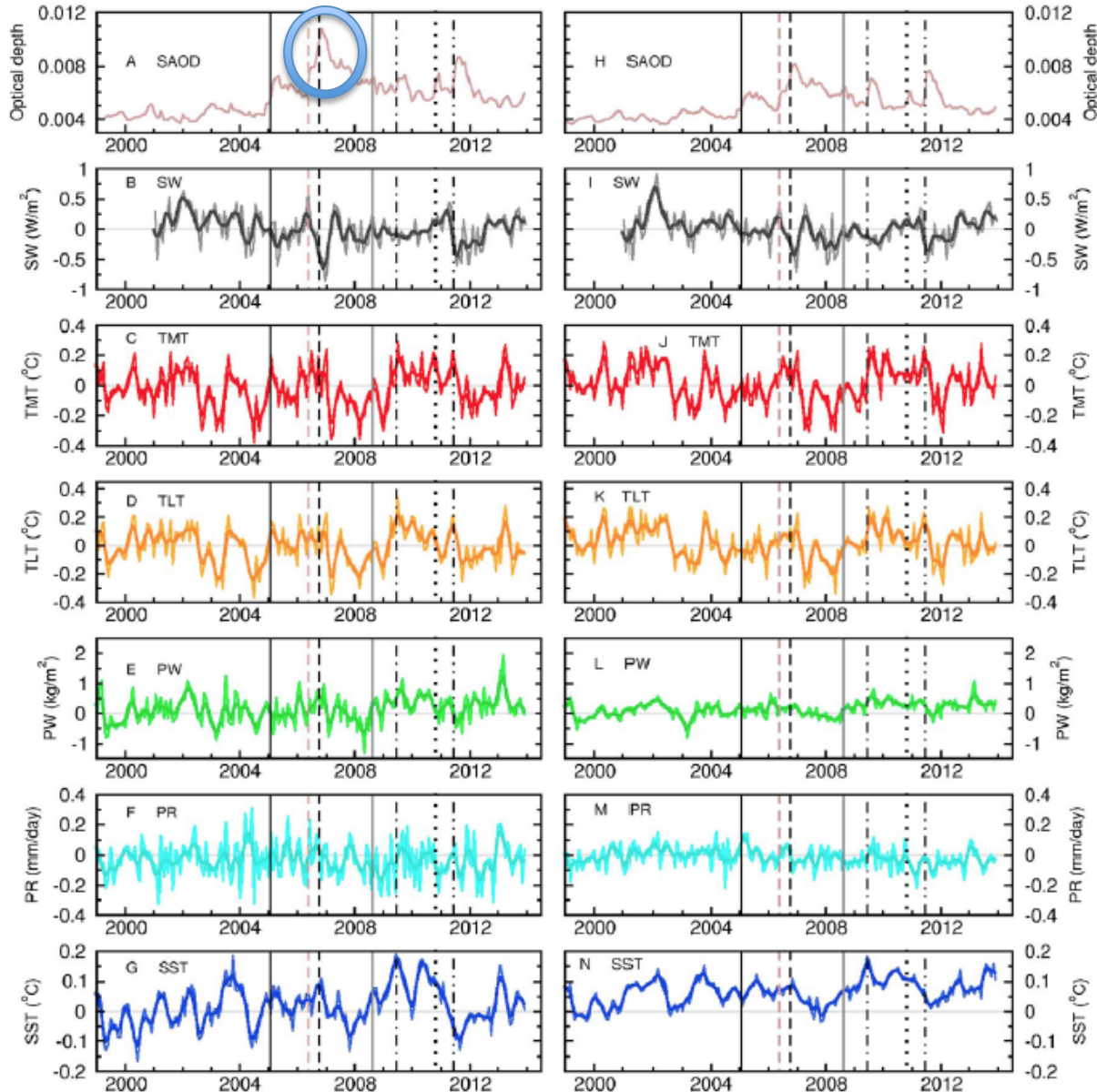
Extinction ratio

SO₂ is key to volcano climate effects



Tropical (20°N-20°S)

Near-global (50°N-50°S)



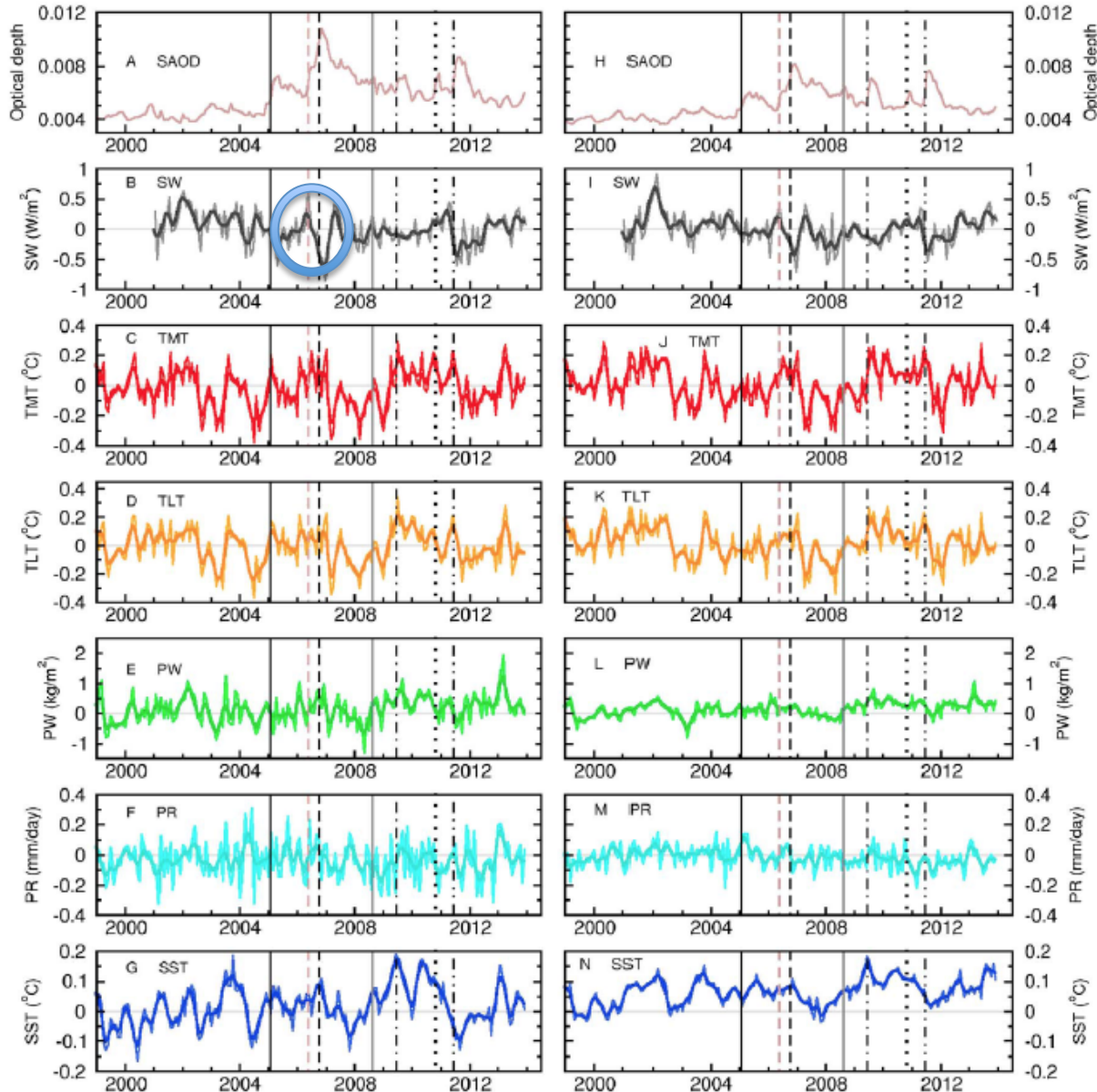
— Manam
- - - Soufriere Hills
- · - Tavorur
- - - Kasatochi
· · · Sarychev
- · · Merapi
- · - Nabro

Small,
common,
eruptions
have
detectable
signals in
aerosol
optical depth

Santer et al. 2015

Tropical (20°N-20°S)

Near-global (50°N-50°S)

