Satellite-derived warm rain fraction as constraint on the cloud lifetime effect

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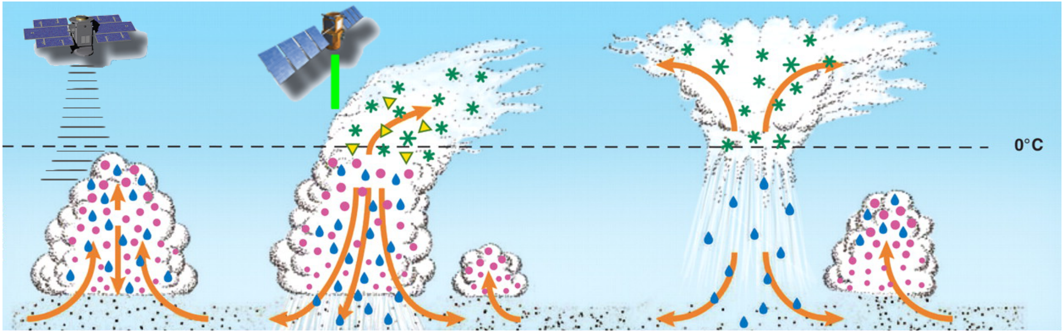
20 September 2016



Precipitation

High radar reflectivity of rain drops

→ CloudSat CPR via 2C-PRECIP-COLUMN *or* DARDAR_MASK

Liquid-topped clouds

High lidar backscatter at cloud top from liquid droplets

→ CALIOP via DARDAR_MASK

Ice clouds

High radar reflectivity of ice particles

→ CPR via DARDAR_MASK

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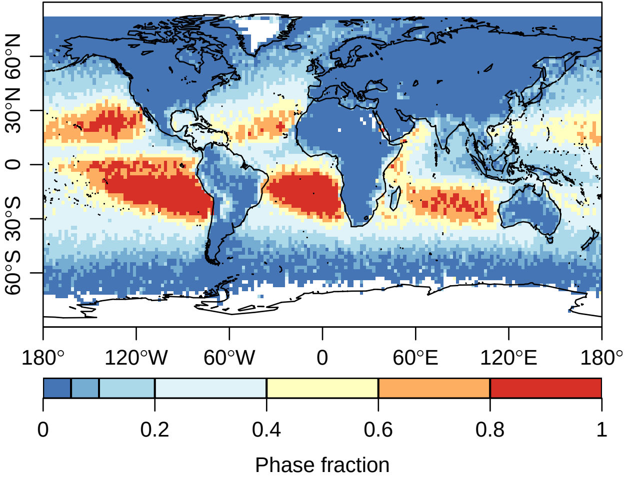
after Rosenfeld et al. (2008), *Science*

Rain from pure liquid clouds ("warm rain") is very rare over the extratropical continents



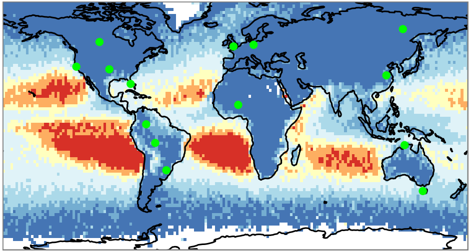
Mülmenstädt et al. (2015), *Geophys. Res. Lett.* 42 (15), 6502–6509, doi:10.1002/2015GL064604

AeroCom project proposal

- Aerosol influence mainly acts on autoconversion in liquid-water clouds in current models
- The more precipitating warm clouds are simulated in a model, the more opportunity aerosols have to influence the precipitation microphysics
- We hypothesize that the strength of the cloud lifetime effect in models is therefore related to the warm-rain fraction
- This hypothesis can be tested in the AeroCom models
- Comparing warm-rain fraction in models against satellites may provide an observational constraint on the cloud lifetime effect

