

# Model Intercomparison of Aerosol Effects on Liquid and Mixed-phase clouds

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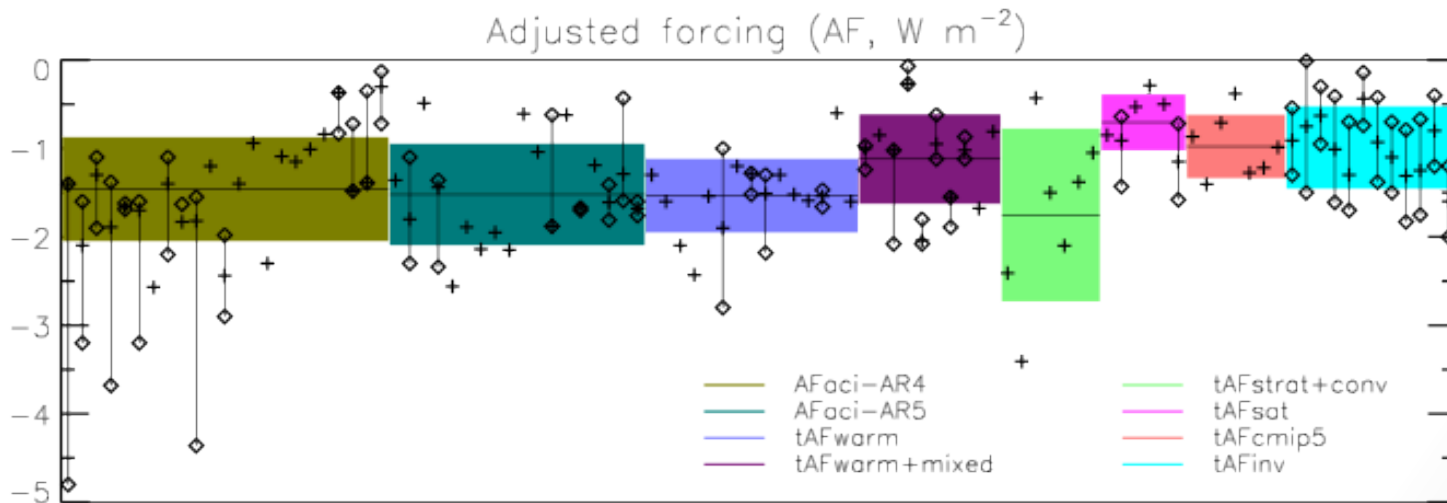
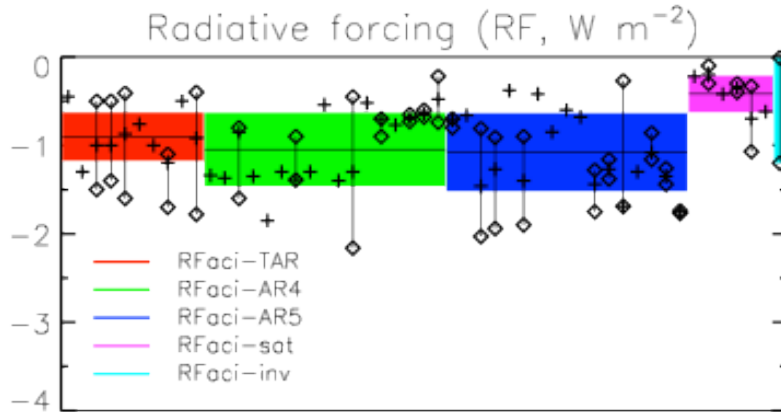
U. Lohmann, ETH-Zurich

T. Takemura, University of Tokyo

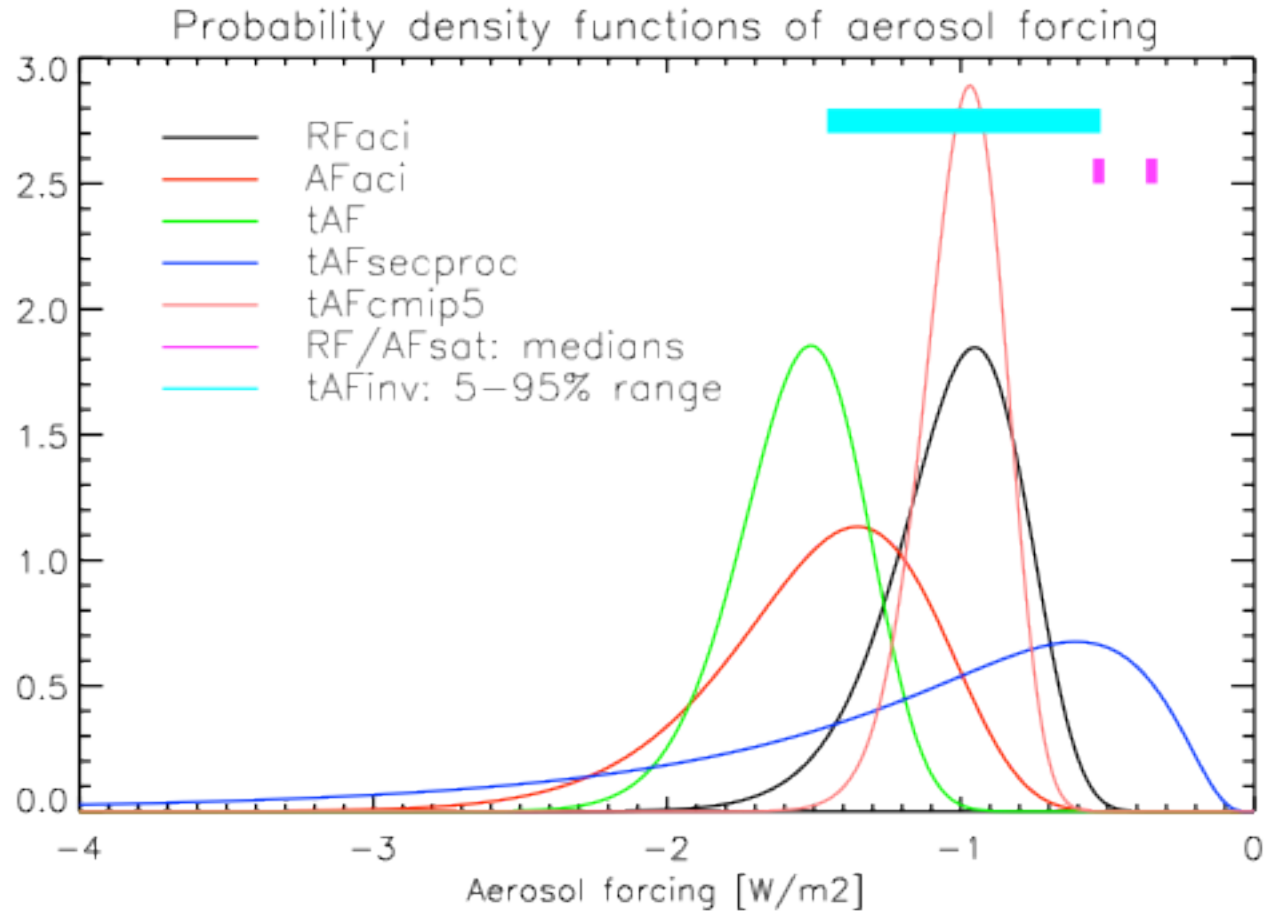
Y. Yun and J. Penner, University of Michigan