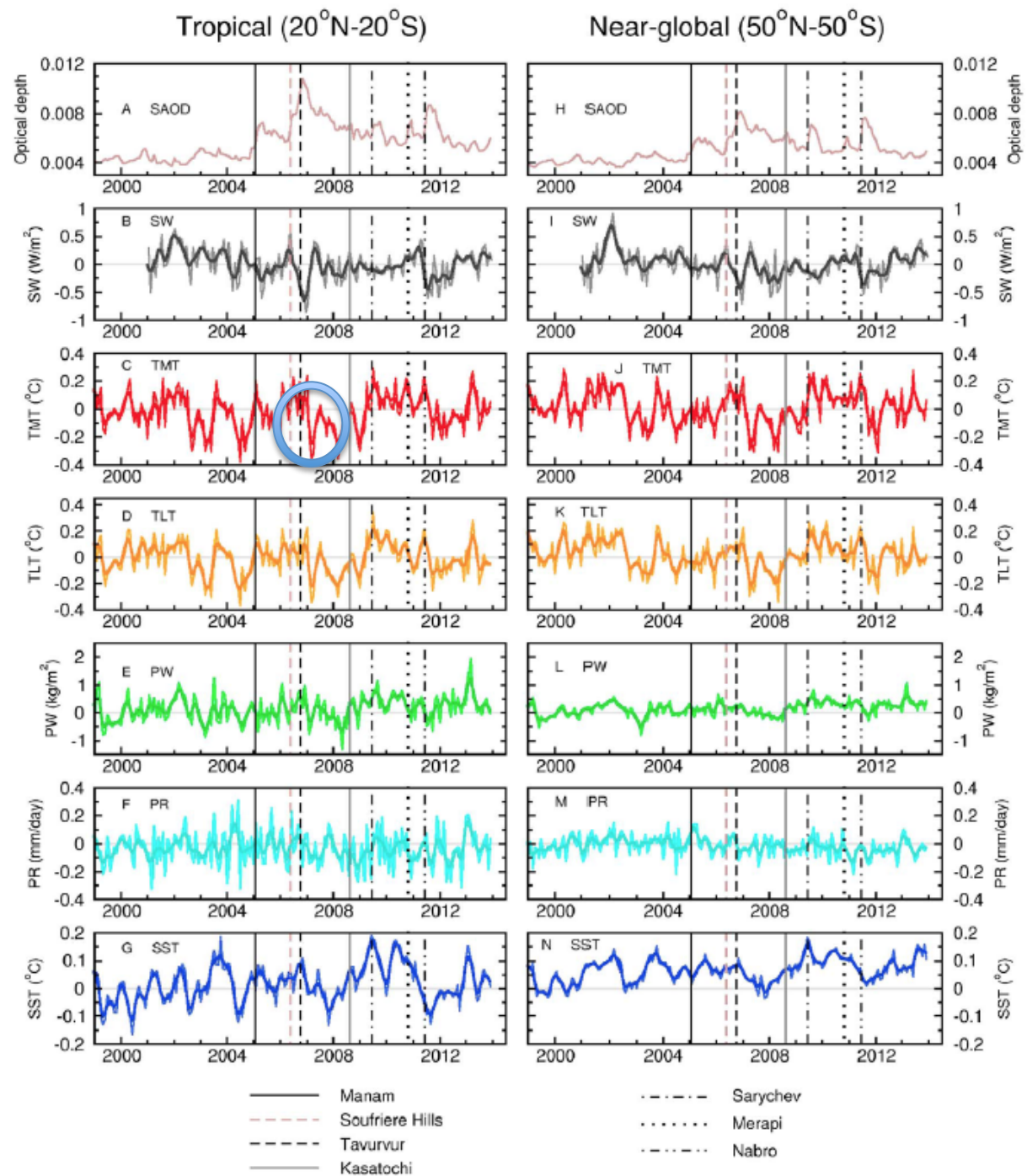


— Manam
- - - Soufriere Hills
- - - Tavorur
— Kasatochi

- - - Sarychev
... Merapi
- - - Nabro

Small,
common,
eruptions
have
detectable
signals in
short wave
flux

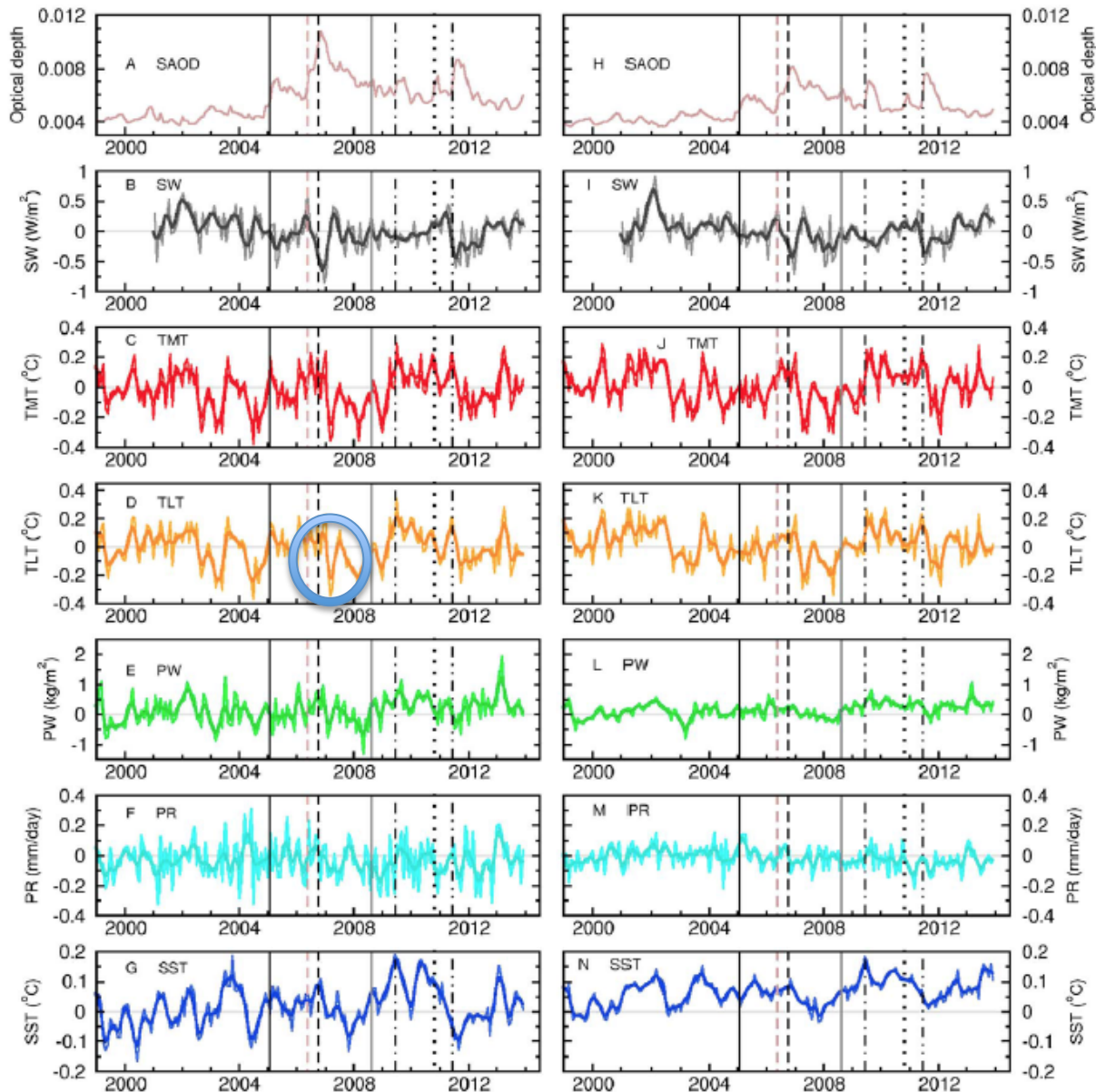
Santer et al. 2015



Small,
common,
eruptions
have
detectable
signals in
temperature
of middle and
upper
troposphere

Tropical (20°N-20°S)

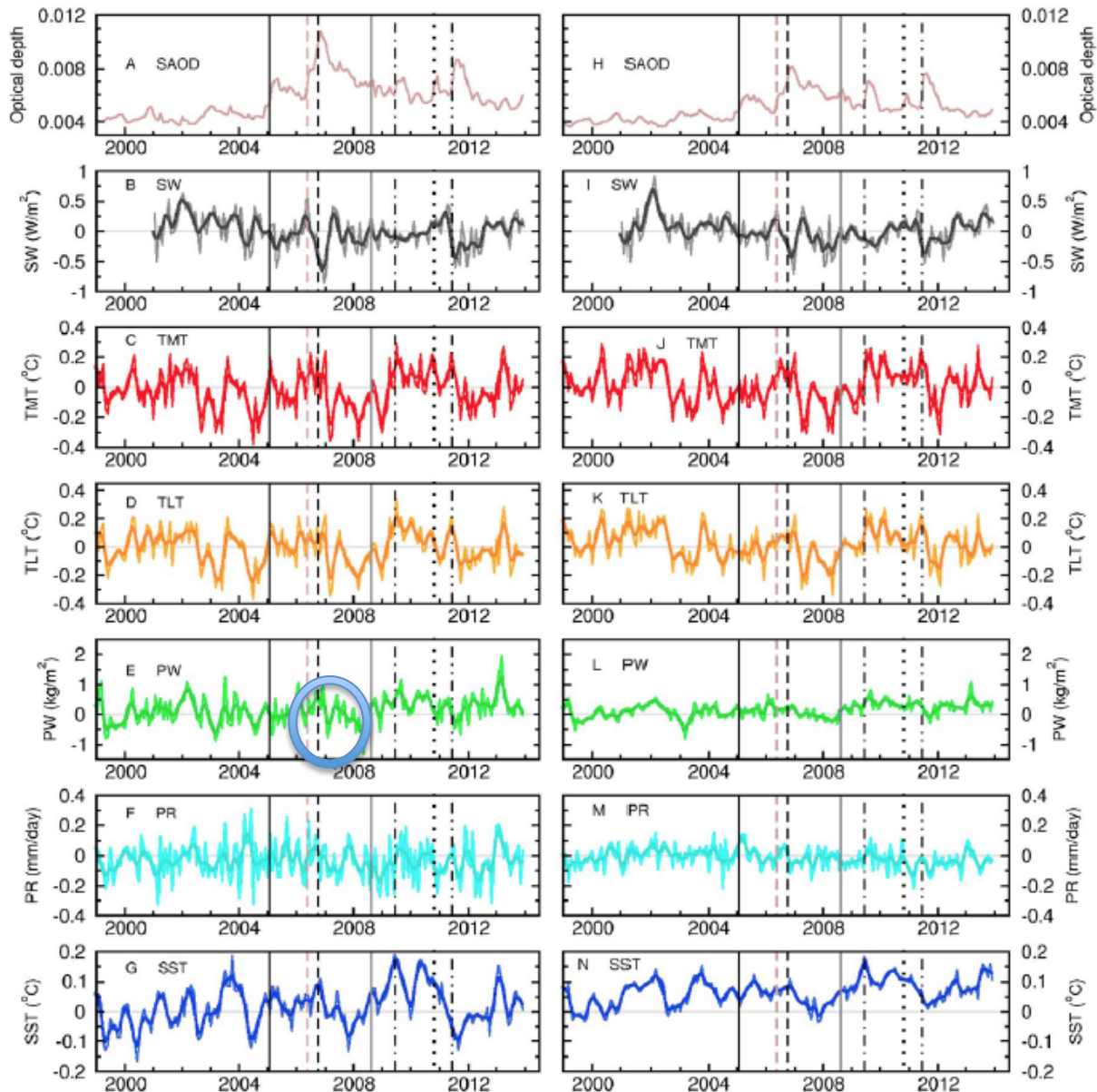
Near-global (50°N-50°S)



Small,
common,
eruptions
have
detectable
signals in
temperature
of lower
troposphere

Tropical (20°N-20°S)