Outline

Compare satellite climatology to CMIP5 cfSites



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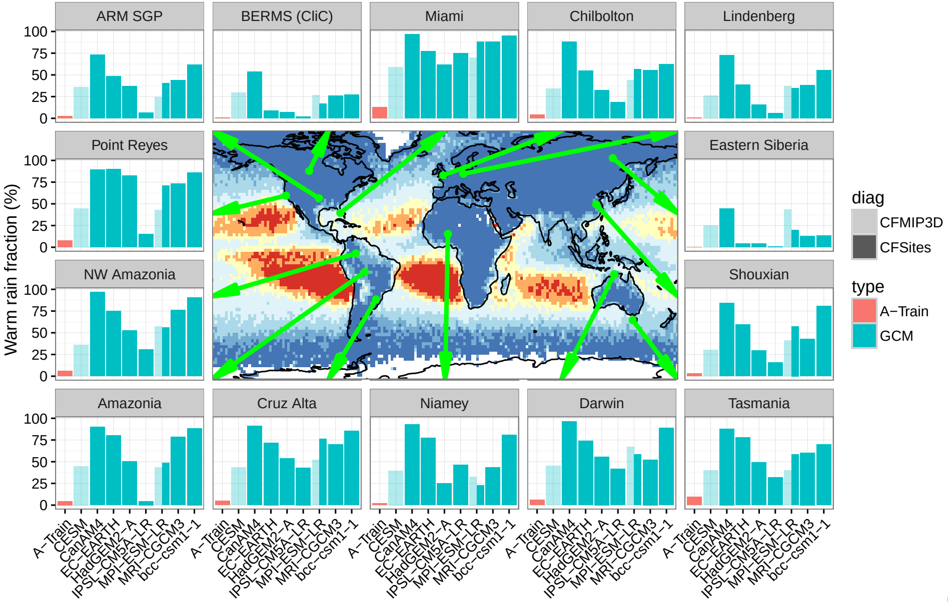
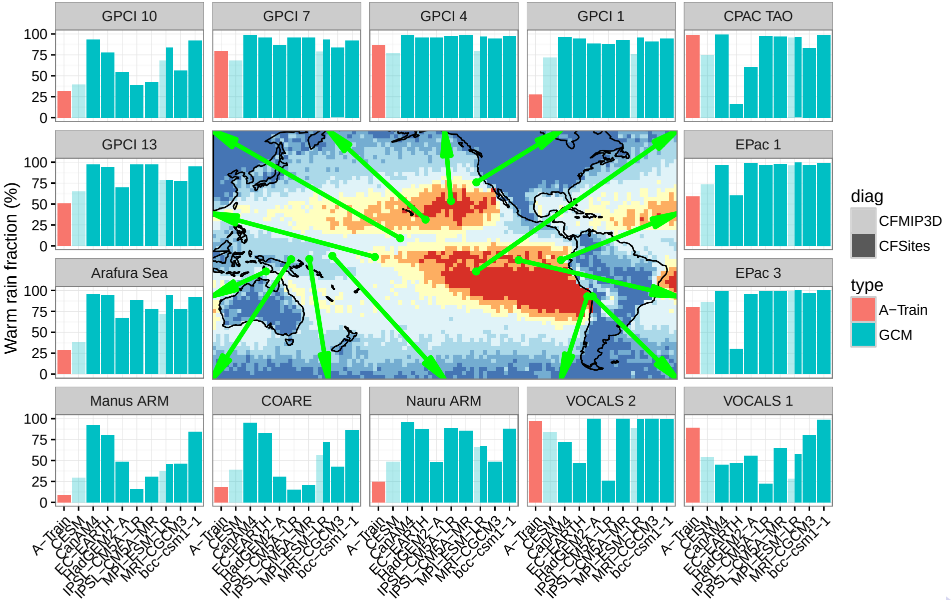
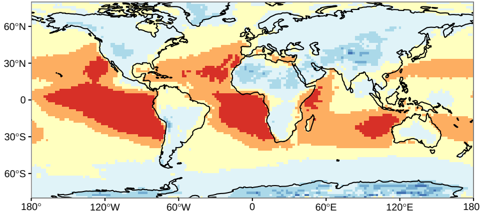