Y. Wang and X.Liu, PNNL

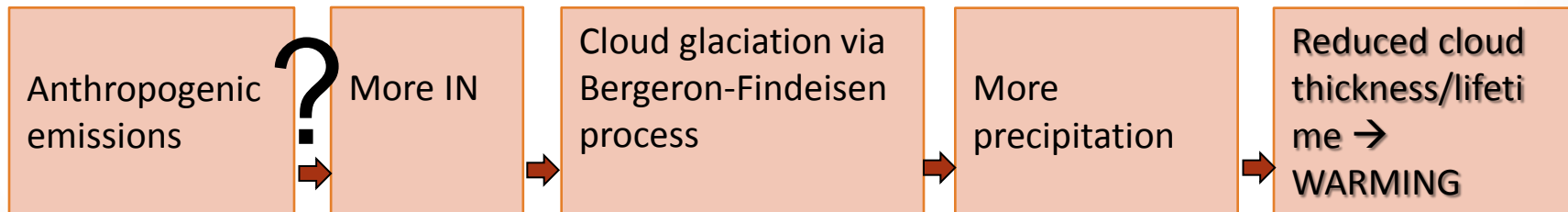
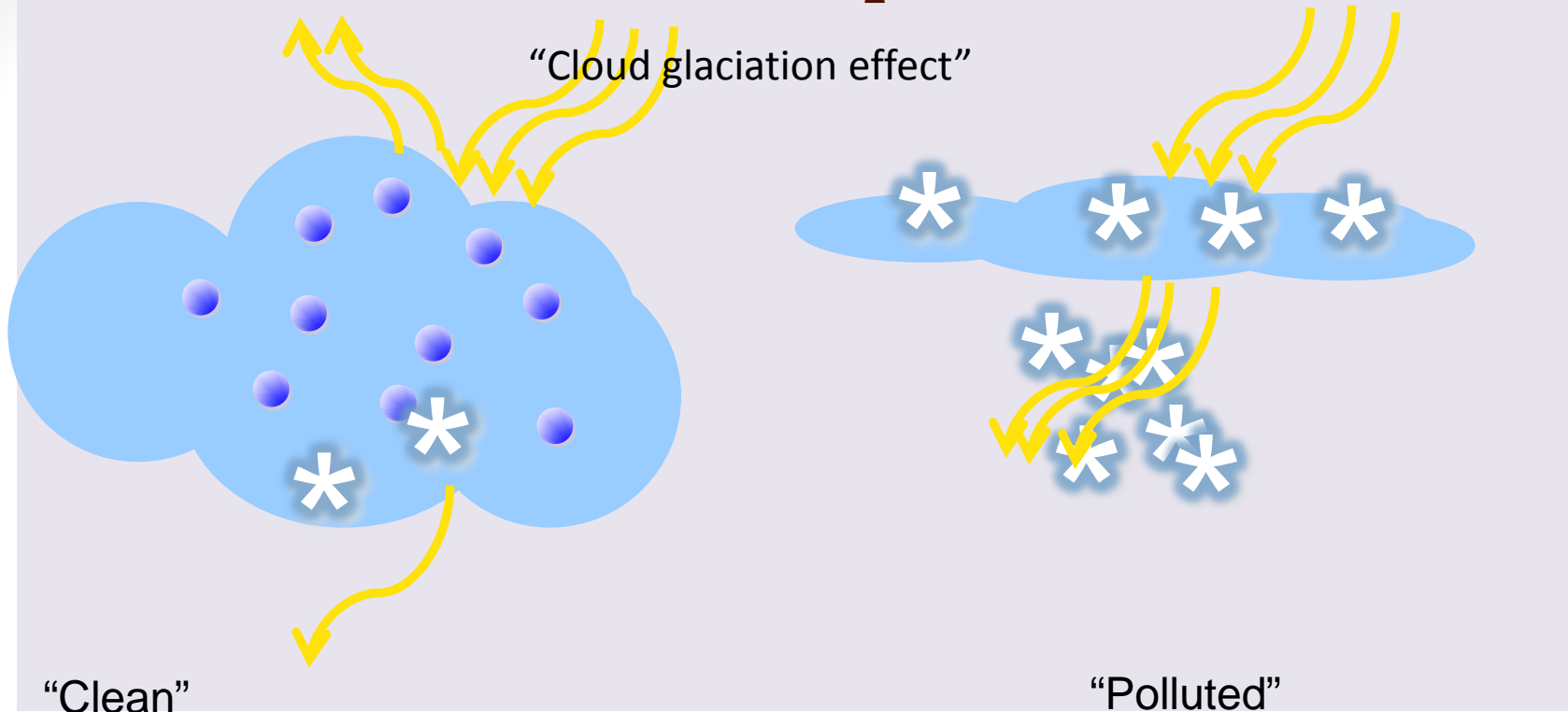
# Aerosol-cloud interactions, AR5



# Aerosol-cloud interactions, AR5



# Aerosol and mixed-phase clouds



## Ways that mixed-phase clouds could influence the total aerosol effect

- 1) Glaciation effect: Anthropogenic emissions = more IN (reversed lifetime effect)
- 2) De-activation effect: Coating of soluble material can “de-activate” IN = fewer IN
- 3) By limiting effects on liquid clouds: many liquid clouds, stronger effect

# Model Intercomparison on aerosol effects on liquid and mixed-phase cloud: Key questions

- Will a subset of models that explicitly represent mixed-phase cloud microphysics yield a lower aerosol adjusted forcing?
- How sensitive is the aerosol indirect effect to the choice of heterogeneous ice nucleation parameterization?
- How does the aerosol forcing change when IN concentrations are kept the same in PD and PI simulations?
- How do the simulated liquid cloud fractions compare to satellite observations

# Experimental set-up

- Experiment 1: Simulations using the DeMott et al. (PNAS, 2010) ice nuclei (IN) parameterization, in which the number concentration of IN active at cloud temperature  $T_k$  (in Kelvin) is given by:

$$n_{IN,T_k} = a(273.16 - T_k)^b (n_{aer,0.5})^{(c(273.16 - T_k) - d)}$$

where  $a$ ,  $b$ ,  $c$  and  $d$  are constants and  $n_{aer,0.5}$  is the concentration of insoluble aerosol particles larger than  $0.5\mu\text{m}$ .  $n_{aer,0.5}$  was prescribed and identical in PD and PI simulations.

- Experiment 2: Simulations using the DeMott et al. parameterization, but this time each model simulated its own concentration of insoluble particles larger than  $0.5\mu\text{m}$ .
- Experiment 3: Simulations in which the models all used their own internal treatment of heterogeneous freezing
- Each experiment was run twice: once with PD, and once with PI Emissions. Emissions were taken from Dentener et al. (Atm. Chem. Phys., 2006). Both direct and indirect effects were included.

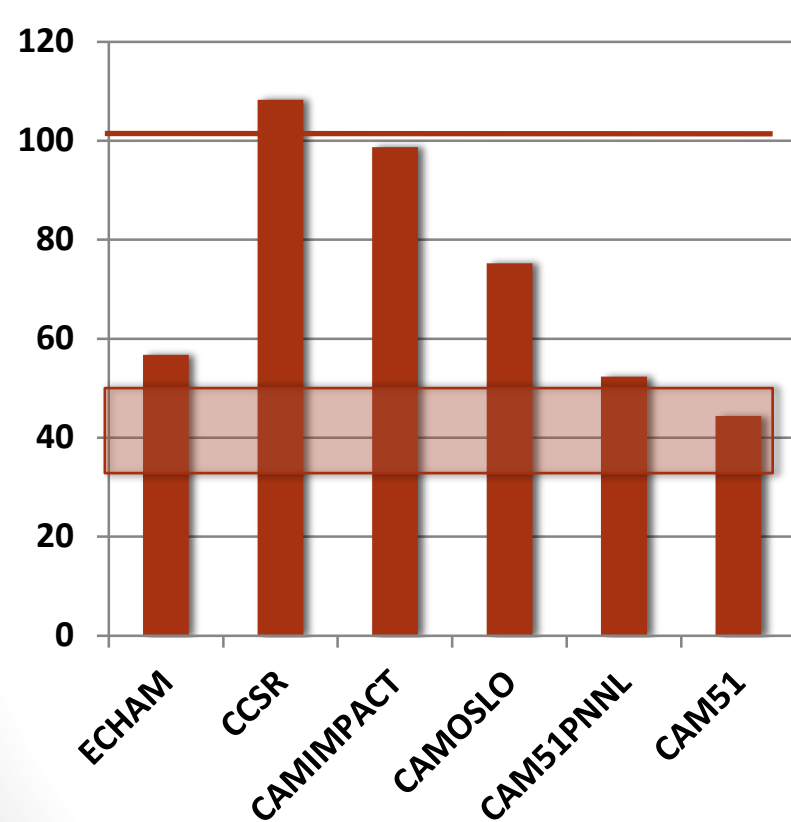
# Model intercomparison

- **Participating models:** ECHAM5-HAM, CAM-Oslo, CAM-Impact, MIROC-SPRINTARS (CCSR), CAM5.1, CAM5.1-PNNL
- **Requested output**
  - 2D fields: LWP, IWP, SWCF, LWCF, Net SW & LW radiation, (TOA and surface), column cloud droplet and ice crystal number concentrations, Cloud cover
  - 3D fields: Ice and water mixing ratios, Ice crystal and cloud droplet number concentrations, Ice crystal and cloud droplet effective radii, Cloud cover, Concentration of particles larger than 0.5 $\mu$ m (only for exp. 2), Number concentration of insoluble particles that are potential IN (only for Exp. 3), Total particle number concentration.
  - Additional output: Additional output on isotherms (-5°C, -10°C, -15°C, -20°C, -25°C, -30°C, -35°C) for comparisons with CALIOP (calculated only in grid-boxes in which: i) the temperature is that of the isotherm +/- 1°C ii) total (liquid & ice) water content exceeds 1.e-10 kg/kg and iii) there is no grid-box above that has an optical depth larger than 3).