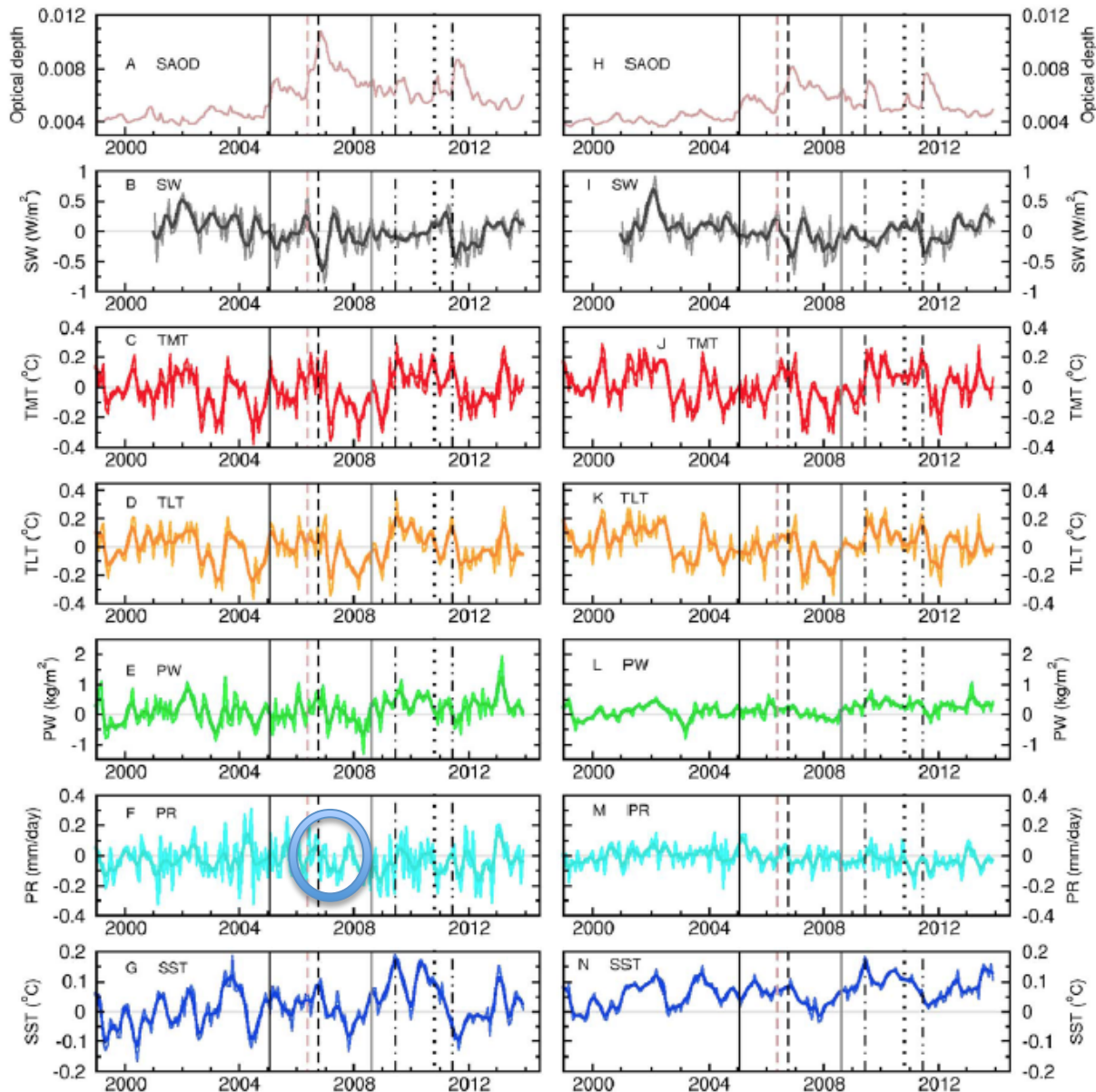
Near-global (50°N-50°S)



Small, common, eruptions have detectable signals in temperature of column water vapor

Tropical (20°N-20°S)

Near-global (50°N-50°S)



— Manam
- - - Soufriere Hills
... Tavorur
- · - Kasatochi

- · - Sarychev
... Merapi
- · - Nabro

Small,
common,
eruptions
have
detectable
signals in
precipitation

Santer et al. 2015

Need to know optical properties to compute radiative changes-emphasis on microphysical models