Compare satellite climatology to CMIP5 cfSites

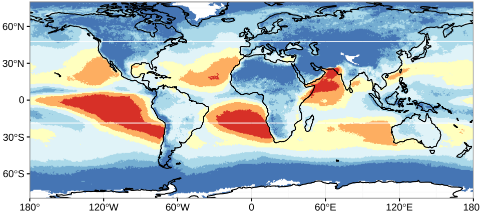


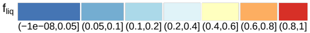
Modeled warm-rain fraction is diverse

ECHAM–HAM

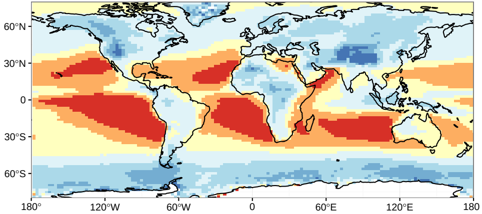


SPRINTARS

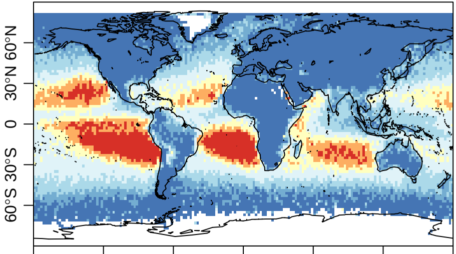




CAM5–CLUBB



Satellite



Outline

Motivation

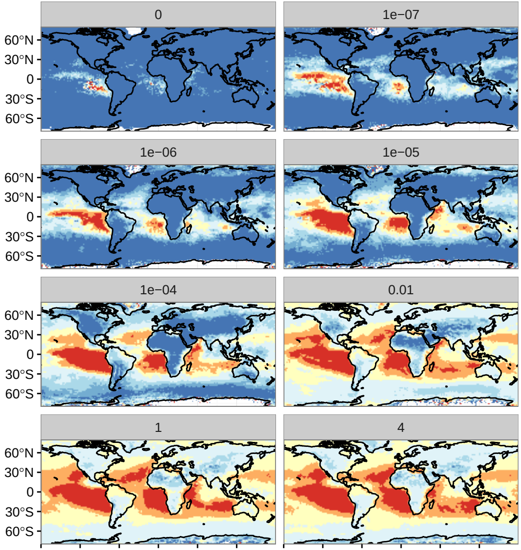
Warm-rain fraction in observations and GCMs

Tuning the warm-rain fraction in ECHAM–HAM

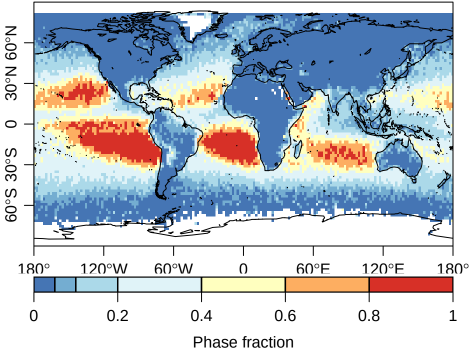
Interactions between the warm-rain fraction and ERF_{aci}

Scale factor on autoconversion rate: $10^{-4} \times Q_{aut}$ reproduces observations