# Liquid Water Path (LWP)

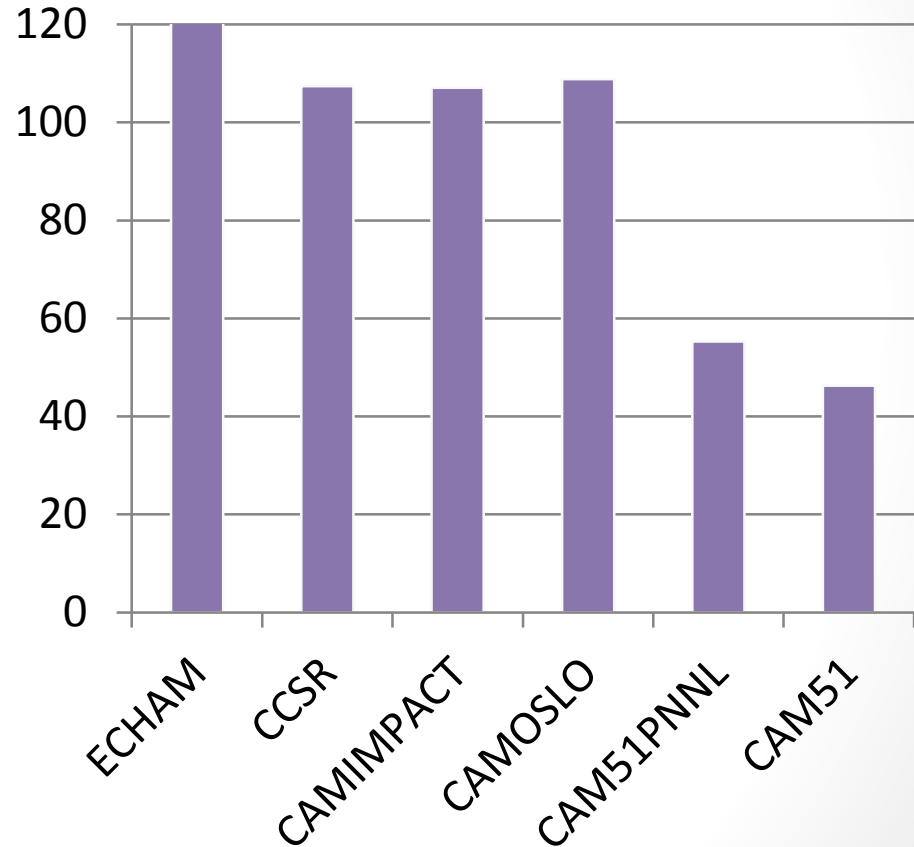
## E3

LWP (g/m<sup>2</sup>)



## E2

LWP (g/m<sup>2</sup>)



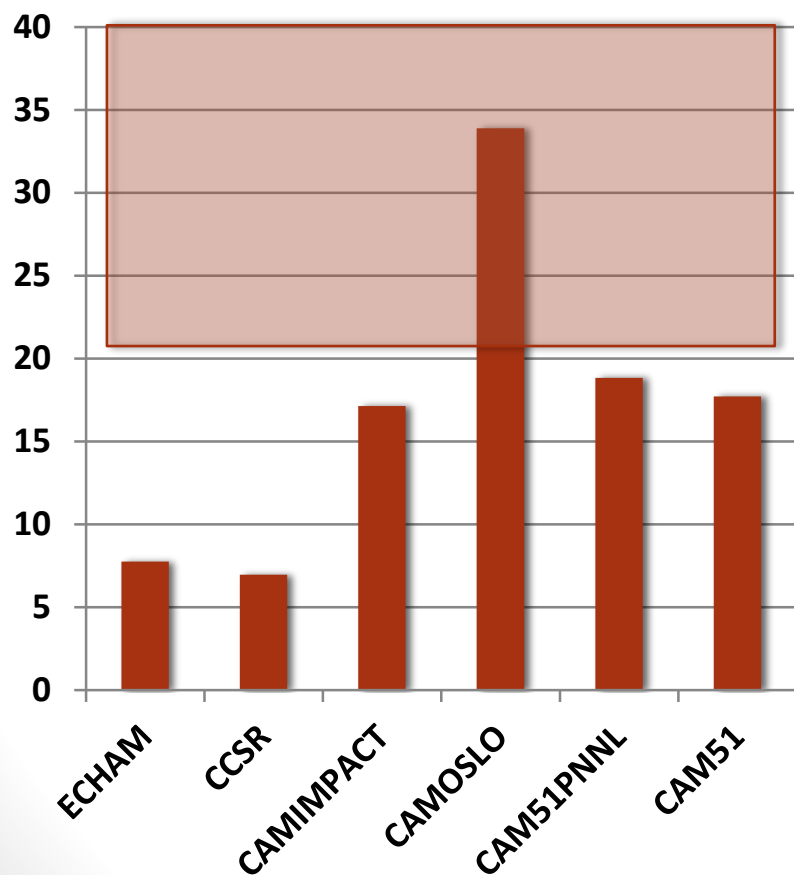
 A-train observations taken from Jiang et al. (2012)



# Ice Water Path (IWP)

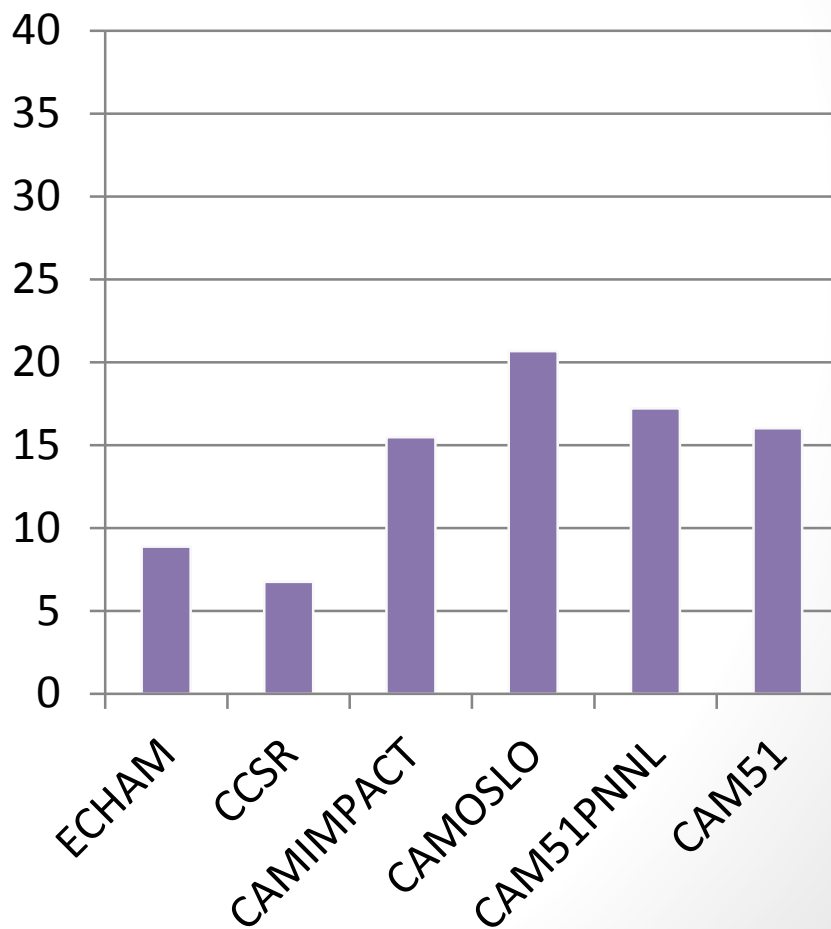
## E3

IWP (g/m<sup>2</sup>)