Particle properties to measure

Composition

Size distribution

Number

Mass

Area

Shape

Optical constants

Extinction optical depth

Scattering optical depth

Absorption optical depth

Scattering phase function

Possible ranges

Dust, sulfates

nm to tens of microns

1 to 10^4

Spheres/fractals

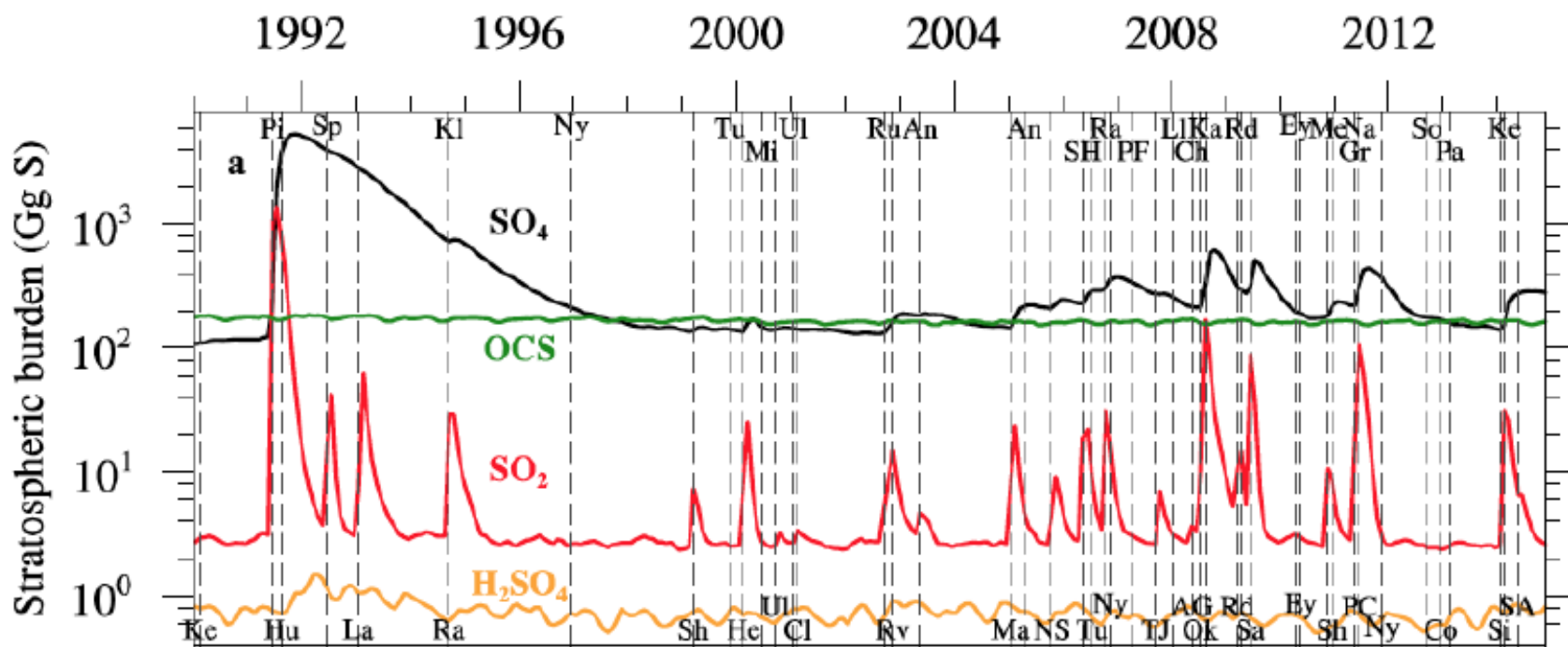
Dust

0.001 to 1

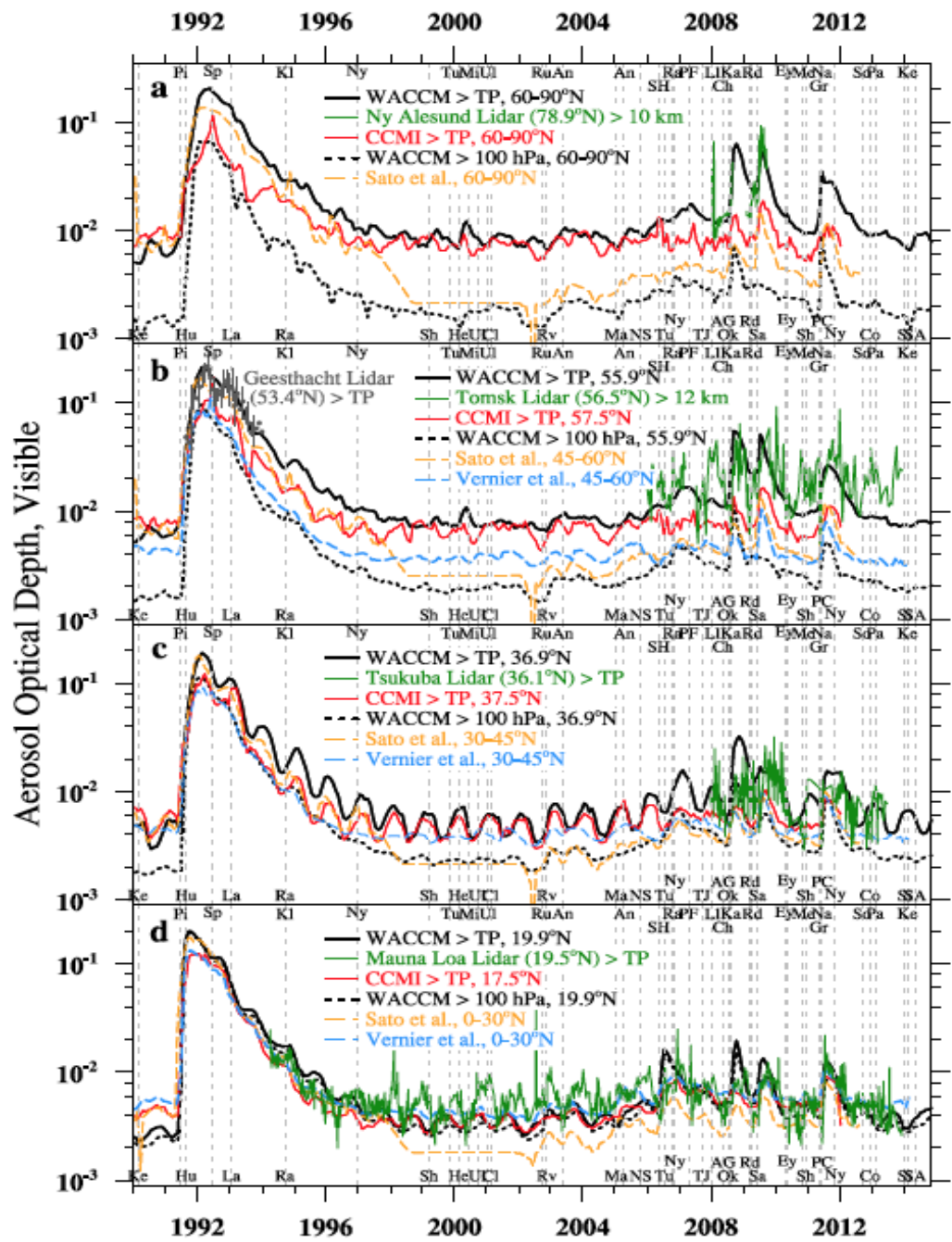
0.001 to 1

0.001 to 1

Microphysical models can predict sulfur chemistry