KK(2000) autoconv with scale factor

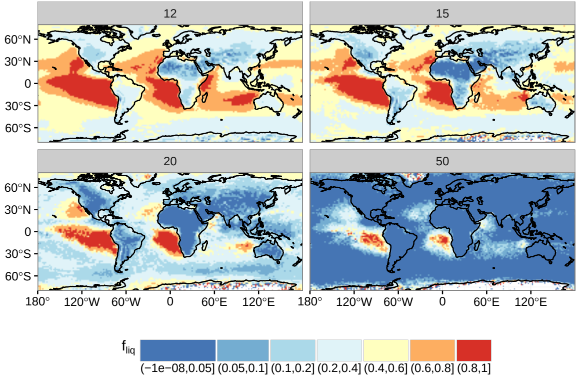


Satellite

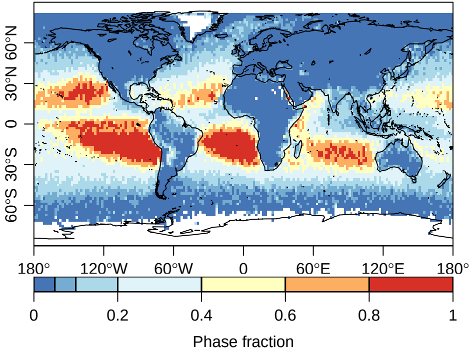


Threshold on autoconversion: $r_e > 20\ \mu m$

KK(2000) autoconv with r_e threshold



Satellite



These modifications are related

Khairoutdinov and Kogan (2000):

$$\frac{\partial q_r}{\partial t} \propto q_l^{\alpha} N^{\beta}, \quad \alpha = 2.47, \beta = -1.79 \tag{1}$$

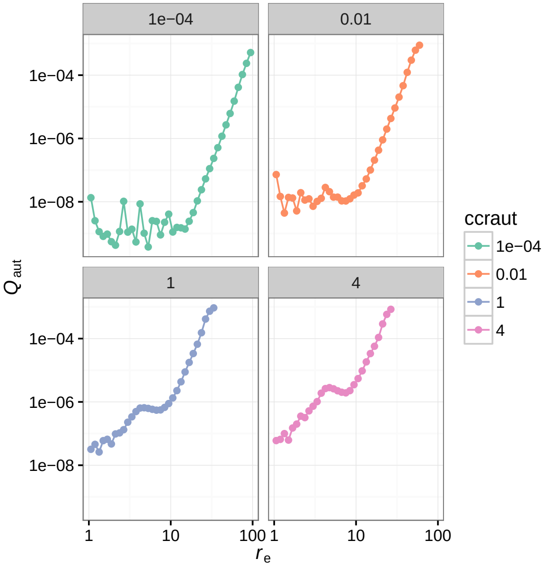
Since

$$q_l \propto r_e^3 N \tag{2}$$

the autoconversion rate can be rewritten as a function of r_e and either of q_l or N:

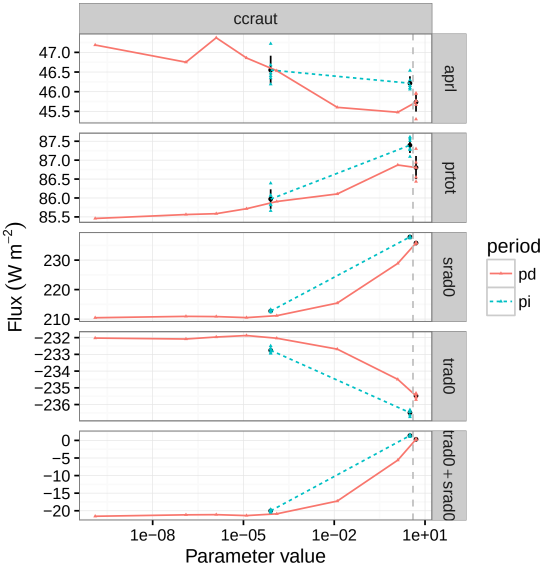
$$\frac{\partial q_r}{\partial t} \propto \begin{cases} r_e^{3\alpha} N^{\alpha+\beta} \\ r_e^{-3\beta} q_l^{\alpha+\beta} \end{cases} \tag{3}$$

Under the simplifying assumption that r_e is uncorrelated with either of q_l or N, we expect the autoconversion rate to scale with $r_e^{5.5 \sim 7.5}$, which effectively sets an r_e threshold.