## E2

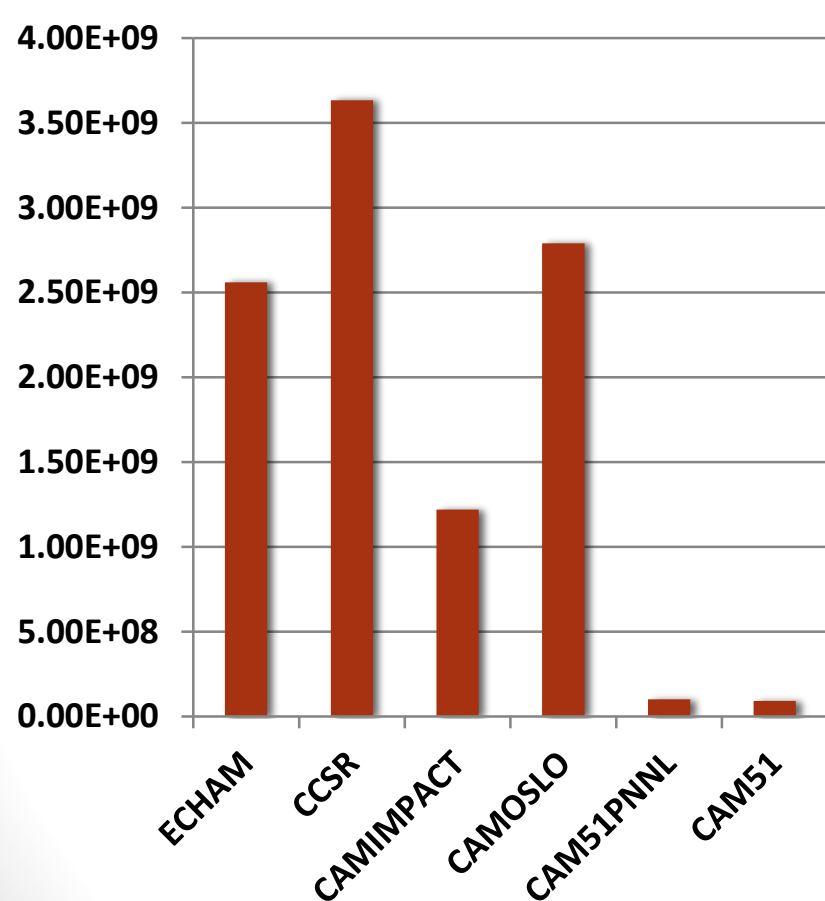
IWP (g/m<sup>2</sup>)



# Column ice crystal number

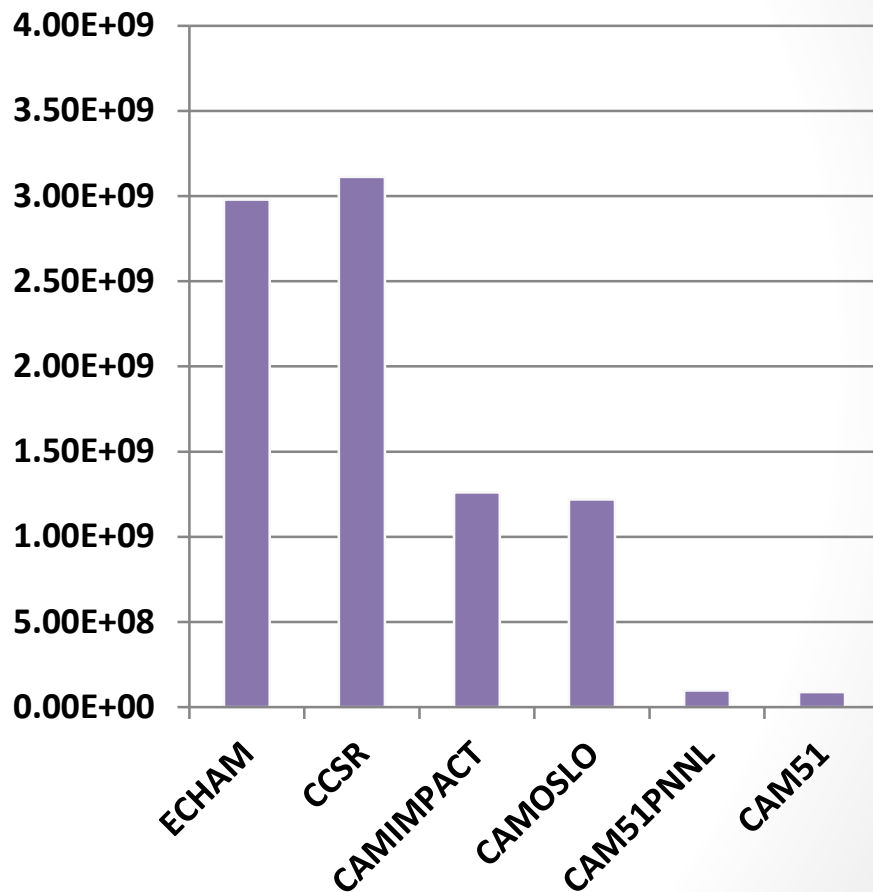
## E3

ICE NUMBER (#/m<sup>2</sup>)