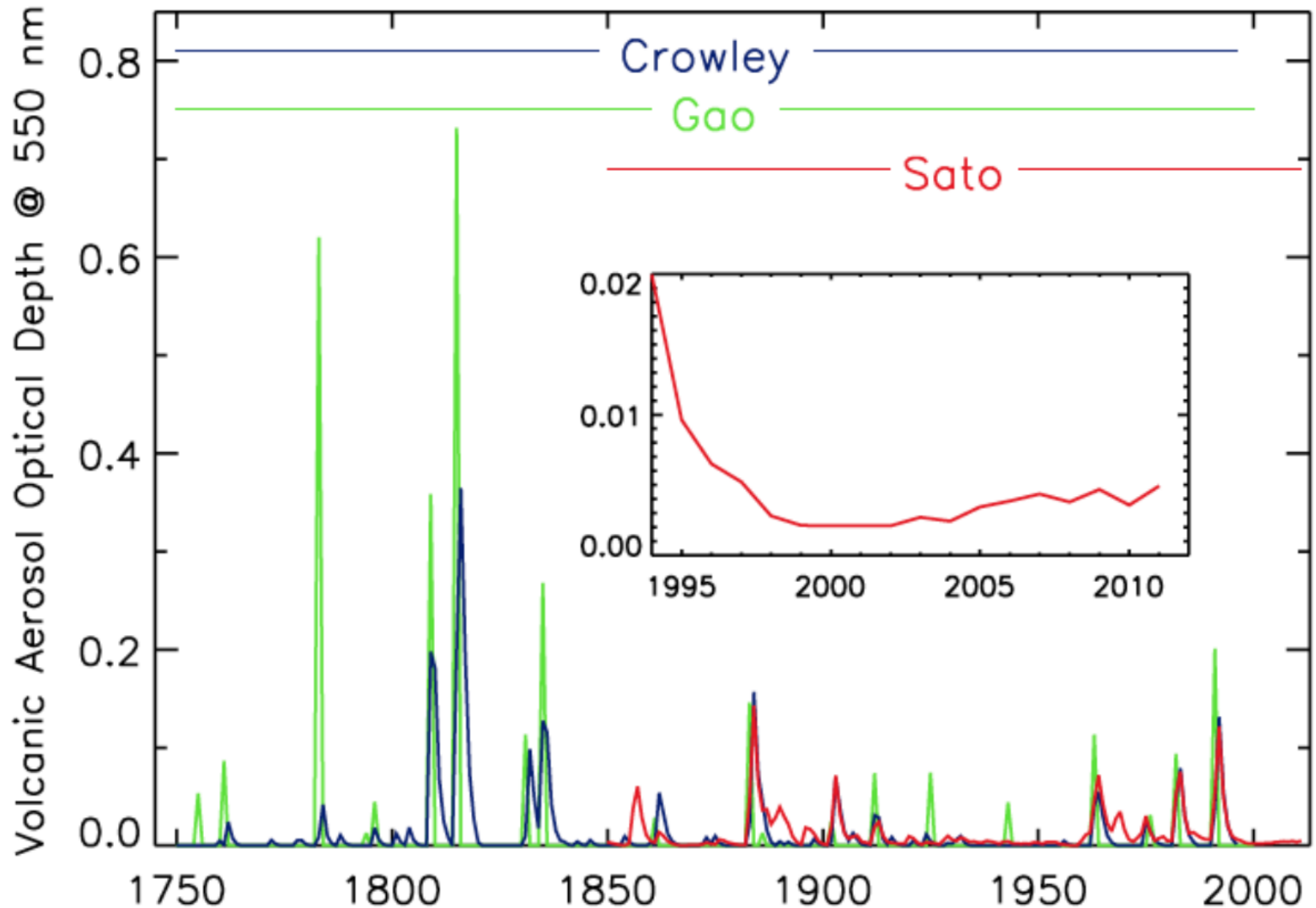


Models can reproduce the optical depth given SO₂ from observations



Mills et al., 2016

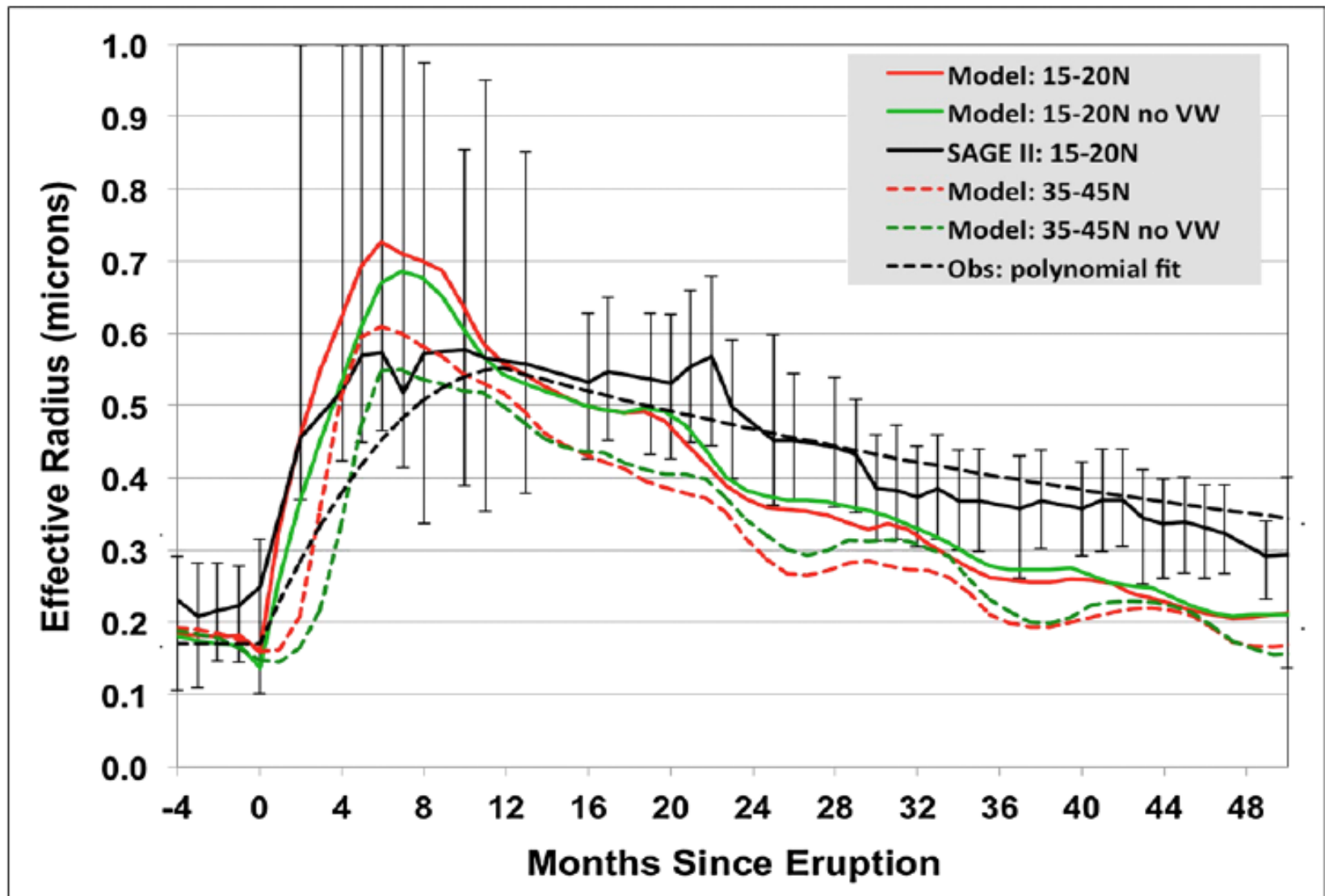
Geologists are not using microphysical models to reconstruct optical depth, what is error source?



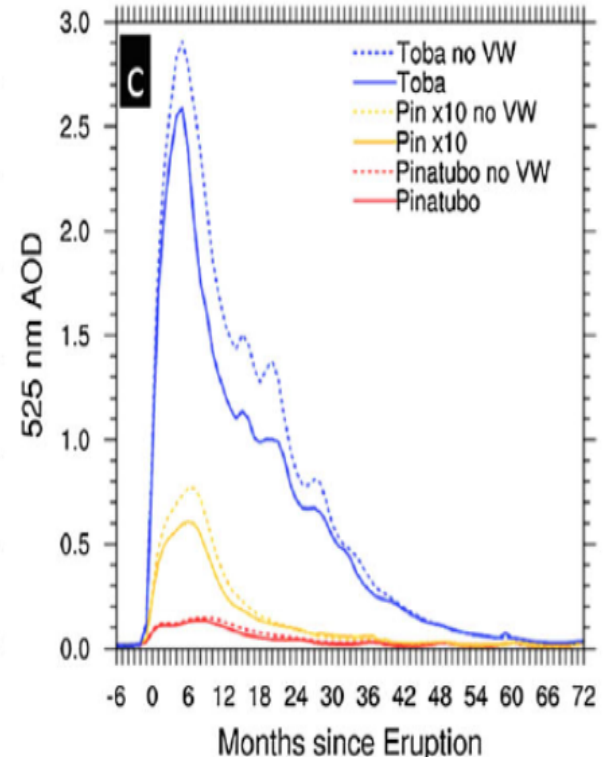
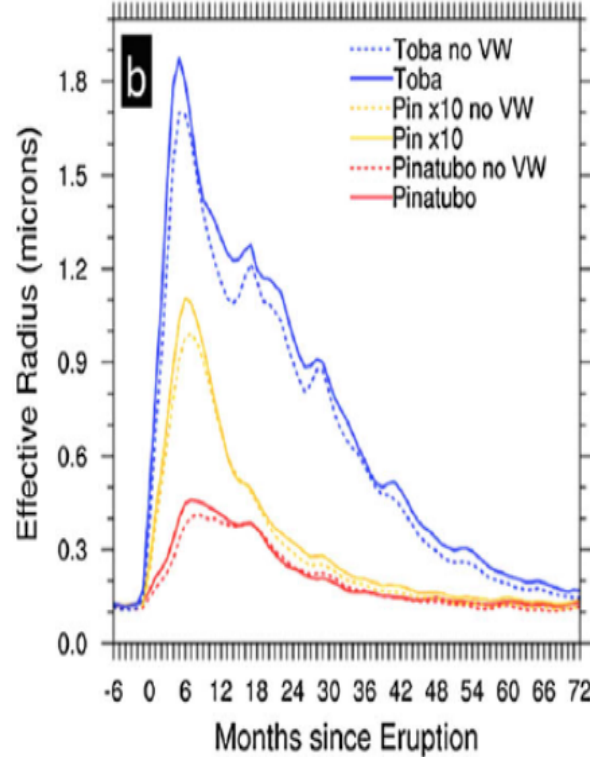
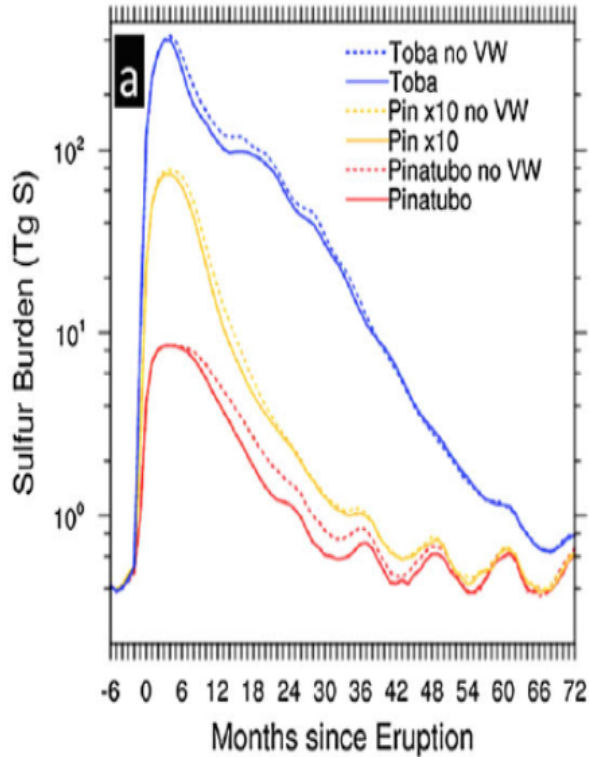
Effective radius is a measure of the particle size that impacts radiation