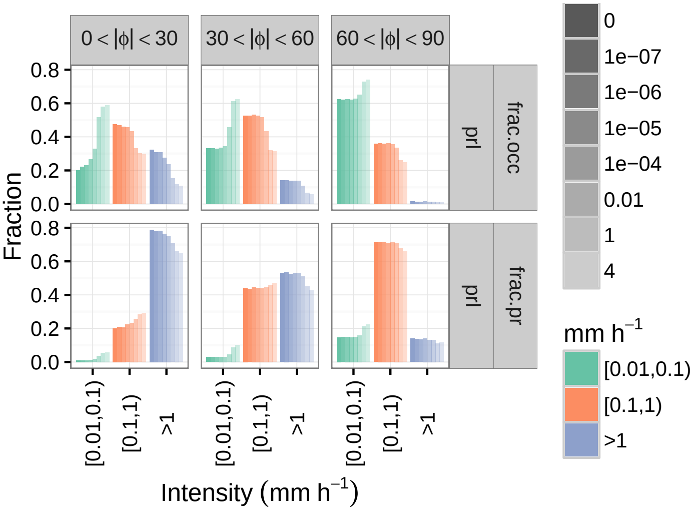
Effect on energy fluxes

- Reducing the warm-rain fraction significantly detunes the TOA energy balance $\rightarrow$ retuning is required (primarily SW)
- (Reducing warm-rain fraction increases large-scale precipitation)



Effect on precipitation intensity distribution

- Reducing the warm-rain fraction also increases the intensity spectrum
- Shown here are large-scale precipitation intensity spectra at different latitude bands
- Decreasing the warm-rain fraction increases the probability of intense large-scale precipitation



Tuning the warm rain fraction in ECHAM–HAM: conclusions

- Satellite warm-rain fraction can be reproduced in ECHAM–HAM by multiplying the Khairoutdinov and Kogan (2000) autoconversion rate by 10^{-4} (default ECHAM–HAM tuning factor: 4)
- Alternative to this drastic scale factor: $r_e > 20\mu$m threshold on autoconversion
- Effect on radiative balance is large (large increase in cloud lifetime)
- Reducing the warm-rain fraction to match the satellite climatology also increases the intensity spectrum
- (Some remaining uncertainty on these numbers because of parameter choices in diagnosis of warm-rain fraction)

Outline

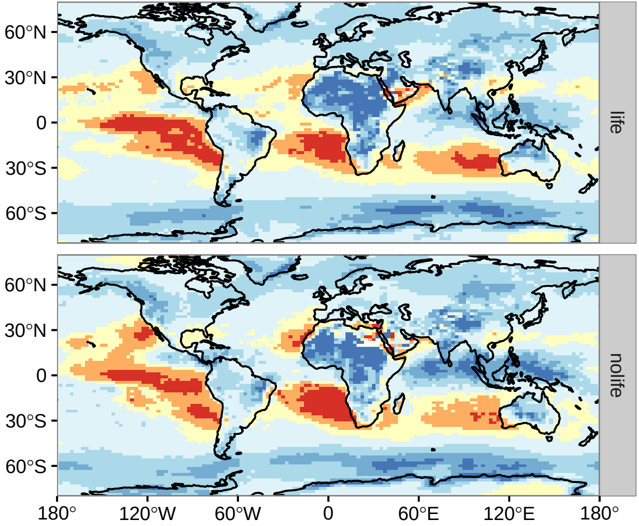
Influence of the warm-rain fraction on ERF_{aer}

Results for ECHAM6.1–HAM2.2, AeroCom II 1850/2000 emissions

`ccraut`	SW PD − PI (W m^{-2})	LW PD − PI (W m^{-2})	SW + LW PD − PI (W m^{-2})
4 (default)	−2.1	1.0	−1.1
10^{-4}	−1.6	0.72	−0.86