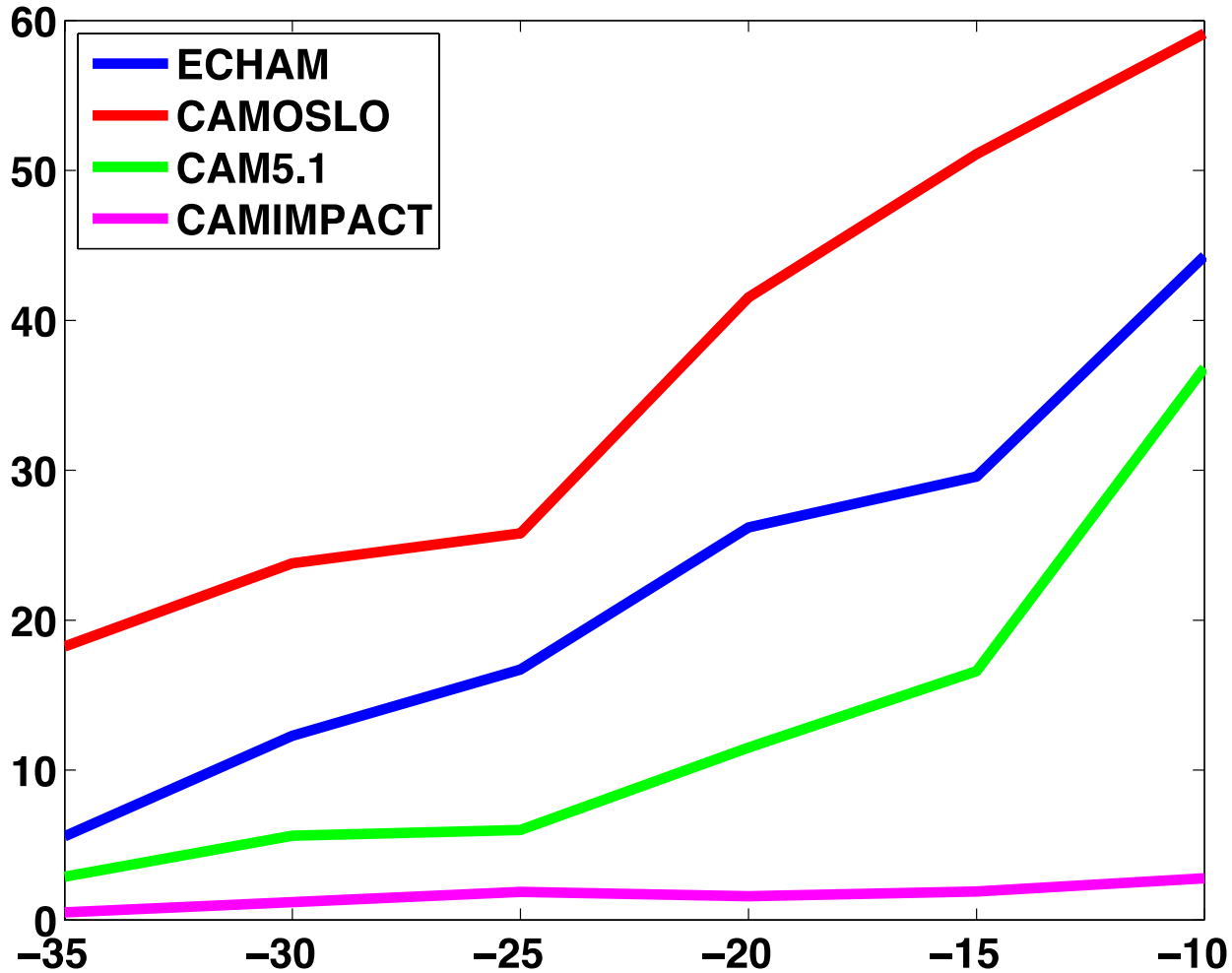


## E2

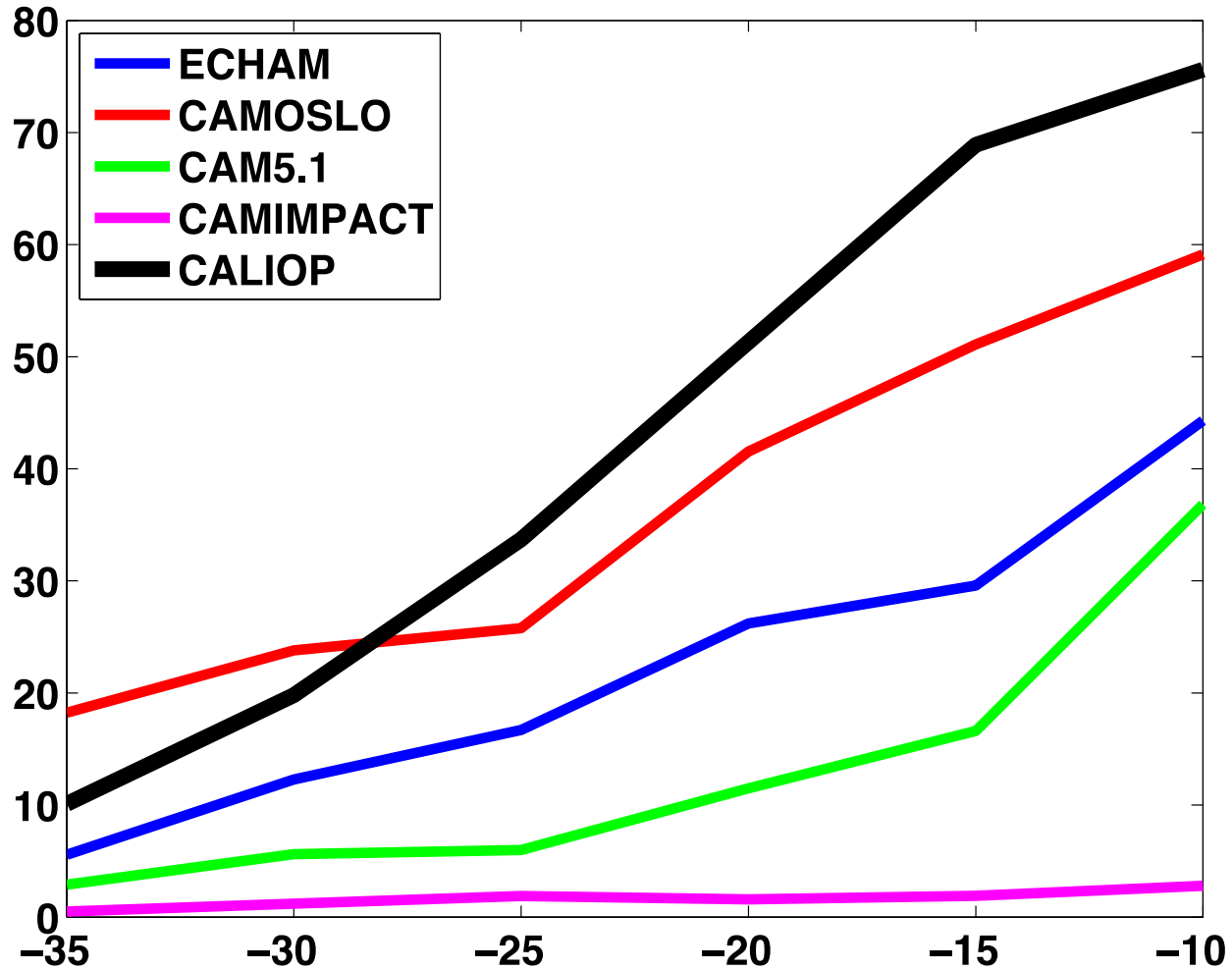
ICE NUMBER (#/m<sup>2</sup>)



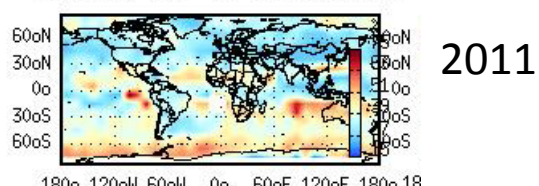
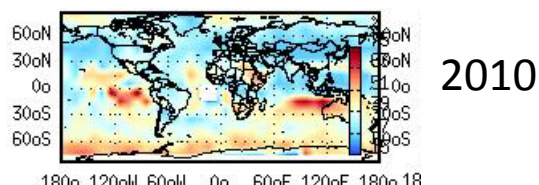
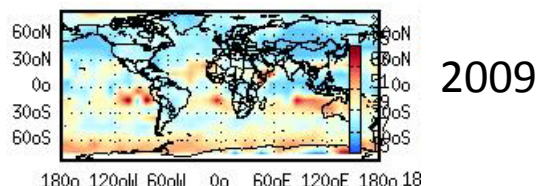
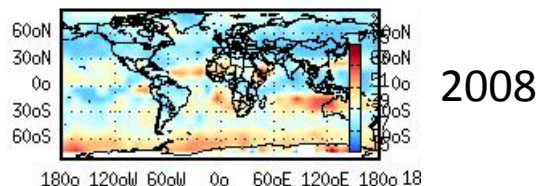
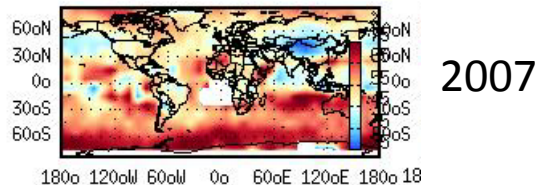
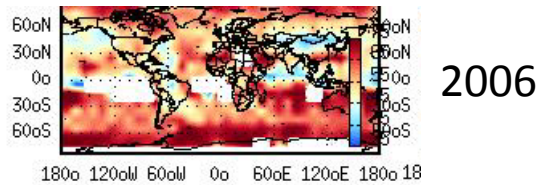
# Global Annual Mean Liquid Fraction



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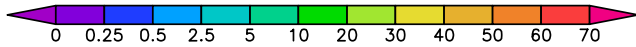
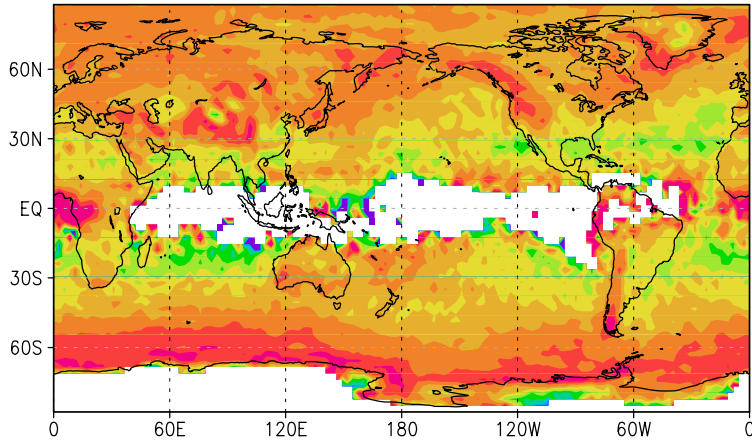
# Validating with CALIOP



- Dramatic change in supercooled liquid fraction (here shown for the  $-20^{\circ}\text{C}$  isotherm) after off-nadir tilt in November 2007.
- For the purpose of the intercomparison, only 2008-2012 data should be used.