$$r_{eff} = \frac{\int r(\pi r^2 n(r) q_{ext}(m, \lambda / r)) dr}{\int \pi r^2 n(r) q_{ext}(m, \lambda / r) dr} \approx \frac{3}{4} \frac{V}{A_c}$$

The effective radius after Pinatubo was not constant in time



Models show optical depth not linear in SO₂ emission or sulfate burden



S in sulfate 37% of S in SO₂
 Pinatubo to 100x Pinatubo

$$R_{\text{eff}} \sim M_{\text{SO}_2}^{1/3}$$

$$\text{AOD} \sim M / R_{\text{eff}}$$

Peak AOD ~ 20% of increase
 in S Pinatubo to 100x
 Pinatubo

Reff not
constant
because
coagulation
limits
number

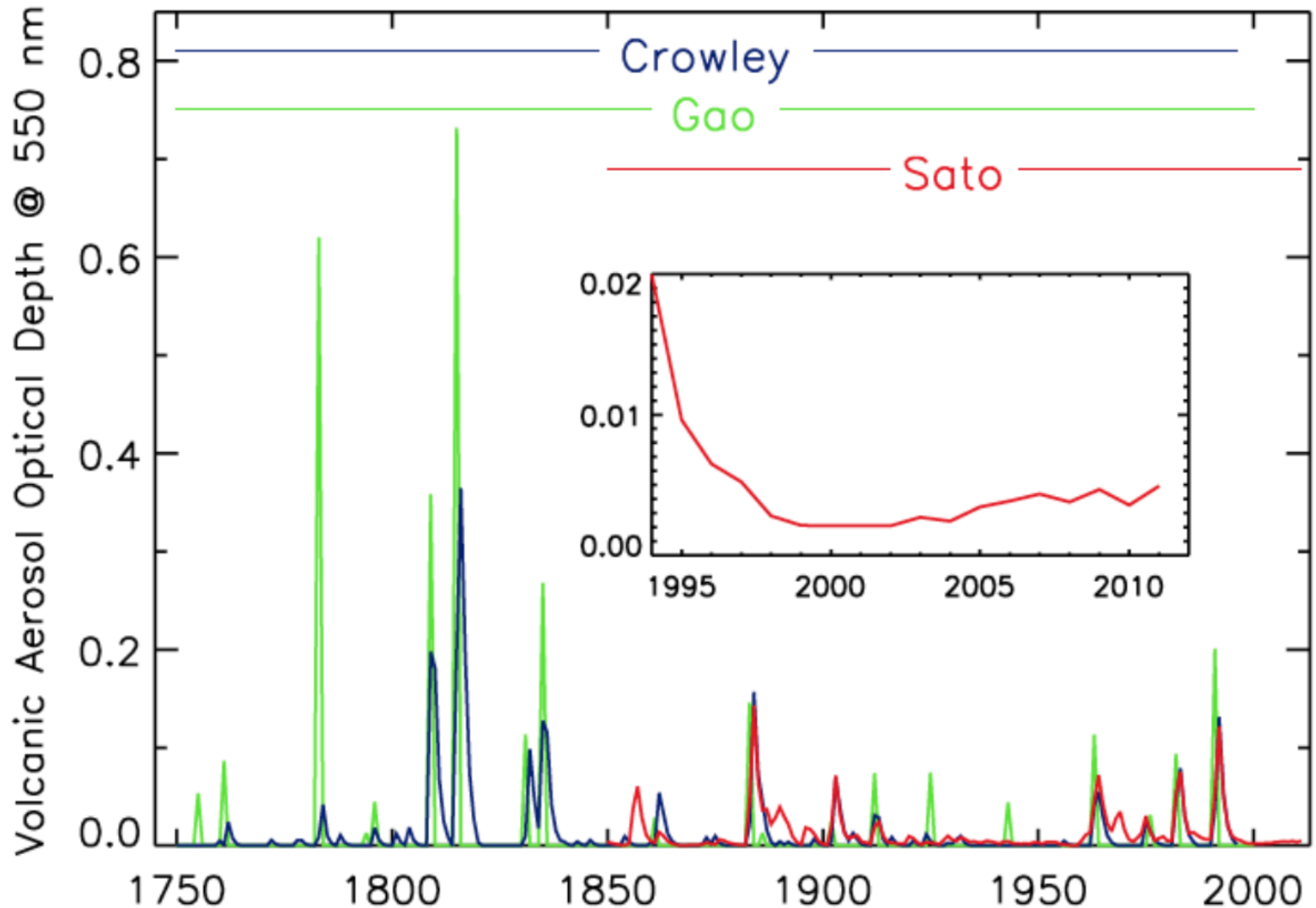
$$\frac{dN}{dt} = -\frac{1}{2} KN^2$$

$$N = \frac{N_0}{1 + \frac{1}{2} KN_0 t}$$

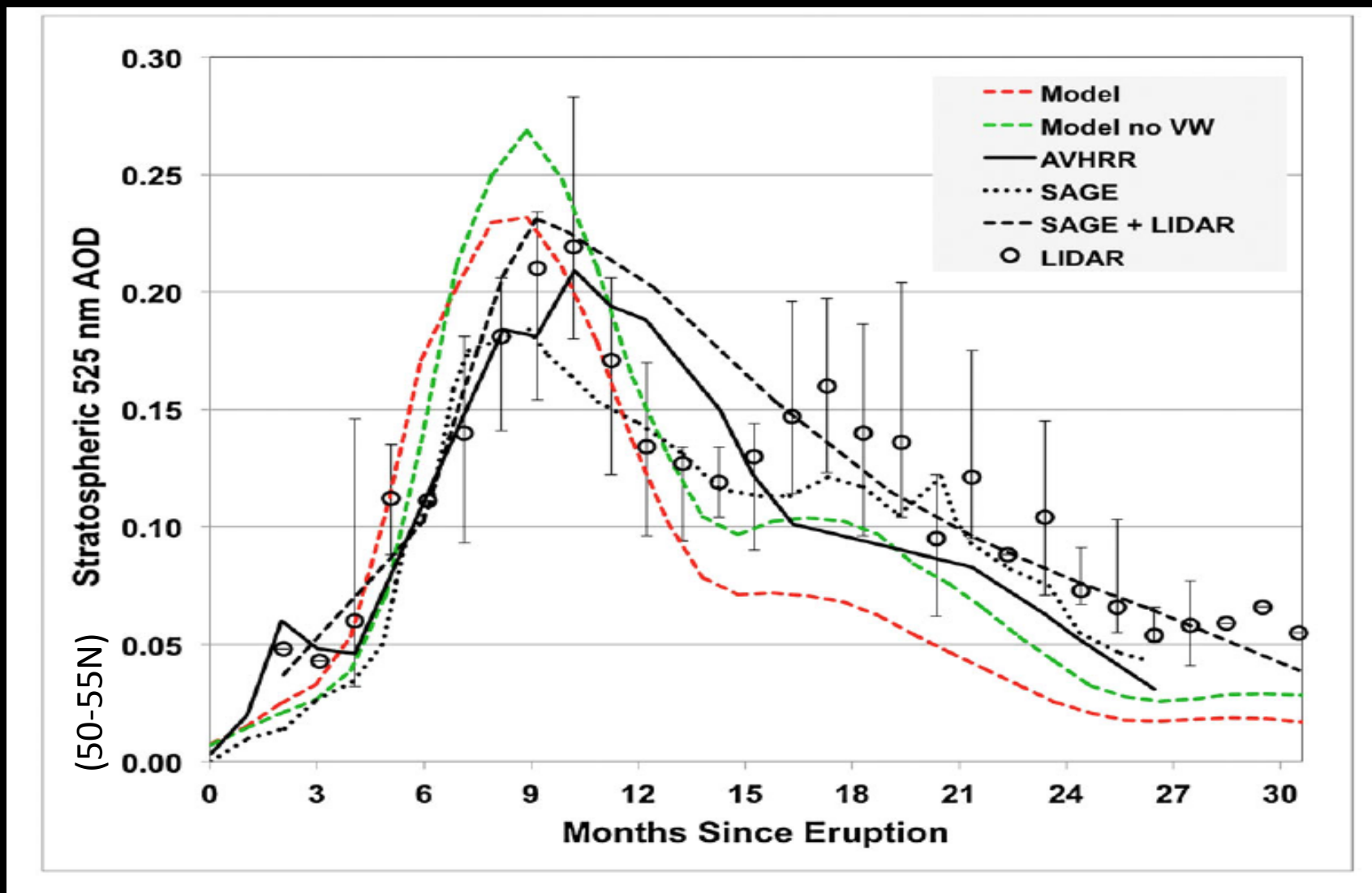
for large time

$$N = \frac{2}{Kt}$$

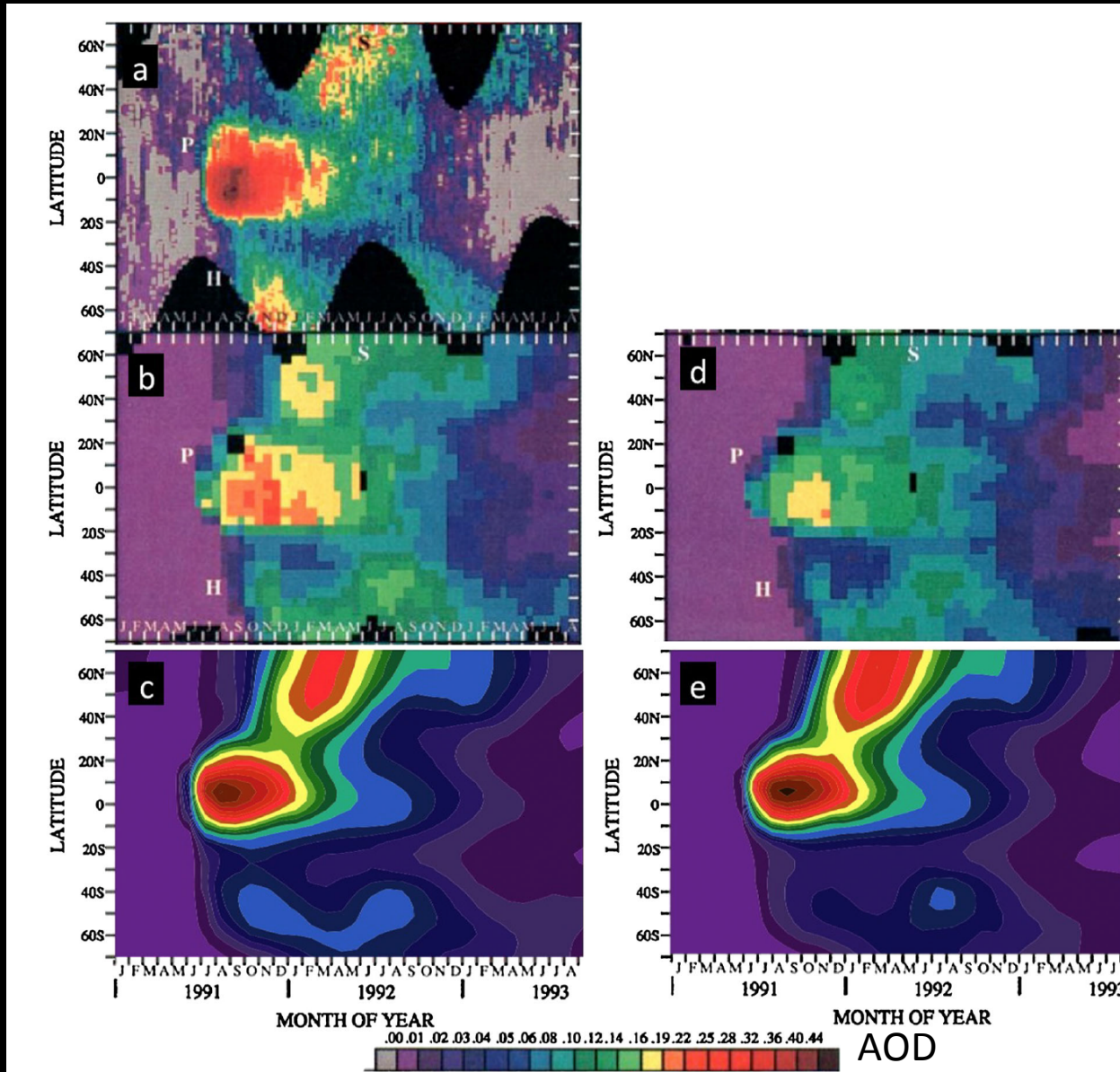
Is a global average optical depth good enough?



Optical depth difficult to predict at fixed location, data have errors, but models have some skill.



The optical depth varies in space and time. A global average makes no sense.