- As hypothesized, the configuration with lower warm-rain fraction has a smaller ERF_{aer}
- The change is −0.5 W m^{-2} SW offset by 0.3 W m^{-2} LW ⇒ plausible that ERF_{aci} change is a large contribution
- (Low-`ccraut` configuration has not been retuned and ERF_{aci} has not been diagnosed separately from ERF_{aer} yet)

Influence of the lifetime effect on warm-rain fraction



- CAM5 runs with and without cloud lifetime effect
- In SE and NE Pacific and Atlantic, lifetime effect decreases the warm-rain fraction, as expected from drizzle suppression
- However, there are also regions where the warm-rain fraction decreases
- Results are very preliminary (still based on non-standard diagnostic algorithm while some more files transfer)

Preliminary conclusions on the relationship between warm-rain fraction and aerosol effects

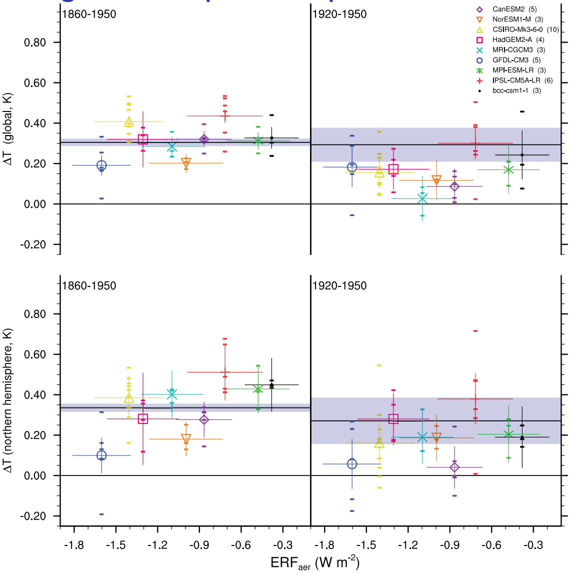
- Changing the warm-rain fraction (in ECHAM–HAM) changes the ERF_{aci}
  $\Rightarrow$ As anticipated, warm-rain fraction is sensitive to aerosol effects

- Lots of model diversity; this observable has not been tuned to death
  $\Rightarrow$ May be useful as an observational constraint

- Next step: investigate relationship between warm-rain fraction and ERF_{aci} across models

- Participation by other models welcome!
  $\Rightarrow$ Required output: snow and rain mixing ratio/flux/path, non-accumulated field, ideally 3h; preferably for a model configuration with known ERF_{aci}

Change of subject: response to Stevens (2015)



Kretzschmar et al. (2016), submitted to *J. Climate*

- Stevens (2015): zero-dimensional global-mean aerosol forcing model with linear ARI and logarithmic ACI terms based on sulfate aerosol
- In this model, ≈50% of ERF_{aer} is already realized in 1950; warming in the early 20th century constrains present-day ERF_{aer}
- In CMIP5 models, the ACI saturates less quickly due to transport from polluted to pristine regions (Rotstayn et al., 2015)
- Realized ERF_{aer} is only ≈25% in 1950
- No strong constraint on ERF_{aer} from early-20th-century warming

Summary

- Warm-rain fraction is very low over continents (especially extratropical NH); details: Mülmenstädt et al. (2015), *Geophys. Res. Lett.* 42 (15), 6502–6509, doi:10.1002/2015GL064604
- Warm-rain fraction can be diagnosed in GCMs and may serve as an observational constraint on precipitation-related processes (including aerosol cloud lifetime effect)
- In ECHAM–HAM, agreement with satellite warm-rain fraction can be achieved with either a drastic rescaling of KK2000 autoconversion or a less drastic r_e threshold
- Either method of tuning the warm-rain fraction intensifies the precipitation intensity spectrum and decreases the ERF_{aci}