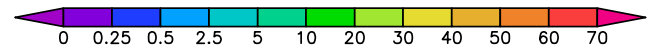
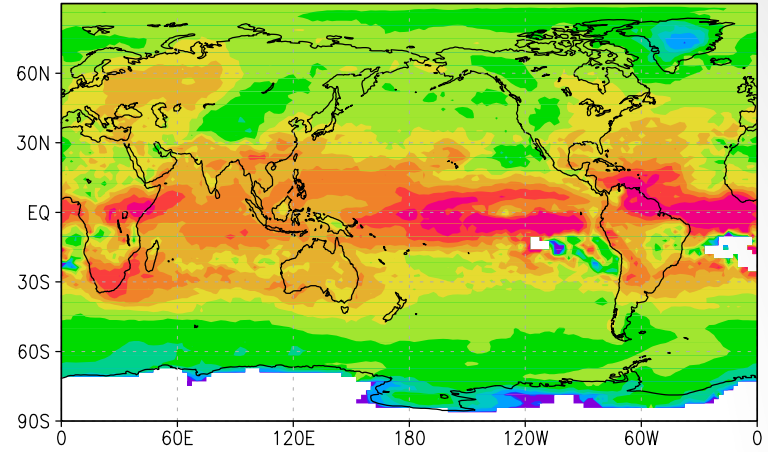
# Simulated supercooled liquid fractions at -10°C

ECHAM



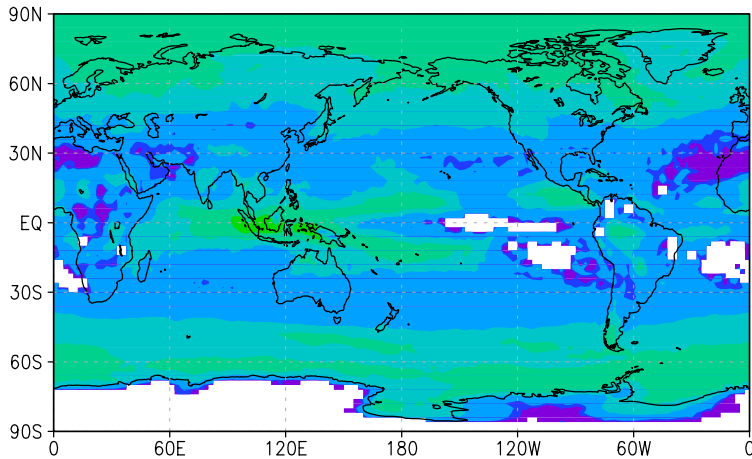
GrADS: COLA/IGES

CAM5.1



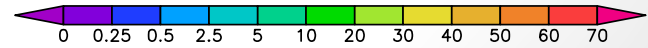
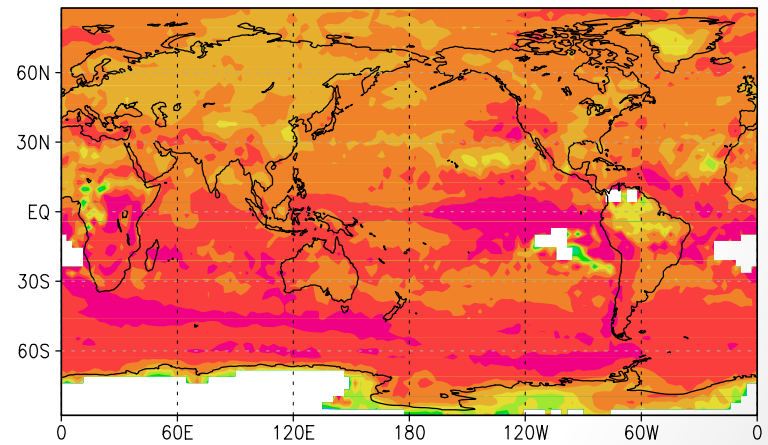
GrADS: COLA/IGES

CAMIMPACT



GrADS: COLA/IGES

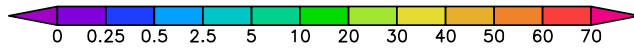
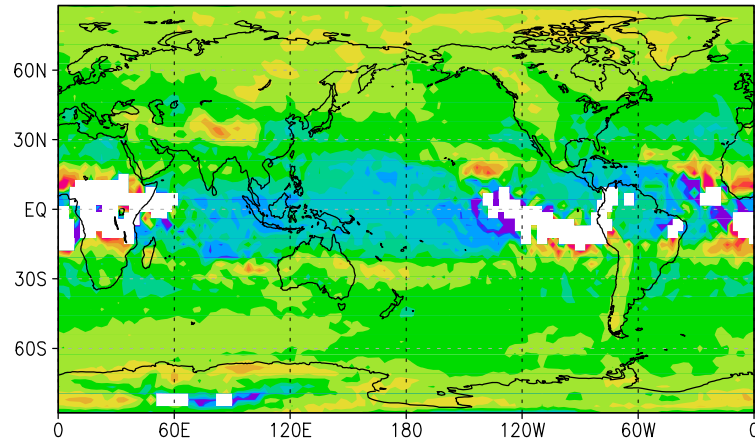
CAMOSLO



GrADS: COLA/IGES

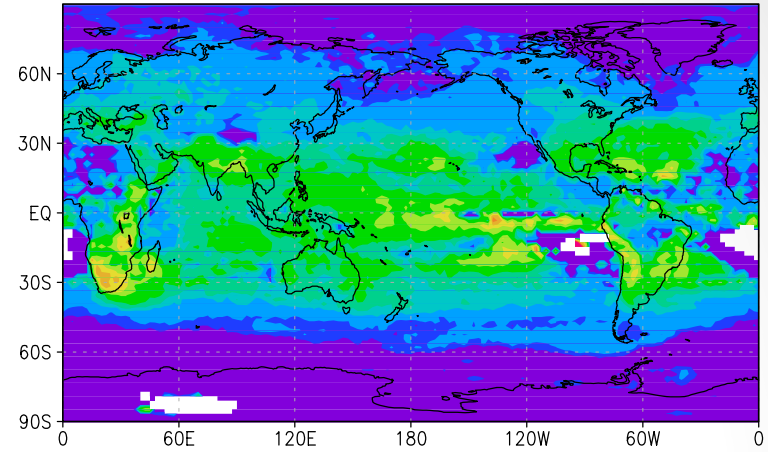
# Simulated supercooled liquid fractions at -25°C

ECHAM



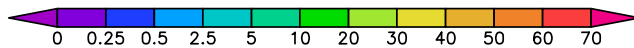
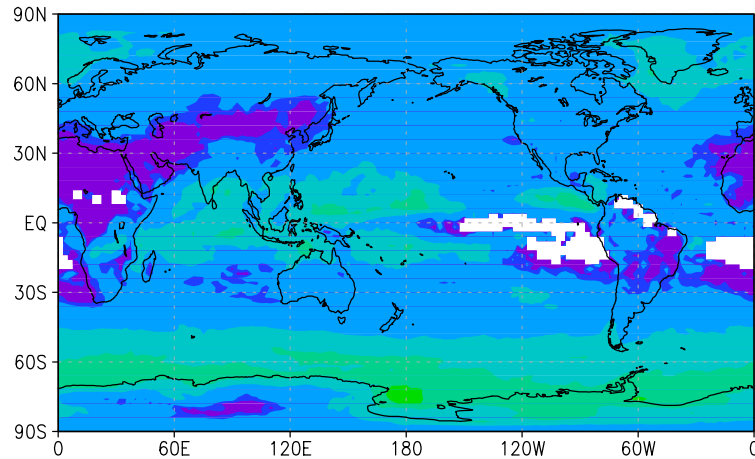
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CAM5.1



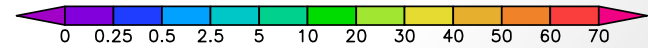
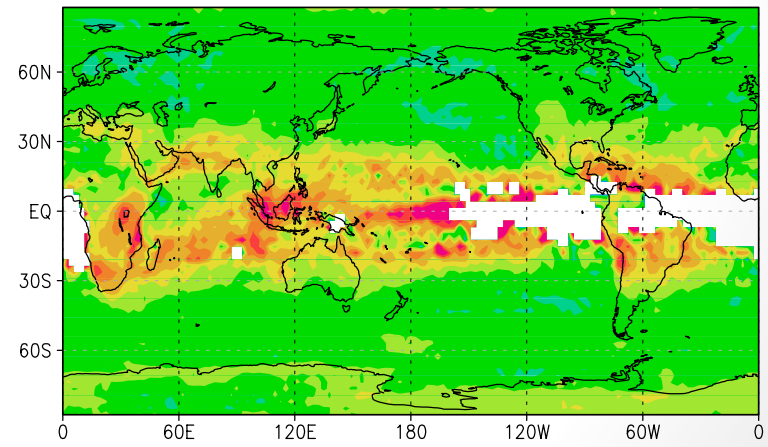
GrADS: COLA/IGES