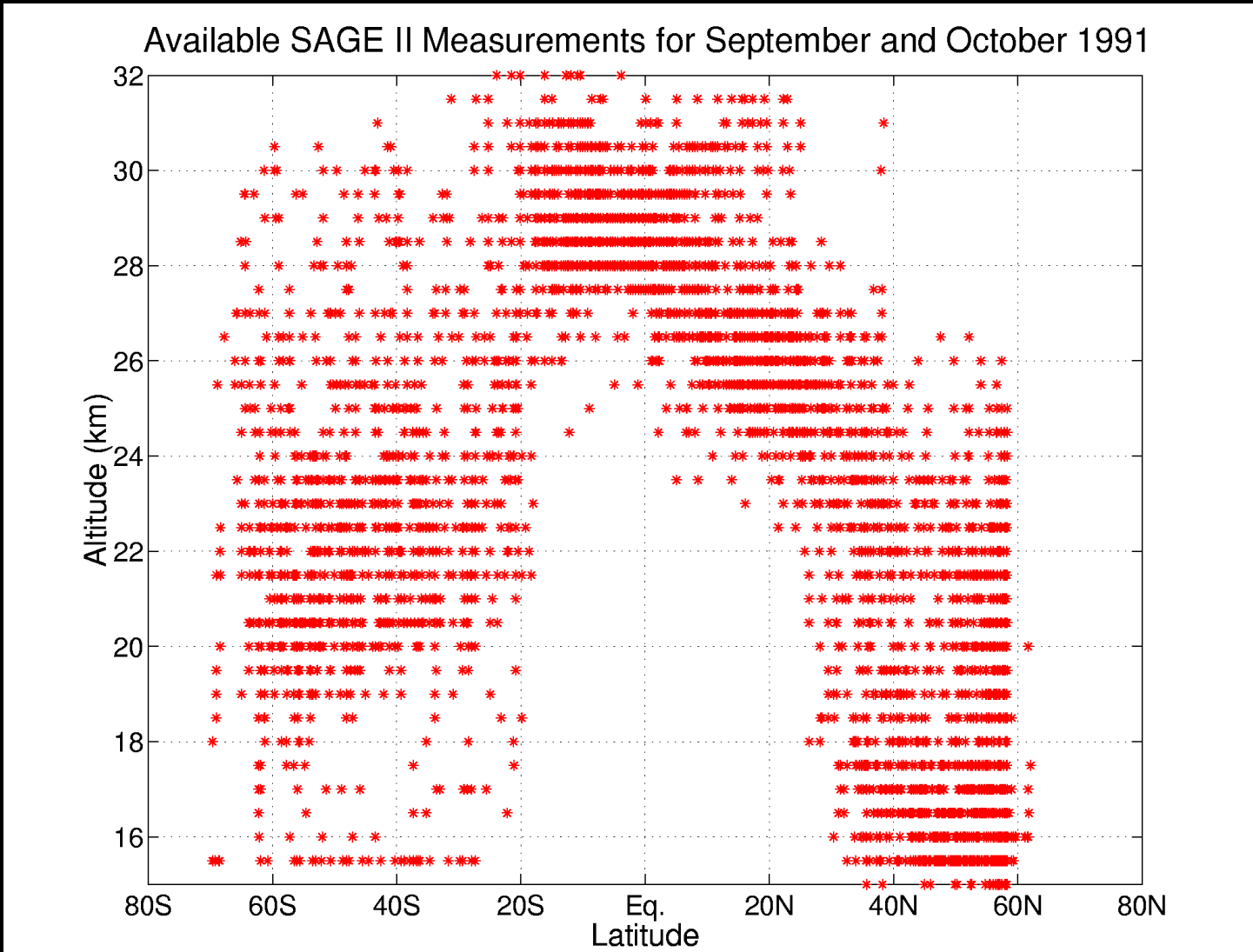


AVHRR, 500nm

SAGE II

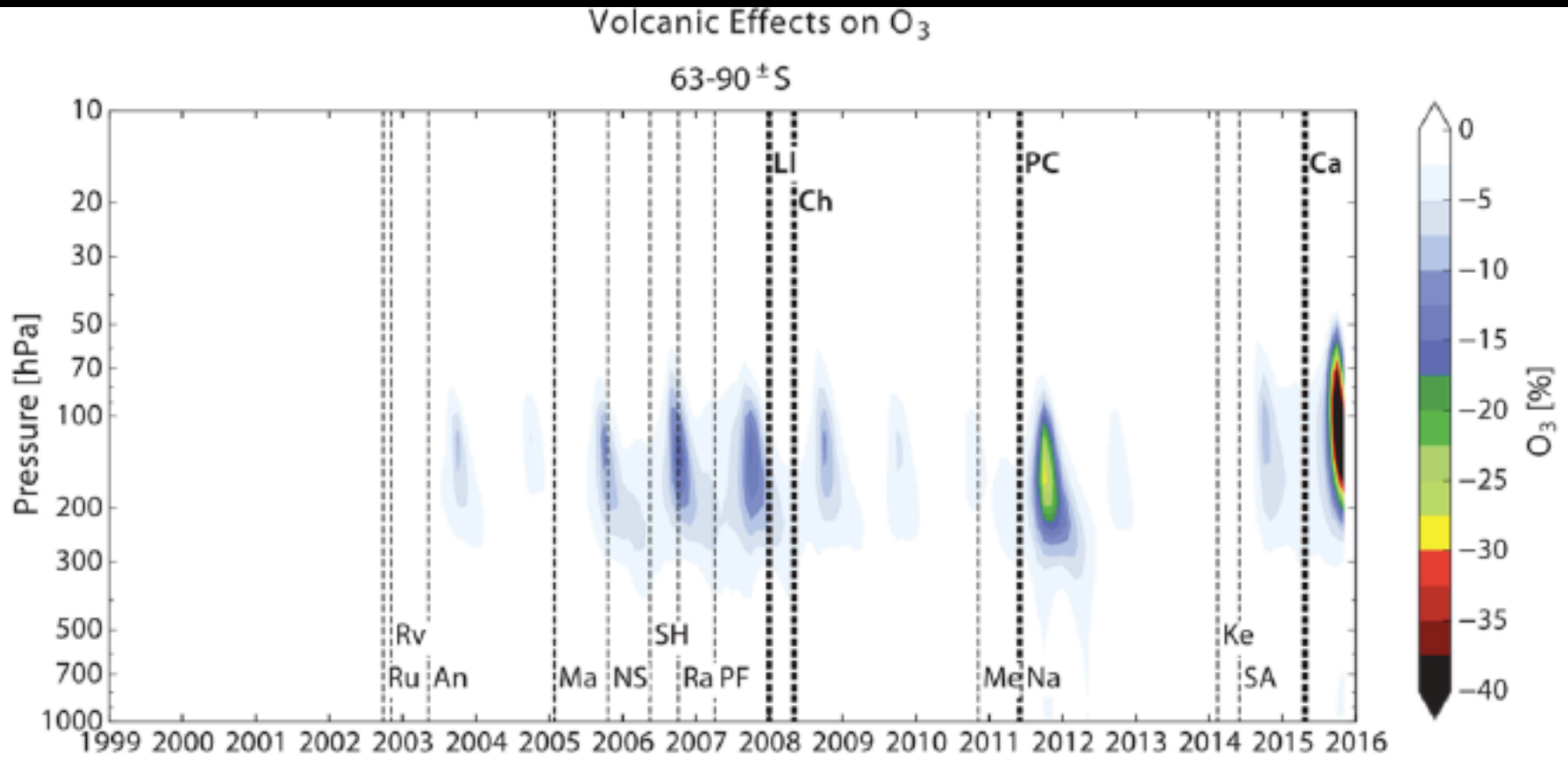
Model 525, 1024nm
English et al., 2013

Limb sounders are blocked by Pinatubo sized events

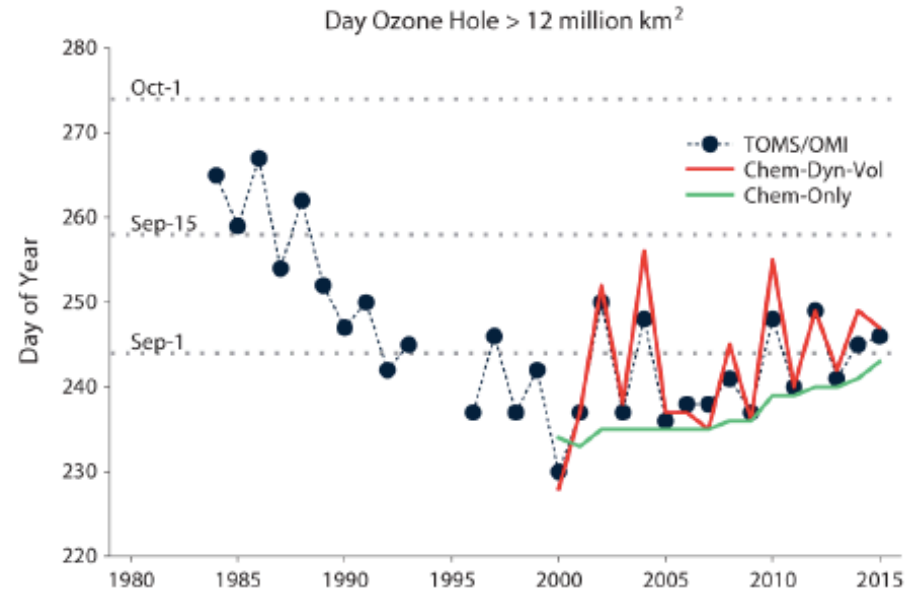
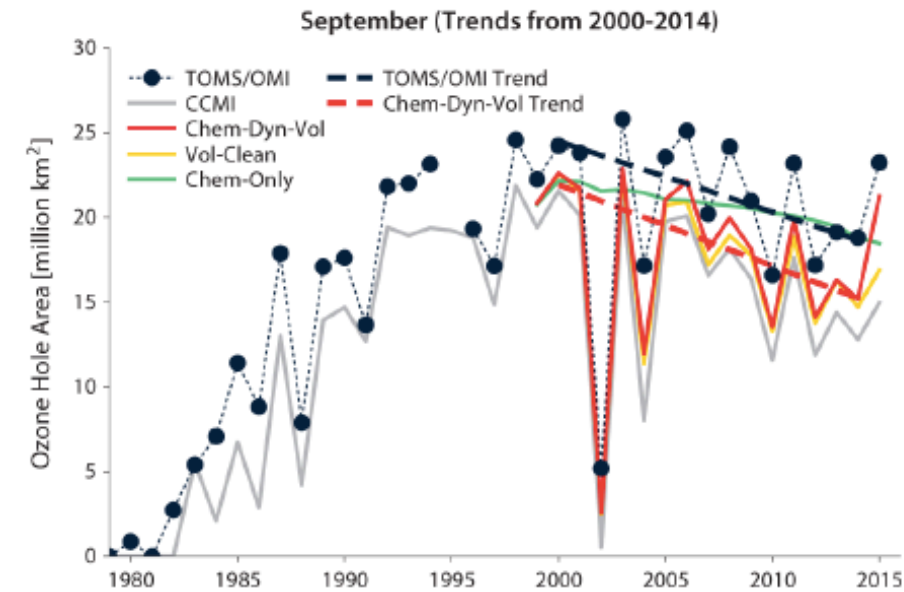


Plot by Juan Carlos Antuña
Courtesy Alan Robock

Small eruptions may cause polar ozone loss

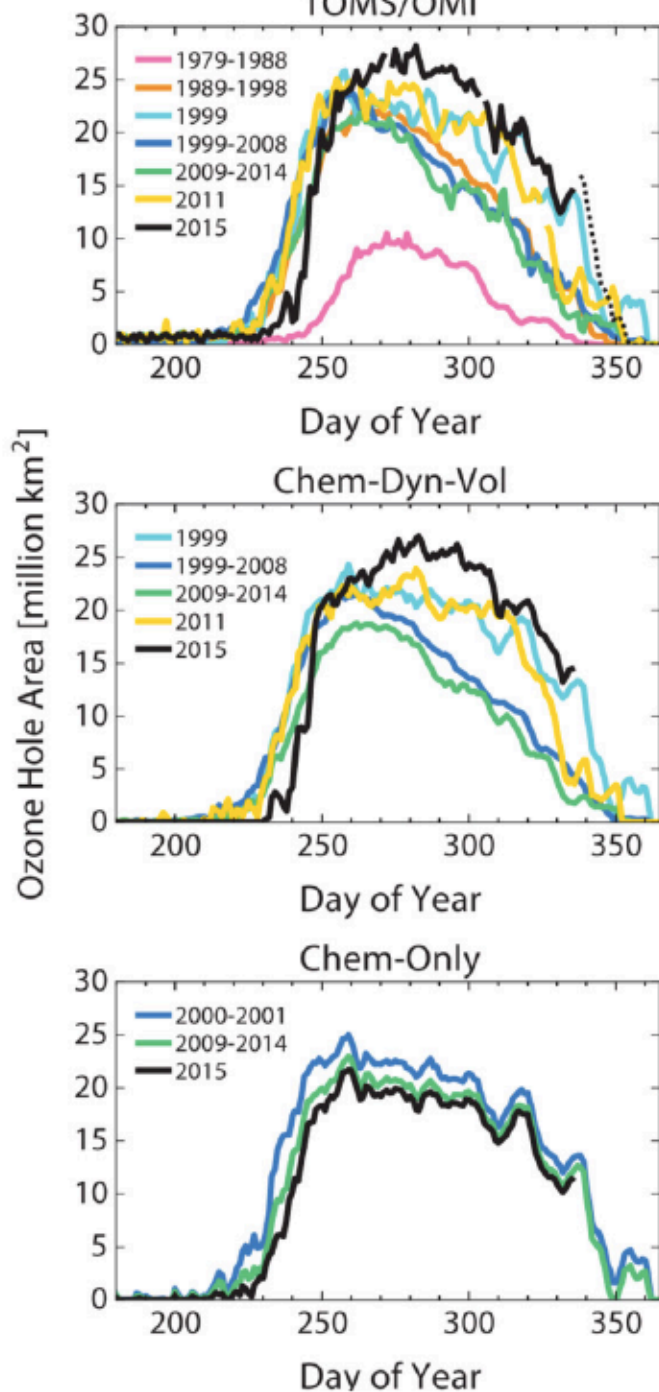


The ozone hole is recovering, but eruptions are delaying it



Solomon et al., 2016

Volcanic eruptions
may cause the ozone
hole to expand, but it
is starting later in the
year



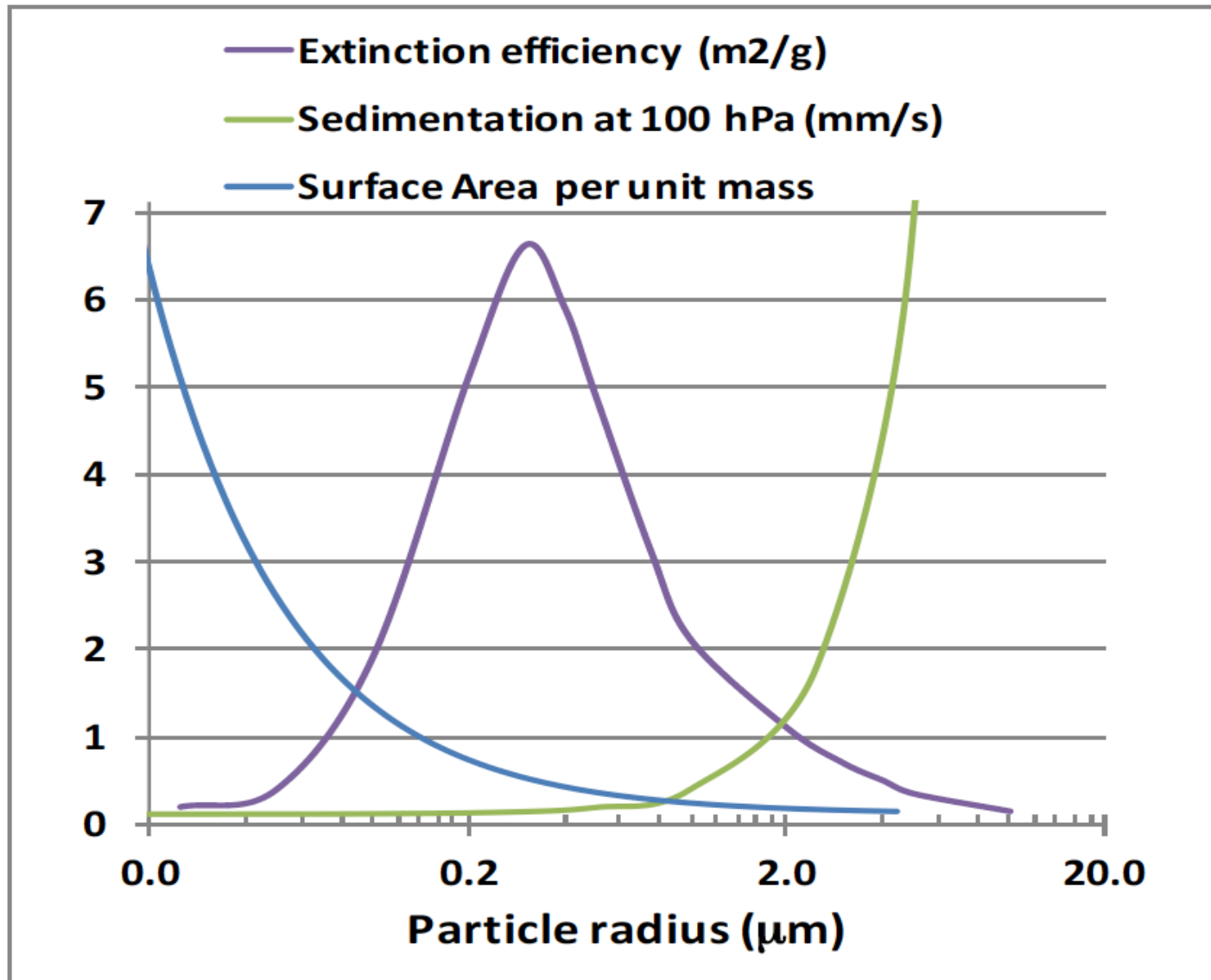
Summary

- Global average optical depths don't correspond to reality.
- The optical depth depends on SO_2 injected not explosivity.
- Optical depth is less than linear with SO_2 mass injected.
- Recovery of the ozone layer retarded by small eruptions
- Large error bars on observed properties of historical volcanic clouds.
- We need more data and we need to be prepared to get it. Small eruptions are important to study. Better models are needed.