CAMIMPACT



GrADS: COLA/IGES

CAMOSLO

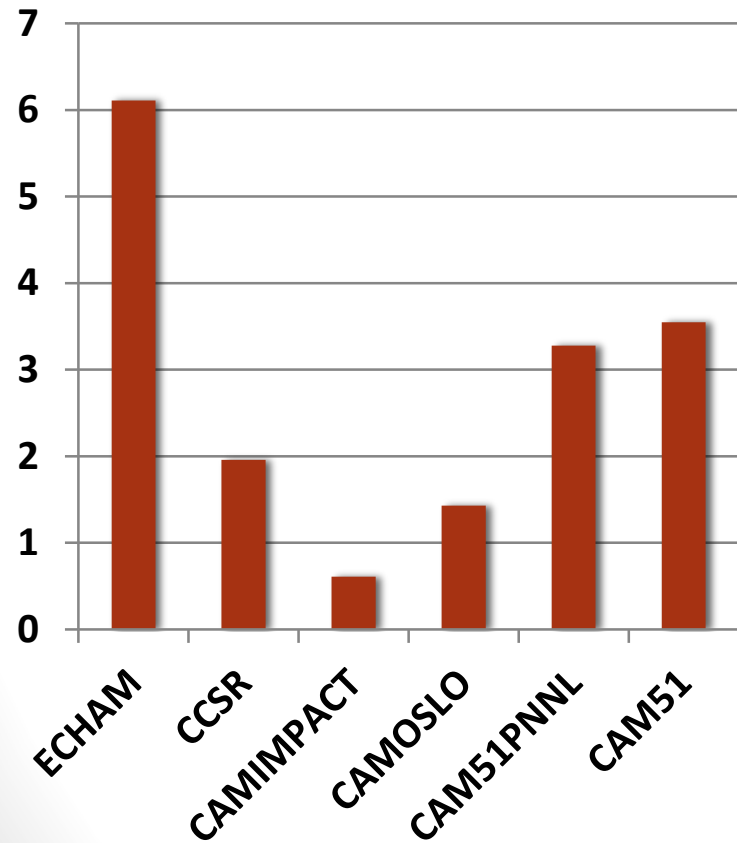


GrADS: COLA/IGES

# Change (PD-PI) in LWP

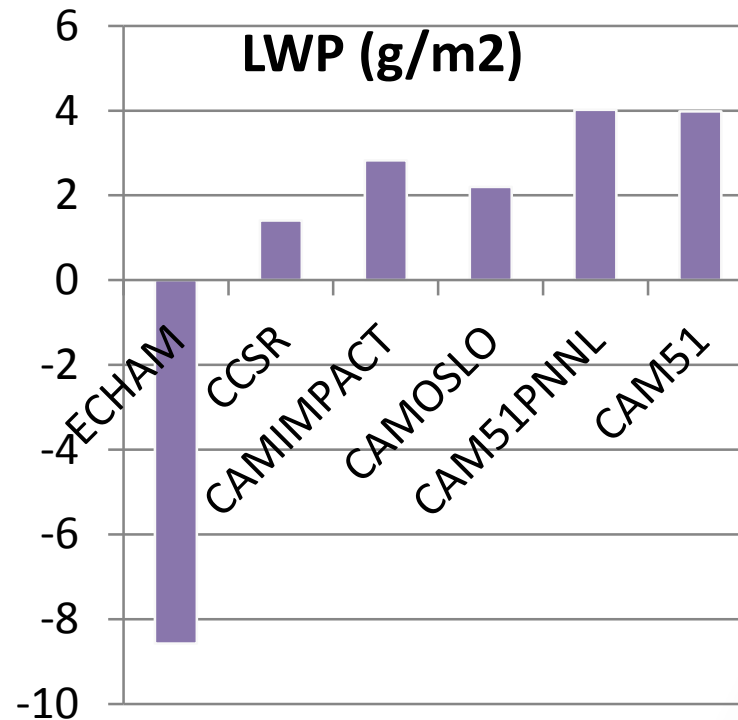
## E3

LWP (g/m<sup>2</sup>)



## E2

LWP (g/m<sup>2</sup>)

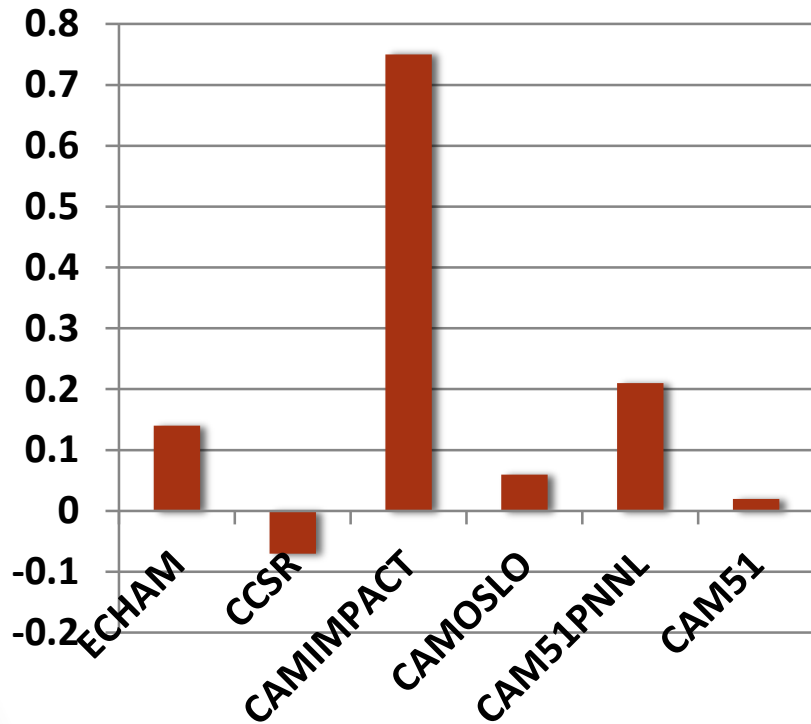




# Change (PD-PI) in Ice Water Path

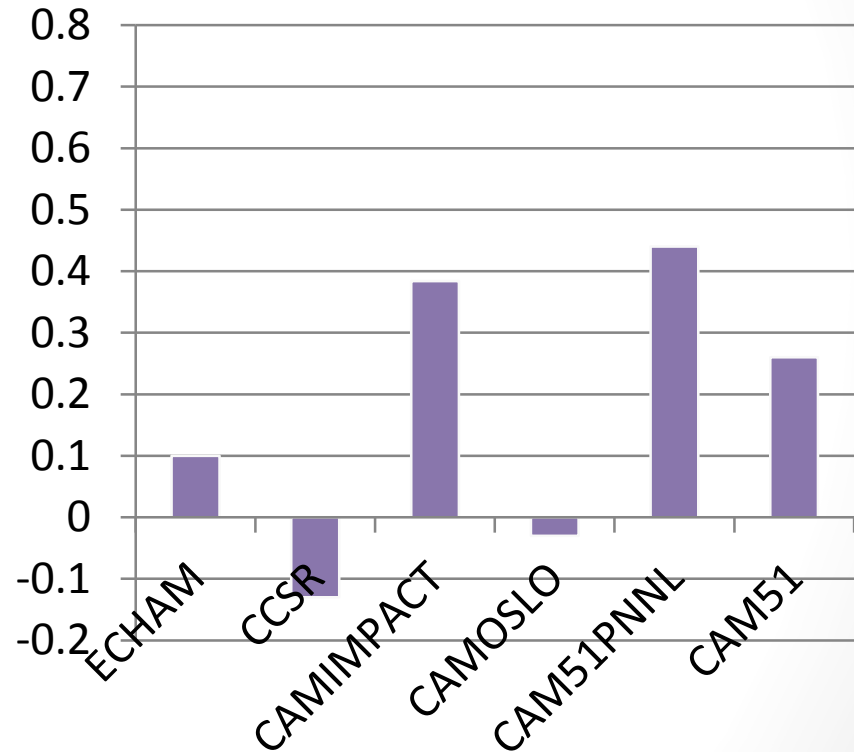
## E3

IWP (g/m<sup>2</sup>)



## E2

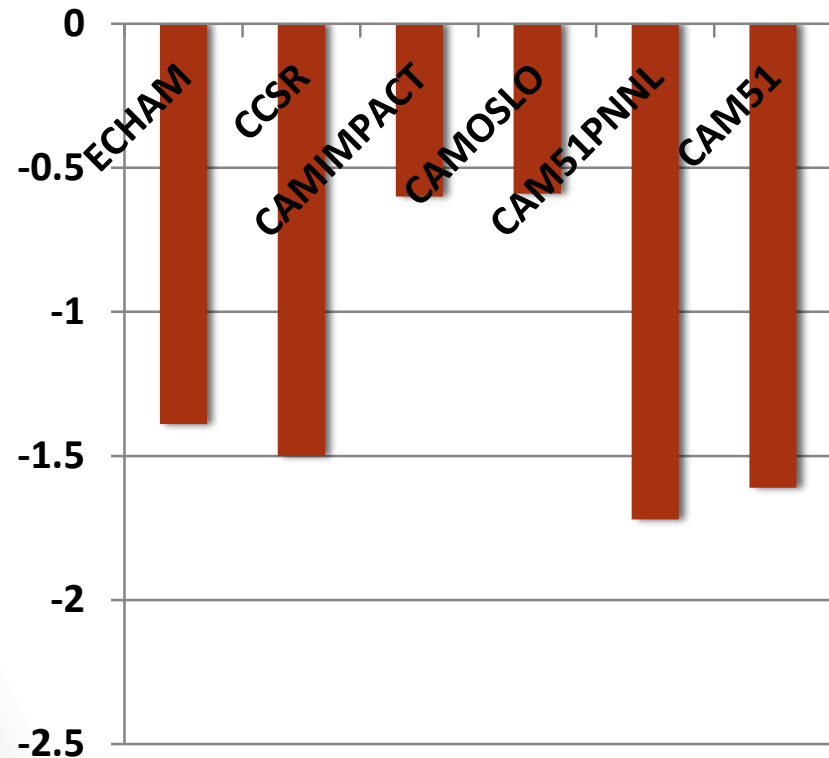
IWP (g/m<sup>2</sup>)



# Change (PD-PI) in SW cloud forcing

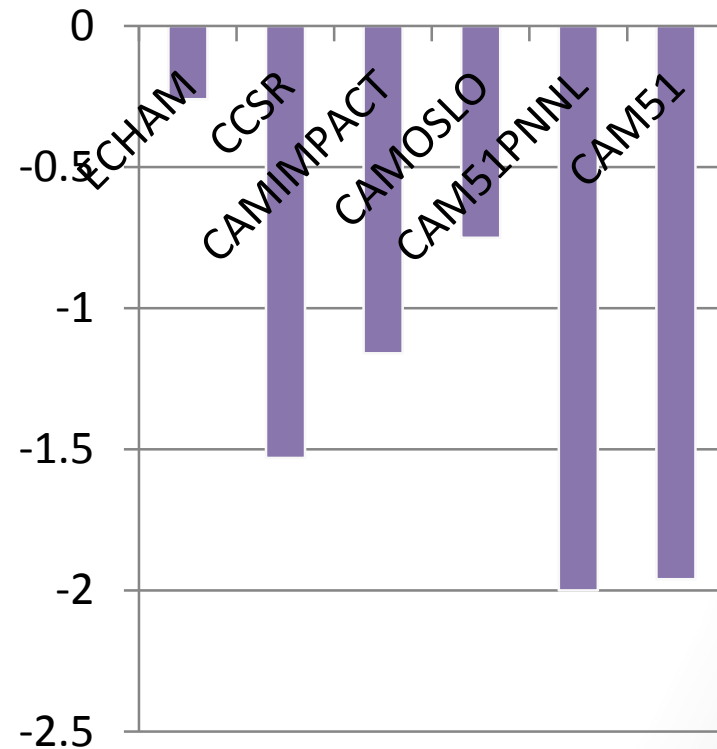
## E3

SWCF (W/m<sup>2</sup>)



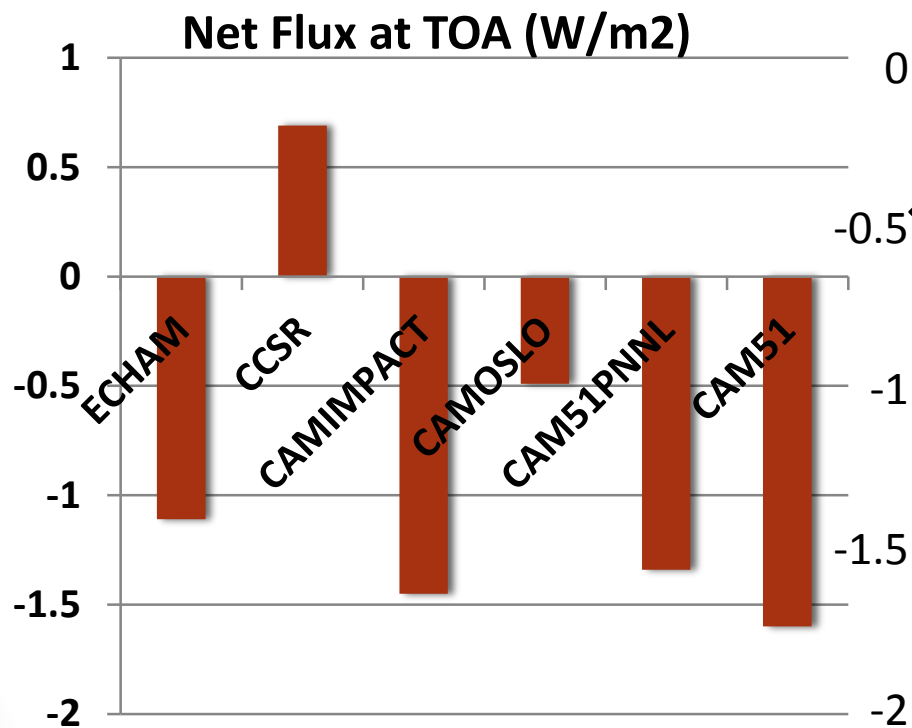
## E2

SWCF (W/m<sup>2</sup>)

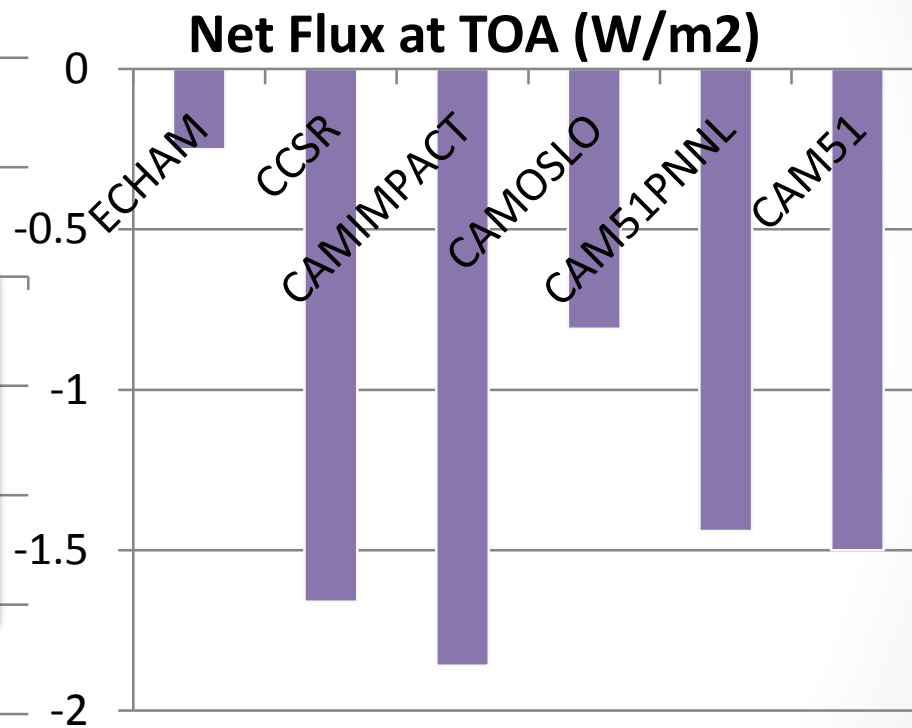


# Change (PD-PI) in net flux, TOA

## E3



## E2



# Preliminary intercomparison results

- Models show a wide spread in simulated IWP, LWP, ice crystals concentrations and supercooled liquid fractions.
- Comparison with satellite data suggests models liquid fractions are on the low side, but CALIOPs sensitivity to off-nadir tilt is a reminder that satellite retrievals have their own challenges.
- Simulated aerosol adjusted fluxes seem to be sensitive to the choice of ice nucleation parameterization (some more than others).
- The simulated total aerosol adjusted flux is lower in the 6 participating models than in the mean of estimates in the literature (compiled for AR5).