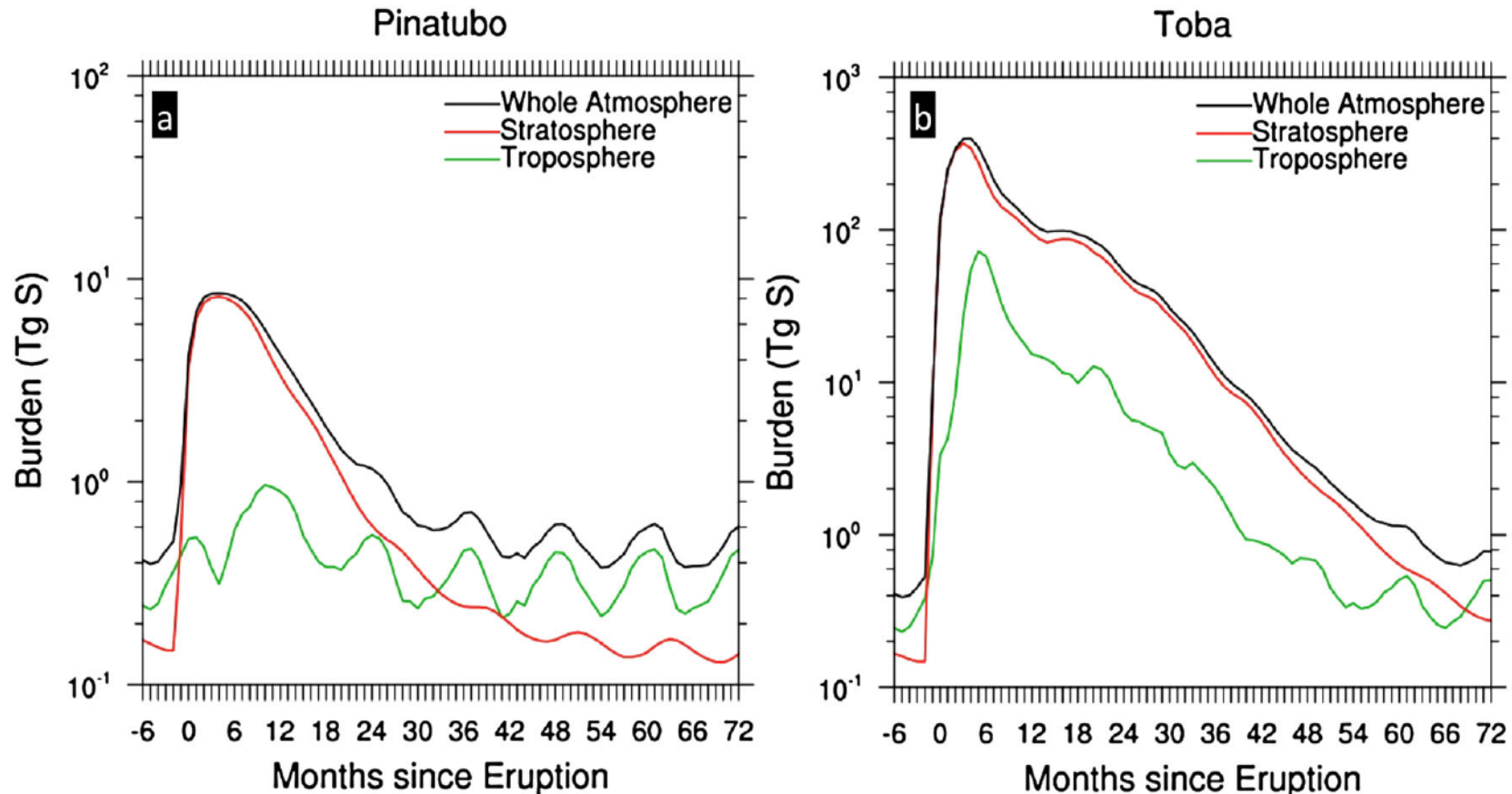
What could you do?

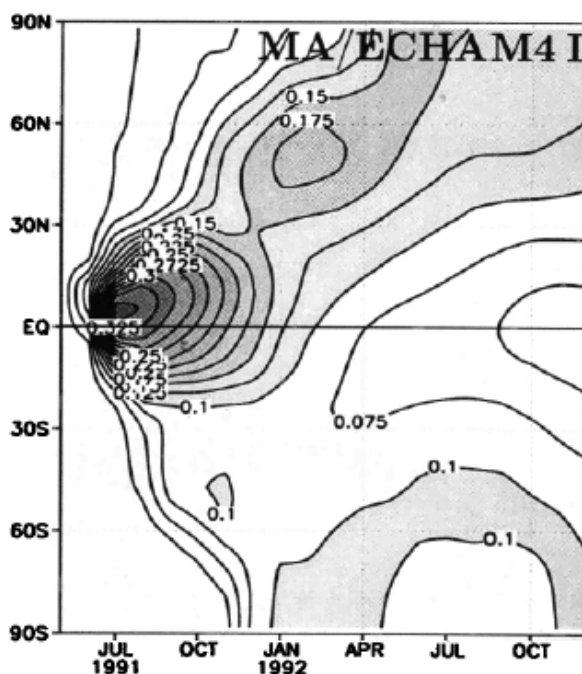
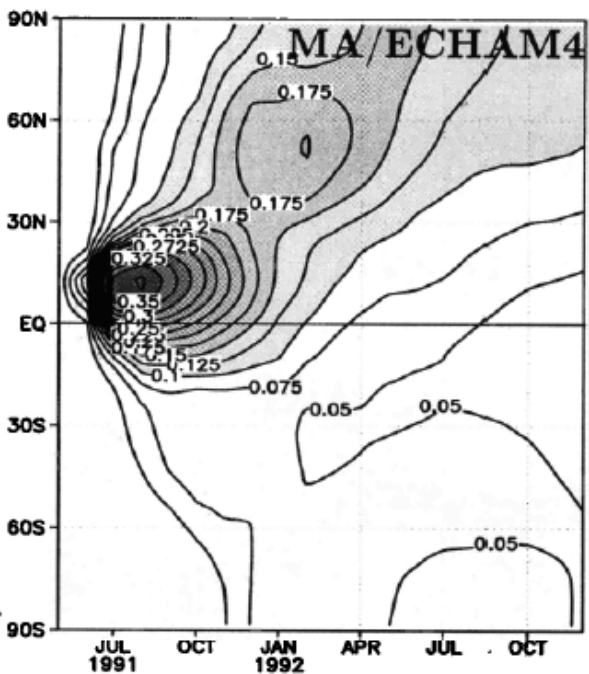
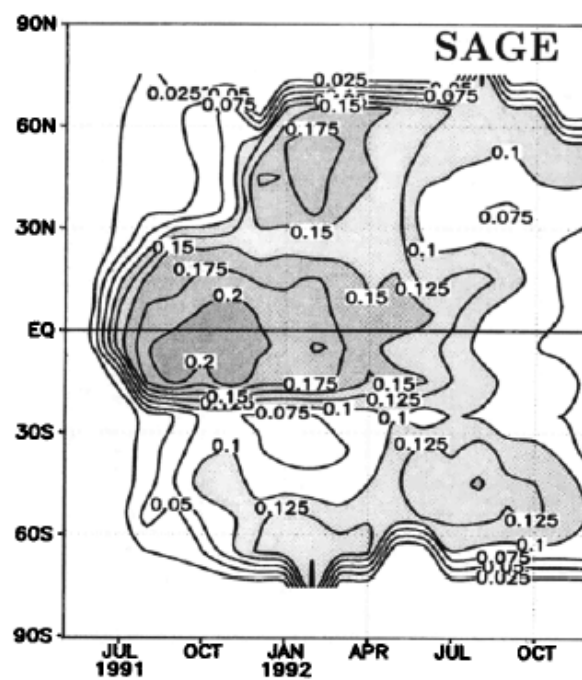
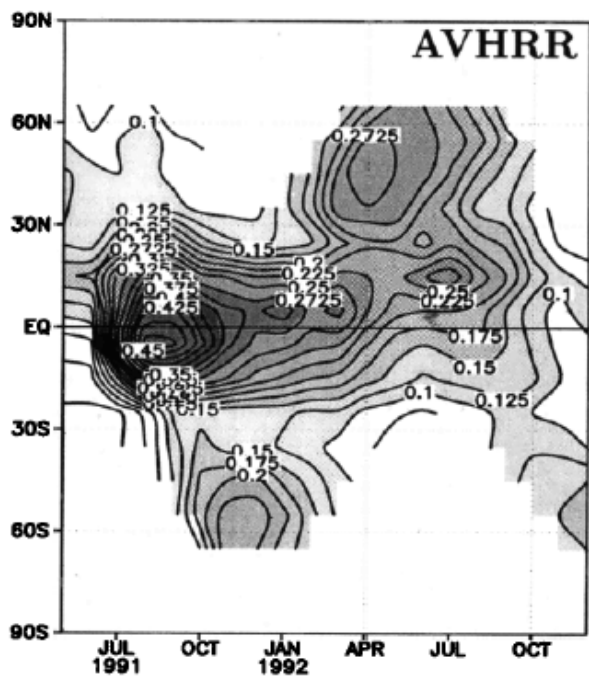
- Formulate plan for national response to measure changes after next eruption.
- Consider how to better determine importance of historic eruptions.
- Make improved models/ observations of volcanic effects on ozone, climate, cirrus

Reff is not enough to compute physics



A lot of sulfate can be present in the troposphere





Radiative heating in Pinatubo cloud affects its spatial distribution

Timmreck et al., 1999